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Preliminary Draft

GEOCHEMICAL BUDGET FOR FILSON CREEK WATERSHED

by

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Note: This includes the first ^{7 parts} ~~three chapters~~ of the doctoral dissertation, "Geochemical Mass Balance of Filson Creek Watershed, Northeastern Minnesota."

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INTRODUCTION AND PURPOSE

The behavior of silicate minerals in the weathering environment and the effects of water-solute reactions on both the minerals and the resulting solution have been recently investigated for ground-water and surface-water systems. Through use of activity diagrams and an assumed equilibrium between solid and aqueous phases, Feth (1964) related ground-water quality of springs discharging from the Sierra Nevada batholith to the weathering products of feldspar. Garrels (1967) compared natural ground-water chemistry from major igneous rock types to calculated chemical quality based on theoretical weathering reactions of silicate minerals to kaolinite. Garrels and MacKenzie (1967) elaborated on this approach and stoichiometrically back-reacted the chemical quality of an average spring water reported by Feth (1964) with kaolinite to reproduce the original mineralogy of the bedrock.

These ground-water studies together have suggested that; most dissolved silica in ground water is derived from the weathering of silicate minerals other than quartz, reactions are so rapid that waters remain in near-equilibrium with one or more phases at all times, aluminum is chemically conserved, and that the chief buffer system is the CO_2 system (Bricker and Garrels, 1967).

Norton (1974) extended these studies to evaluate the surface water chemistry of the Rio Tanama system in Puerto Rico. Changes in river water quality were related to reactions between surface-water and rock-forming minerals in the watershed. Norton also calculated from activity-concentration plots a solubility product for montmorillonite which favorably compared with previously published data.

Bricker et al. (1968) and Cleaves et al. (1970) utilized a mass balance approach coupled with silicate weathering models for a small watershed in Maryland. From their geochemical model they were able to determine detailed input-output budgets for major ions and the relative rates of chemical and mechanical weathering in the watershed.

Other chemical budget studies of small watershed have generally used a paired-watershed approach to evaluate disturbances to ecosystems by various stresses, such as logging or forest fires, and have not evaluated cation losses from the watersheds with comparable geochemical detail. For example, the classic studies by Johnson et al. (1969) and Likens et al. (1967, 1969) qualitatively evaluated the reactions controlling the release of major cations from Hubbard Brook watershed by comparing ratios of yearly cation losses to the chemical composition of the weathered and unweathered bedrock. Other studies by Johnson and Swank (1973), Fredrickson (1970, 1972), and Wright (1974) calculated cation budgets for small watersheds but did not address the mechanisms of mineral

to account for net cation losses.

The cited watershed and ground water have not had to incorporate the kinetics of silicate weathering to geochemical models because aquifers or watershed soils and streambed materials contained silicates with similar stability. However, the rates of silicate weathering are different for the major silicate groups. Silicate susceptibility to weathering decreases in a series from olivine to quartz Goldich (1938), analogous to Bowen's Reaction Series and the Stunz (1941) classification of silicates from the least complicated structures, the neosilicates (e.g. olivine), to the most complicated structures, the tectosilicates (e.g. feldspar).

As opposed to earlier investigations, this dissertation extends geochemical modeling for a watershed which is underlain by a composite of igneous minerals including all major silicate groups. The results of this work will provide:

- 1). a means to quantitatively assess the relative rates of chemical weathering between major silicate groups under the same hydrological and geochemical conditions in the natural environment.
- 2). a means to determine the practical extent to which chemical mass balances and geochemical

models can be applied to a natural system considerably more complicated than those previously studied.

In addition, because the watershed includes a mineralized zone containing copper and nickel sulfide minerals, this study will integrate, if necessary, both sulfide oxidation as well as hydrolysis of the silicate minerals in the geochemical analysis.

LOCATION AND PHYSIOGRAPHIC DESCRIPTION OF THE WATERSHED

Filson Creek is located on the Superior Upland Province of Minnesota and is approximately 8 miles southeast of Ely (fig. 1). It was chosen for study because:

- 1). Filson Creek is a small stream with a continuous U.S. Geological Survey stream gauge near its mouth.
- 2). Previous soils, water chemistry, geological and vegetational studies are available for the watershed.
- 3). All waters draining from the basin originate as precipitation on the watershed.
- 4). The watershed is relatively undisturbed by man.
- 5). The watershed is underlain by glacial till composed predominantly of a mixture of igneous minerals derived from the underlying mafic Duluth Complex and nearby acid lithologies of the Giants Range Granite.
- 6). The watershed is predominantly forested.

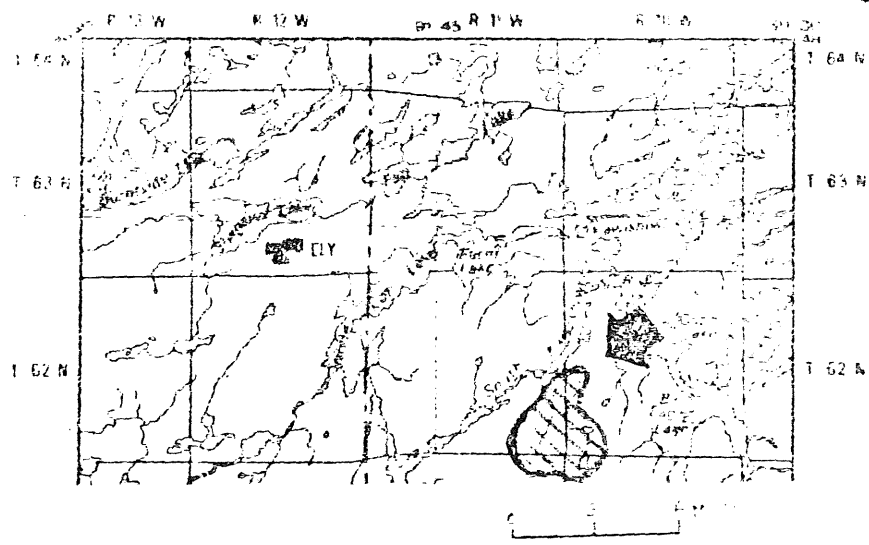
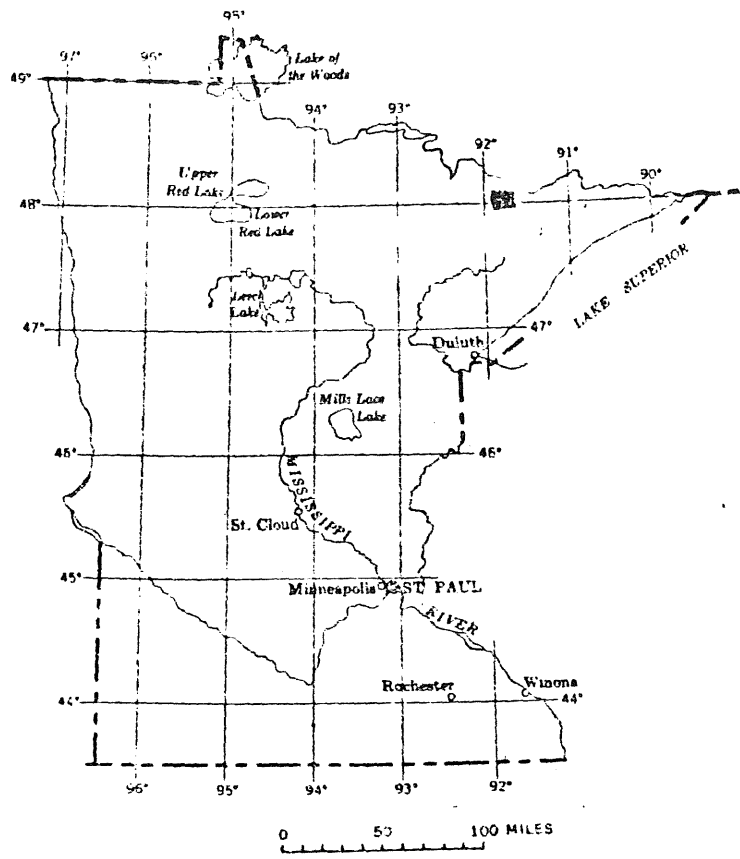


Figure 1 Location of Filson Creek Watershed

Filson Creek drains westward to the Kawishiwi River and has two tributaries, designated North Filson and South Filson for this study (fig. 2), which join approximately one-tenth of a mile upstream from the U.S. Geological Survey stream gauge on Spruce Road.

Total relief in the watershed is 317 ± 10 feet, with elevations ranging from greater than 1,760 feet east of Bogberry Lake in the eastern part of the basin to less than 1,443 feet near the mouth of the stream.

The general topographic grain trends northeast-southwest and reflects bedrock ridges in the Duluth Complex as well as the general direction of Wisconsinian glaciation. South and southwest of Bogberry Lake, this general trend is broken by an east-west bedrock ridge which partly outlines the southern boundary of the watershed.

Between five and seven percent of the watershed is exposed bedrock.

South Filson Tributary

South Filson drains 2.42 square miles of fens and upland forest bordered on the west and east by bedrock ridges mantled by a thin veneer of till. The southern

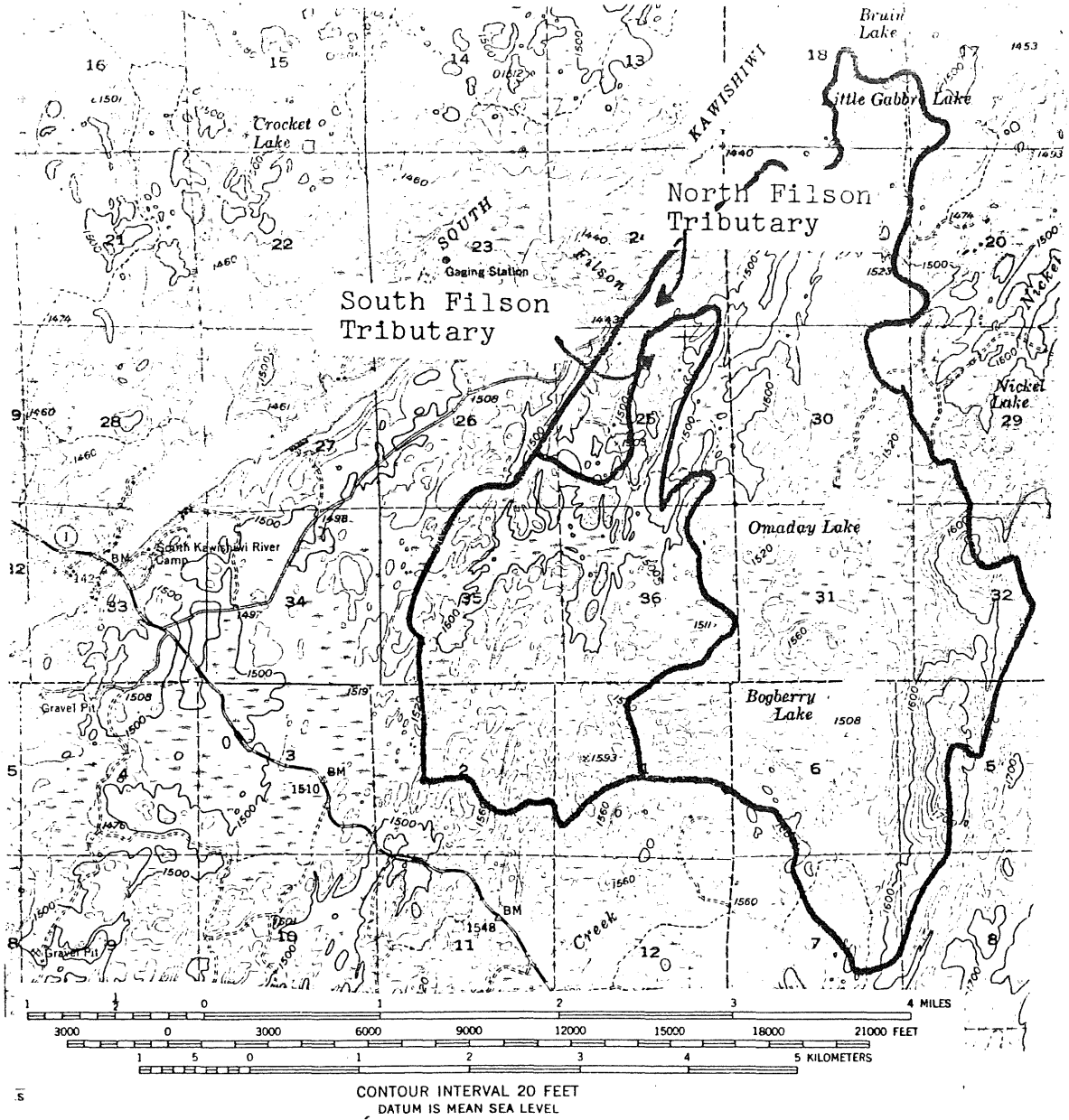


Figure 2 Topography of the Filson Creek Watershed

boundary of the watershed is indistinct and estimated from surface topography in the wetlands. South Filson is about 3.3 miles long and meanders in a linear fen which ranges from 500 to 1,250 feet in width.

Small rapids are located directly upstream from its confluence with North Filson and at other locations where bedrock ridges increase local gradients in the channel. The general gradient for South Filson watershed is 20 feet per mile.

North Filson Tributary

North Filson drains a 7.34 square mile area including Bogberry Lake (94 acres) and Omaday Lake (34 acres). The eastern and southern boundaries are well defined by forested topographic ridges having relief from 140 to 250 feet. The western margin consists of en'echelon bedrock ridges having relief from 20 to 60 feet.

Bogberry and Omaday Lakes range in depth from three to five feet, depending on the season, and are almost completely surrounded by wetlands. Bogberry Lake drains northward to Omaday Lake through a perennial stream.

North Filson has an average gradient of 14 feet per mile, one-third less than that of South Filson. As with South Filson, small riffles and rapids are located at nick points where the channel crosses bedrock ridges or leaves beaver dams. One particularly long rapids extends for about 600 feet upstream from the intersection of the creek with Spruce Road in T.62 N., R.11 W., section 9.

CLIMATE

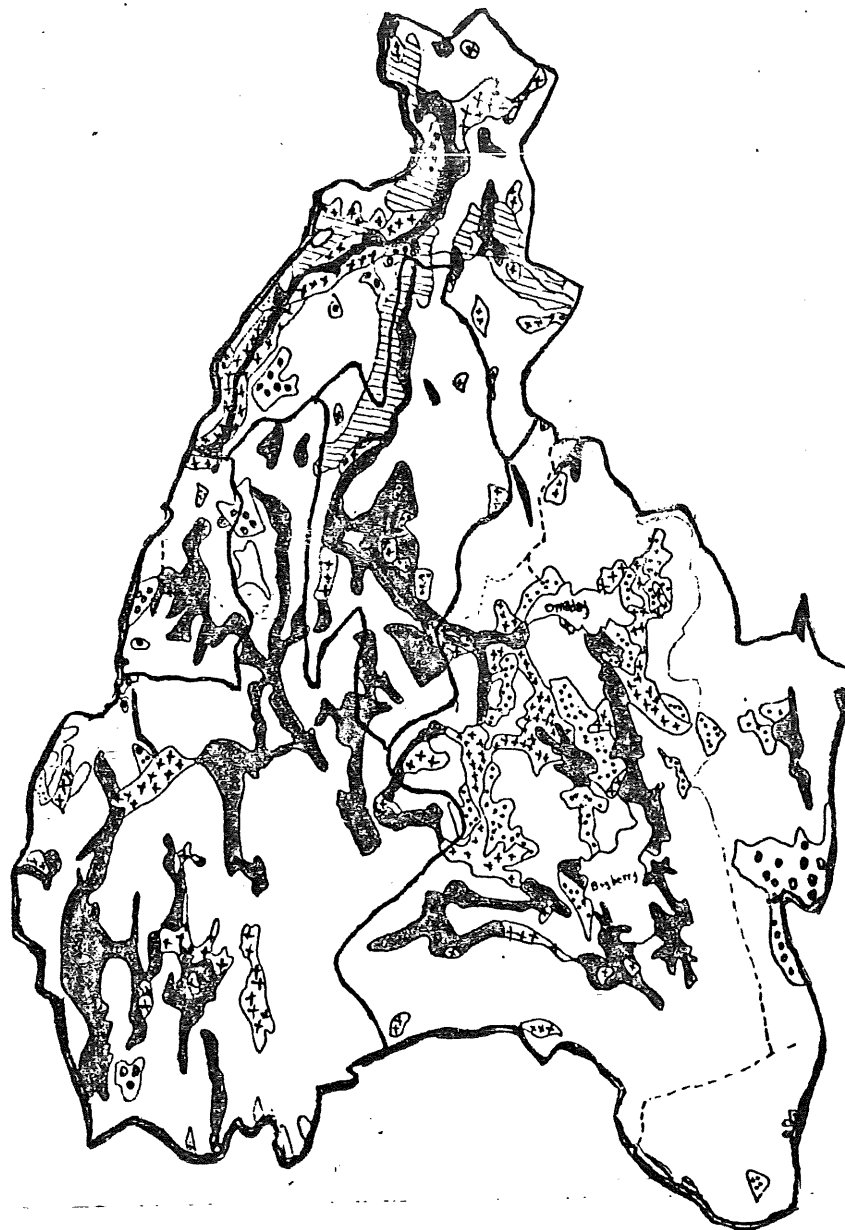
Filson Creek watershed has a mid-continental climate with long, cold winters and warm, short summers. The mean annual precipitation is about 27 inches, with monthly average temperatures ranging from 14°F in January to 62°F in July (Ahlgren, 1969). Precipitation occurs in the form of brief storm events of varying intensity which occur most frequently during the early spring and fall.

VEGETATION

During the past 50 years, much of the watershed has experienced various degrees of logging. Much of the original post-glacial climax vegetation of Jack Pine and White Pine (Marschner and others, 1974) has been selectively cut. With the exception of parts of the high ridges marking the eastern boundary of the watershed, large scale clear-cutting and rock-raking techniques have not taken place. Those areas which were clear-cut are now generally covered with a secondary succession of aspen, birch, mixed coniferous species and various shrubs.

The vegetation in the watershed is composed of several distinct natural forest types, as well as some areas covered by planted species (fig. 3). About 75 percent of the watershed is classified as a mixed upland forest, 15 percent as fens and Black Spruce bogs, and the remaining 10 percent as natural or planted stands of Jack, Norway and Red Pine (*Regional Camp Study Site*, 1976).

The mixed upland forest consists of an assemblage of aspen, birch, fir and other conifers. The dominant species along the margins of the watershed are aspen and birch, ranging from 30 to 70 feet in height, while the more mixed



EXPLANATION

- Mixed Upland Forest
- Fen
- Black Spruce
- Natural Pine Stands
- Pine Plantations

After Sather (1977),
 Unpublished map, Regional
 C-11 Study Staff and
 Unpublished Forest Survey map,
 Vancouver, July, 1976.

Figure 3 Vegetation in the Filson Creek Watershed

assemblages including conifers are found generally in the central part of the watershed where there are wetlands and lakes. Two small stands of Jack Pine occur on the north side and about one-half mile due south of Omaday Lake. In the extreme northern part of the watershed are tracts of planted Red and Jack Pine. The northern reach of North Filson meanders in a spruce bog for about one and one-half miles before it enters a reed-sedge bog near the mouth of Filson South.

It is assumed from the vegetation that the biomass in the watershed is near dynamic equilibrium. This implies that biomass on the average is being replaced at the same rate at which it is being dispersed, Odum (1975). Evidence that suggests such an equilibrium operating in Filson Creek Watershed is similar to that described by Cleaves et al. (1970):








- 1). Trees of all size classes from saplings to stands up to 70 feet high occur in the watershed.
- 2). Deadfalls are present.
- 3). Major clear-cutting disturbances have not occurred during the past twenty years.
- 4). There is a general lack of heavy first order undergrowth except around wetland areas.

BEDROCK GEOLOGY

The bedrock geology (fig. 4) of the watershed mainly consists of the mafic lithologies of the middle pre-cambrian Duluth Complex. The northern part of the watershed is underlain by the eastern edge of the Giants Range Granite as well as xenoliths of hornfels and Biwabik Iron-formation found within the Duluth Complex.

The eastern part of the Giants Range Granite is composed mainly of medium to coarsely crystalline, porphyritic, hornblende rich, plutonic rocks that range in composition from adamellite to diorite, but dominantly are adamellite and granodiorite. (Sims and Viswanathan, 1972). Green (1970) divided the eastern part of the Giants Range Granite into the Farm Lake and Clear Lake facies. Only the Farm Lake facies is found in the watershed, and consists mainly of porphyritic, hornblende adamellite and monzonite. The rock typically is hypidiomorphic-granular in texture, having microcline phenocrysts ranging up to two centimeters in cross section. Grading with the porphyritic rocks are minor amounts of non-porphyritic varieties.

EXPLANATION

-  Troctolite
-  Anorthosite
-  Anorthositic gabbro
-  Hornfels and Biwabik Iron-Formation
-  Mineralized contact zone
-  Giants Range Granite
-  Lake

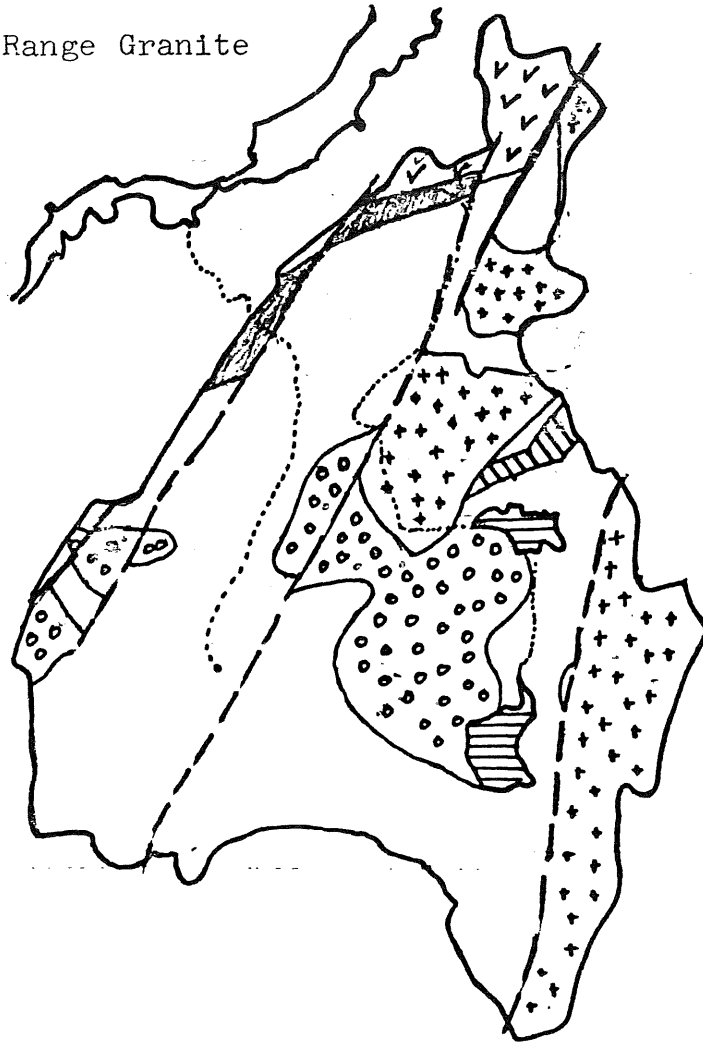


Figure 4 Bedrock Geology of the Filson Creek Watershed

The quartz content ranges from 15 to 25 percent, while the major mafic mineral, green hornblende, composes between 6 and 10 percent of the rock. The plagioclase type is oligoclase which ranges from An_{16} to An_{24} . In the field, oligoclase appears cloudy due to kaolinization which occurred prior to the intrusion of the Duluth Complex (Green, 1970).

The contact between the Duluth Complex and the Giants Range Granite along the northern margin of the watershed consists of a mineralized zone of gossen which is one quarter to one half mile wide. Mineralogy of the rock is transitional between the amphibolite, the hornblende-hornfels and the pyroxene-hornfels metamorphic facies (Green, 1970), and includes copper and nickel sulfide minerals which have potential economic value.

The Duluth Complex underlying approximately 90 percent of the watershed was first mapped by Green et al. (1966), with subsequent modifications by ^{Murray} Cooper (1978).

The center and eastern parts of the watershed are underlain by lithologies ranging from poikilitic augite troctolite to anorthosite, with plagioclase (An_{60} to An_{65}) content ranging from 65 to 90 percent, olivine (Fa_{37-40})

content from less than 10 to 20 percent, and the remaining percentage composed of interstitial iron oxides and augite.

The eastern margin of the watershed is underlain by anorthositic gabbro with plagioclase (An_{60-88}) comprising between 75 and 95 percent of the rock. The remaining percent consists of interstitial augite and hypersthene.

Due north of Omaday Lake, the bedrock includes an elongated xenolith of metamorphosed Biwabik Iron-formation and hornfels.

Average modes in volume percent of the major rock types discussed are given in Table 1. Table 2 gives the percentages of the watershed underlain by major rock types.

Table 1 Average modes in volume percent of major rock types in
the Filson Creek Watershed

	Troctolite ¹ (33)	Anorthosite ¹ (9)	Anorthositic Gabbro ¹ (35)	Giants Range Granite ² (3)
Plagioclase	71.2	95.2	78.9	49.3
Augite	4.5	0.9	8.9	---
Orthopyroxene	0.6	0.6	1.1	---
Olivine	19.3	1.3	3.2	---
Opakes	2.0	1.3	2.9	0.2
Biotite	0.6	tr	2.5	2.0
Hornblende	---	---	---	8.7
Quartz	---	---	---	11.7
K-spar	---	---	---	23.7

¹
from Phinney, 1972

²
from Sims and Viswanathan, 1972

Numbers after each rock type indicate number of thin sections used to calculate average mode

Table 2 Percentage of Filson Creek Watershed
Underlain by Major Bedrock Type

<u>Bedrock Type</u>	<u>Percentage of Watershed</u>
Troctolite	56.7
Anorthositic Gabbro	20.3
Anorthosite	12.9
Giants Range Granite	3.5
Contact Zone	2.7
Hornfels	1.4

GLACIAL HISTORY

Northeastern Minnesota in the vicinity of Filson Creek experiences two advances from the north and northeast by the Rainy Lobe of the Laurential Ice Sheet during Wisconsin time (Wright, 1972).

The first advance, termed the St. Croix phase, deposited the Toimi Drumlin Field and St. Croix Moraine south of the study area. The second advance, termed the Autumba phase, deposited thin ground moraine and minor ice contact deposits in the watershed. The ground moraine consists of bouldery, sandy till (Stark, 1978; Winter and others, 1973; Olcott and Siegel, 1978).

Numerous test holes by the author, (Prettyman, 1976 ; Olcott and Siegel, 1978; and Stark, 1977) indicate that the bouldery till is the major drift type in the watershed. Glaciofluvial sand and gravel constitute no more than 10 percent of the surficial deposits.

Omaday Lake is underlain by ^{about 22} feet of organic sediment, which covers sand and gravel of indeterminate thickness. The sand and gravel form a small beach on the northern side of the lake and may underlie wetlands near the mouth. U.S. Geological Survey test holes in and around Filson Creek (Olcott and Siegel, 1978) also indicate that wetland areas can be underlain by sand and gravel.

The total thickness of surficial materials, including postglacial wetland deposits, ranges from less than one foot over bedrock ridges to up to fifty feet in wetlands.

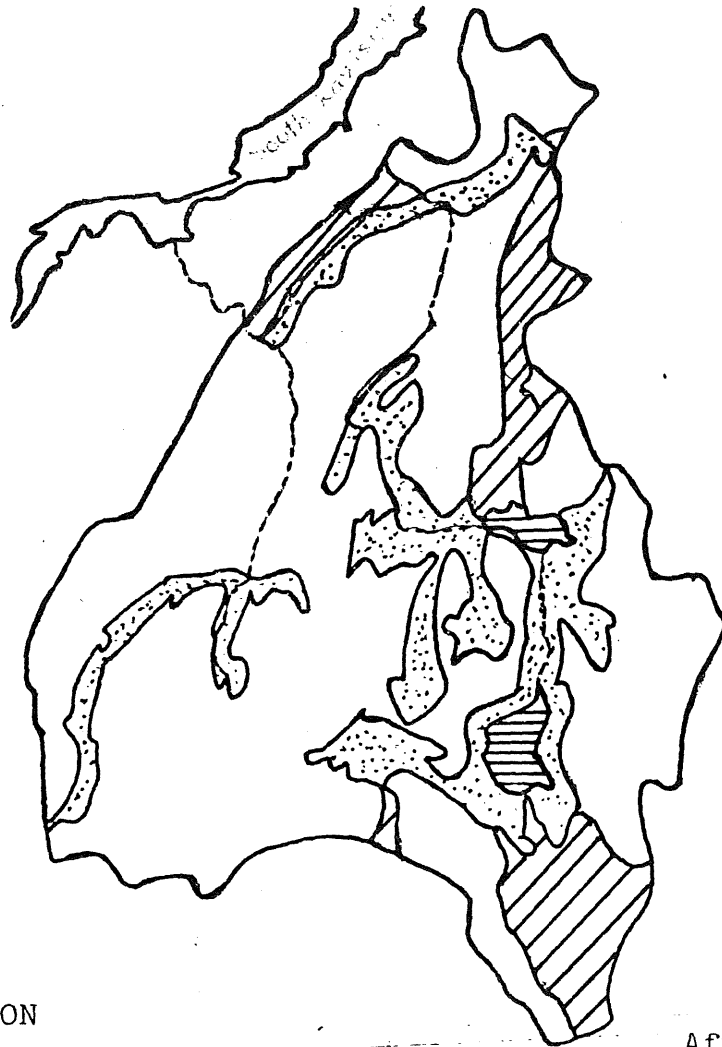
The thickness of the bouldery till is generally less than three feet thick over much of the watershed (Prettyman, 1976 ; Stark, 1977; Olcott and Siegel, 1978). It ranges, however, from only a few inches to an estimated 15 feet thick near the contact zone between the Duluth Complex and the Giants Range Granite. Till on the western flanks of the bedrock ridges bordering the eastern side of the watershed is as much as 5 feet thick.

SOILS DESCRIPTION





Soils in the watershed (fig. 5) are a mixture of immature, shallow and gravelly loams which occur over and adjacent to bedrock ridges. Peat and muck occur in the wetlands (Prettyman, 1976). Soil development is poor, with parent till generally less than one foot below the A-horizon (Grigal, personal communication, 1976).

The A-horizon consists of up to three inches of fine, organic-rich material containing few rock fragments and ranging in color from black to dark red. The poorly developed B-horizon ranges in color from yellow-brown to dark brown like its parent till and contains coarse rock fragments and boulders. In work along the western margin of the area, Alminas (1975) reported a clay-rich layer at the base of the B-horizon where till thickness exceeded two feet. The discontinuous clay layer, when found, is yellow-olive to yellow-gray.

Large clasts in the till reflect local source areas. For example, stone counts of over one hundred clasts of greater than pebble size indicate that underlying and nearby bedrock types constitute most of the rock fragments (fig. 6).



EXPLANATION

-  Bog
-  Sandy Loam, ≥ 40 " thick
-  Sandy Loam, < 40 " thick
-  Lake

After Prettyman(1976)

Figure 5 Soils in the Filson Creek Watershed

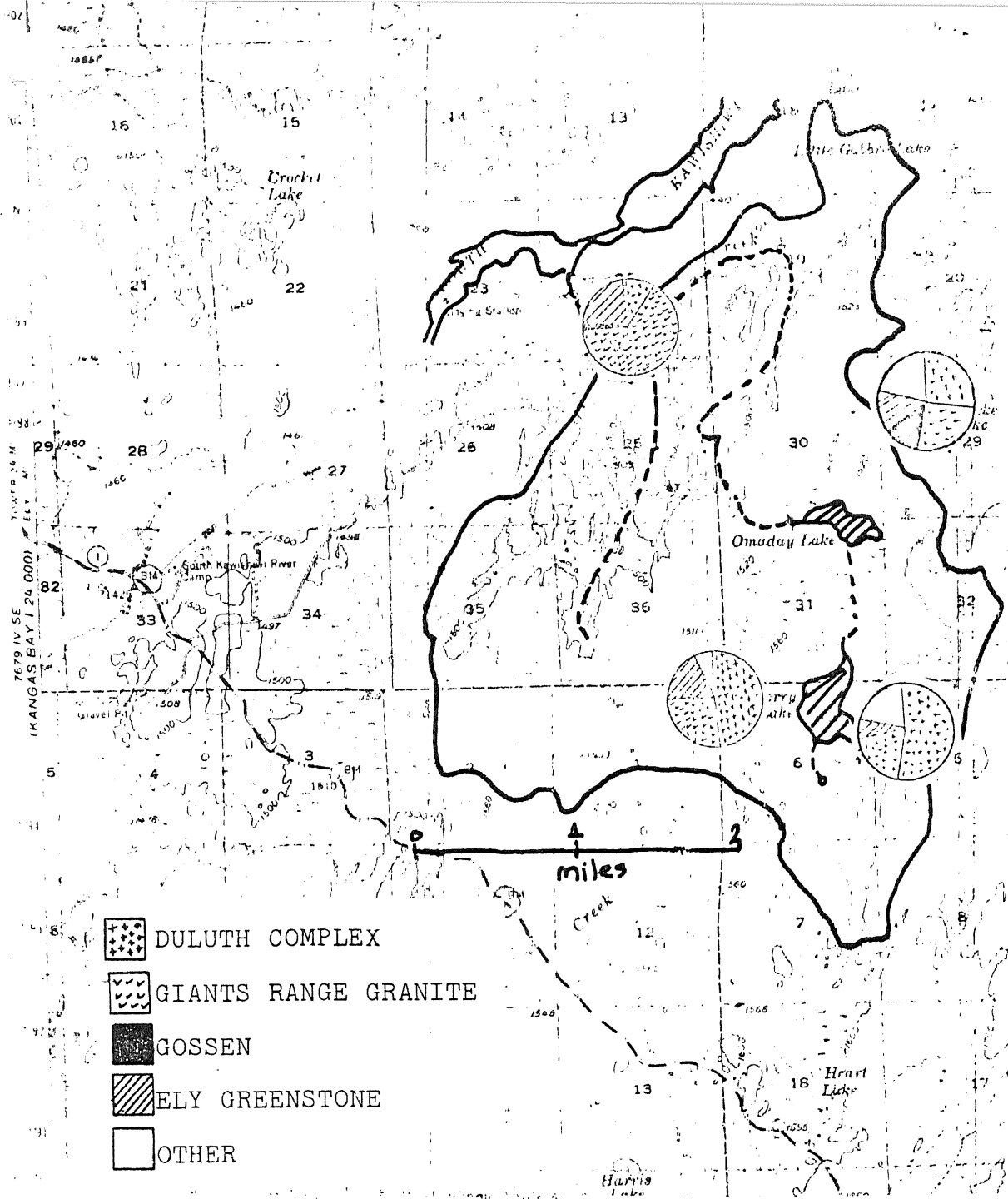


Figure 6 Composition of the **Greater than Pebble Size** Clasts in the Rainy Lobe Till

Because the most reactive part of the till is the finer-grained material having the greatest surface-area-to-volume ratio, 25 till samples were collected for complete sedimentological and mineralogical analysis of the sand and finer-sized fractions.

Methods

(fig. 1)

Sedimentological analysis of till samples included particle size analysis, petrographic description and quantitative analysis of the fine sand-size fraction (Wentworth, 1922), and x-ray diffraction for the identification of mineral species in the clay and silt size ranges.

Till samples were collected at least one and one half feet below the A-horizon to ensure that parent material was sampled.

Standard methods outlined by Folk (1974) were followed for the grain size analyses of till samples of approximately fifty grams each. Dry sieving was done at half-phi intervals for the fine sand to granule size fractions while pipette analyses were performed for the clay to silt size fractions.

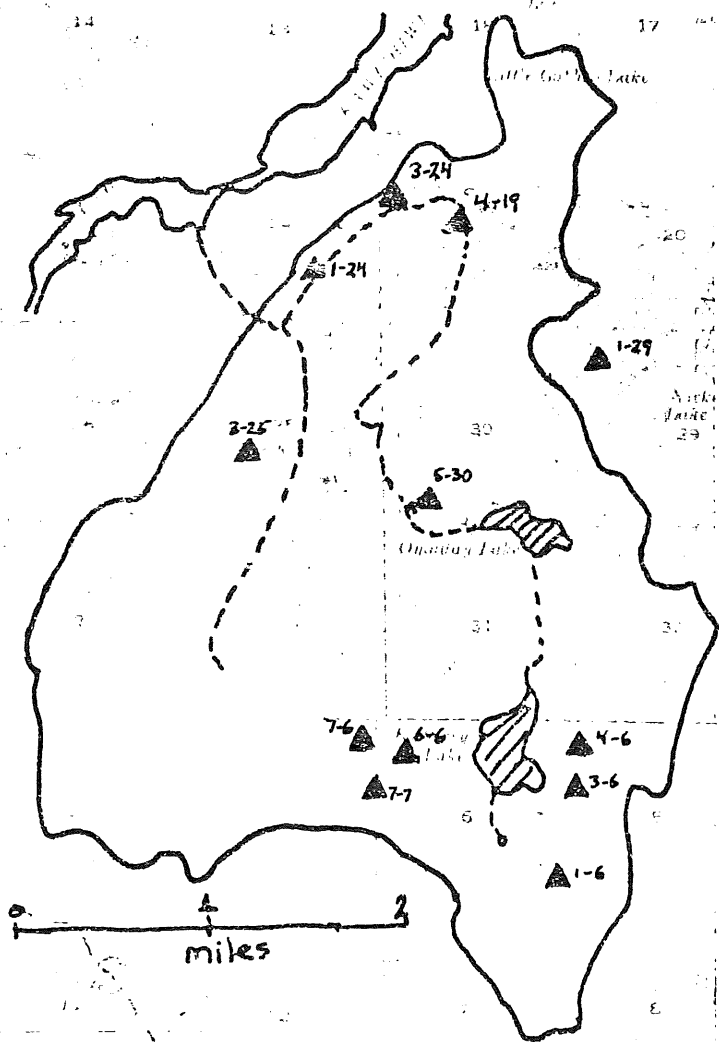


Figure 7 Till Sample Sites in the Filson Creek Watershed

Point counts of at least 400 grains were performed to determine the mineralogy of the fine sand size fraction of the till. Petrographic identifications were made through the use of such references as Krumbein and Pettijohn (1938) and Milner (1954).

X-ray diffraction for identification of clay mineralogy followed the procedures outlined by Carroll (1970) and Whittig (1965). Clays were first treated with hydrogen peroxide to remove organics and then with appropriate reagents to remove free iron oxides before x-ray diffraction. Patterns were run on a Phillips type 42273 x-ray diffractometer using Cu α radiation and a scan speed of one degree 2 θ per minute.

Mineralogical distribution in the very fine sand size fraction

Minerals found in the very fine sand size fraction (0.125 to 0.0625 mm) of the till are a composite of the minerals found in the Giants Range Granite and the Duluth Complex. (Appendix 1).

The major minerals are the plagioclases and biotite, with lesser amounts of olivine, hypersthene, quartz and opaques. Sodic plagioclases from the Giants Range Granite are identified by intensive kaolinization which masks the polysynthetic twinning. Calcic plagioclases are clear and generally free of any alterations. Measurements of extinction angles from cleavage fragments place calcic

plagioclase composition between An₅₀-An₈₀. Hypersthene shows excellent pleochroism from pink to green. Hornblendes are green under plane polarized light and pleochroic from green to brown. Opaques are usually unidentifiable except as aggregates of iron oxide. Olivines are clear to light green and have high relief. Biotite is brown to green in plane polarized light and occasionally "bleached."

Mean percentages of some minerals found in the very fine sand size fraction of the till differ from their modal percentage in source bedrock types (Table 3). Fifteen times as much biotite is found in the till as in either the Duluth Complex or Giants Tange Granite. Only trace amounts of microcline are found in the till, compared to a modal percentage in the Giants Range Granite of about 25 percent. This is due to the difference in susceptibility of the minerals to glacial erosion and transport.

Dreimanas and Vagners (1965, 1971) have determined that during glacial transport and erosion, rocks become abraded to two size modes of two groups of modes. One mode consists of rock fragments and the other consists of mineral fragments within the till matrix. Each mineral can be potentially worn down to a "terminal grade size" dependent upon the original crystal or grain size and its resistance to abrasion. The data suggest that biotite,

Table 3 Comparison between mean mineral composition of the Rainy Lobe till and the mean mineral modes of the Duluth Complex and Giants Range Granite in the Filson Creek Watershed

Major minerals in Duluth Complex	Percentage in size of till (n=11)	very fine sand	Normalized Percentage	Average Mode (77)
Plagioclase	45.9		76.7	89.9
Olivine	2.3		3.8	4.1
Clinopyroxene	0.6		1.0	0.7
Orthopyroxene	5.1		8.5	13.1
Opagues	5.9		9.9	0.9
<hr/>				
Major minerals in Giants Range Granite				
Oligoclase	19.4		48.9	49.3
Quartz	3.7		9.3	11.3
Hornblende	4.2		10.5	8.7
Biotite	12.4		31.2	2.0
K-spar	tr		tr	23.7

which has a prominent basal cleavage, is less resistant than other minerals in the source rocks. The absence of microcline compares favorable with work by Harrison (1960) who found the terminal grade size of orthoclase to be greater than fine sand size.

The larger amount of opaques in the till is because they occur in the mineralized zone in higher percentages than in the Duluth Complex in general.

Clay mineralogy

Kaolinite and chlorite were identified from diffractograms of the finer than +8 ϕ size fraction of six till samples, and were differentiated by heating the samples to 600°C to determine if peaks at 7.15Å and 3.75 Å persisted (Carroll, 1970). At this temperature, kaolinite becomes disordered and these peaks, if present, are attributed to chlorite.

Clay-sized quartz was identified by a strong peak at 3.42 angstroms. Other minor peaks that occur at 3.32, 3.25 and 4.37 angstroms are probably reflections from small quantities of unidentified silicate minerals. Figure 8 shows representative x-ray diffractograms for typical till-clay size particles.

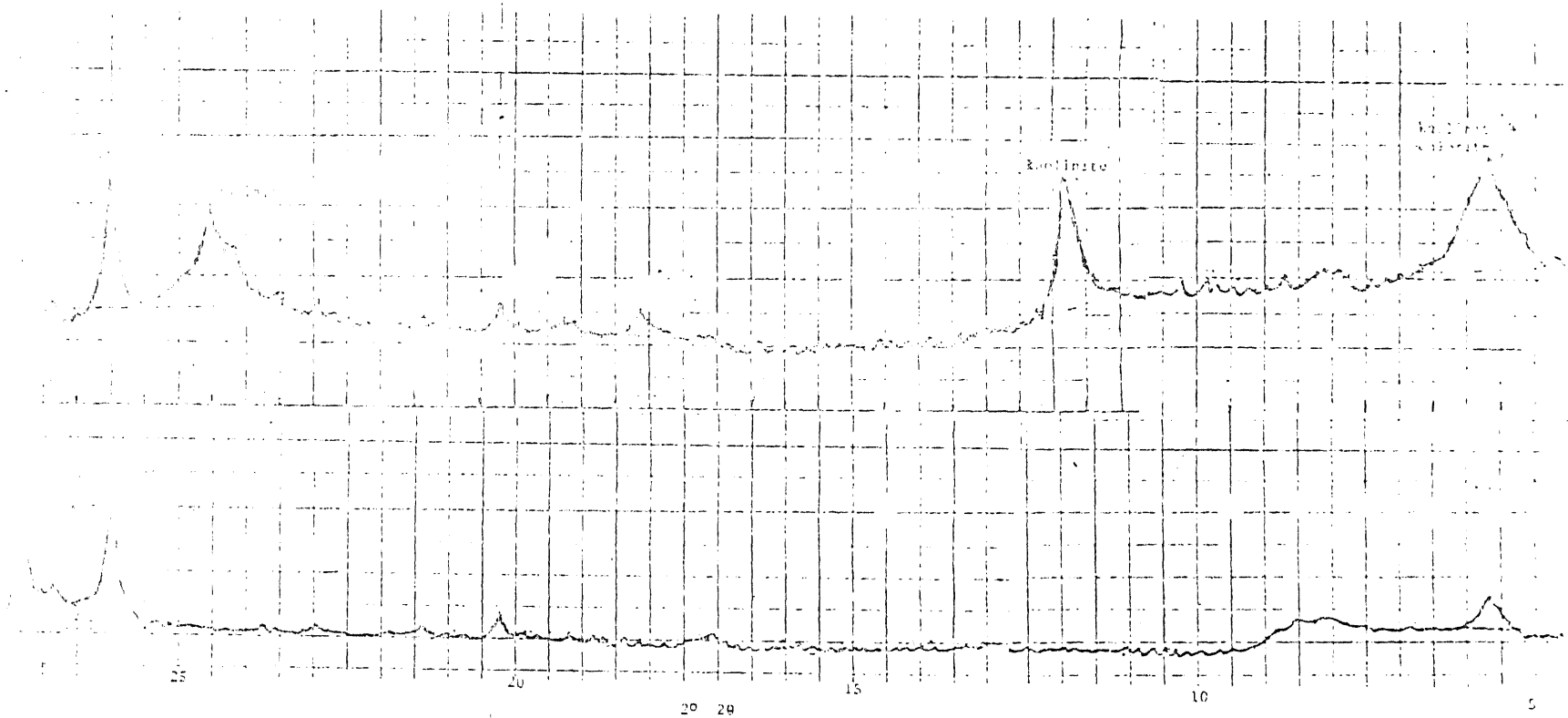


Figure 8 Typical x-ray diffractograms from clay sized fraction of Painy Lobe till collected from the Filson Creek Watershed. Sample 7-6t was glycolated for 1 hour prior to analysis (a), and then heated to 600°C to separate the kaolinite peaks from chlorite (b). Diffractogram showed no peaks beyond 28° 2θ

HYDROLOGY

Introduction

Measurements of precipitation and stream runoff coupled with concentration data for dissolved solids, provide the means to calculate net loading rates in the watershed over the period of the study.

Runoff is composed of three parts: surface runoff, interflow, and ground water runoff. Of these components, ground water runoff is considered to be negligible from the thin till cover. The wetlands store ground water, but because peat ^{is} effectively impermeable ^(Duchon and Holsinger, 1971) below a few feet (Boelter, 1969; [^]), sustained groundwater discharge to the creek is minimal and not enough to even keep the creek open during the winter months.

Interflow through the peat and till occurs during storm events and is incorporated into surface runoff discharged by the creek.

By measuring lake levels, precipitation, lake evaporation, and surface runoff, monthly hydrologic budgets can be estimated from the equation:

$$P = SRO + E_L + \Delta ST_L + ET + \Delta ST_{GW}$$

where: P = monthly total precipitation

SRO = monthly total stream runoff

E_1 = monthly lake evaporation

ΔST_1 = monthly change in lake storage

ET = monthly evapotranspiration

ΔST_{gw} = monthly change in groundwater storage

The first of these gauges (OL-1) was near the mouth of Omaday Lake where North Filson flows out of a small wetland. The second location (F3X) was one-tenth of a mile upstream from the location where North Filson initially crosses the mineralized contact zone. The third location (SF) was at the mouth of South Filson where it enters a reed-sedge bog.

Discharge was measured near Omaday Lake to determine the amount of runoff contributed to the North Filson by the Bogberry-Omaday Lake system. Discharge data at TX location, when combined with chemical quality data, will enable chemical budgets to be calculated for North Filson before it is influenced by the reactions at the mineralized zone. Similar calculations with data from South Filson will give chemical budgets unaffected by either mineralized zone or lakes.

Stages were measured by a Leupold + Stevens level recorder at the OL-1 gauge and Friez Type FA recorders at the SF and F3X gauges. The Stevens recorder was equipped with a four week battery operated clock, and the Friez recorders were driven by weighted clock cables mounted high enough on the gauge house to record 14 days of continuous record before needing to be reset. Stilling wells were constructed of 10 inch diameter stove pipe attached to a platform mounted in a gauge housing

Methods

Precipitation was measured by a U.S. Forest Service recording rain gauge located at the North Central Forest Experimental Station on the South Kawishiwi River and three U.S. Forest Service non-recording gauges located west and south of the watershed (fig. 9).

Linear regressions between the precipitation gauge at the Kawishiwi Laboratory and the non-recording gauges gave high enough correlation coefficients to consider the Kawishiwi Laboratory data as being representative of the Filson Creek Watershed (fig. 10) to within 10 to 20 percent accuracy. Because record at the Kawishiwi Laboratory was terminated at the end of September, precipitation for October, 1977 was taken as the average of that measured at Winton and Babbitt (Climatological Data, 1977).

Stream discharge was measured by the U.S. Geological Survey continuous recording gauge (FlX) near the mouth of the creek and at three other locations by gauges installed for this study (fig. 9).

EXPLANATION

- Lower Filson Precipitation gage (U.S. Forest Service)
- FIX Stream gage (U.S. Geol. Survey)
- SF Stream gage (This study)
- Staff gage
- H-29 Piezometer (U.S. Geol. Survey)
- BSSW Piezometer (This study)

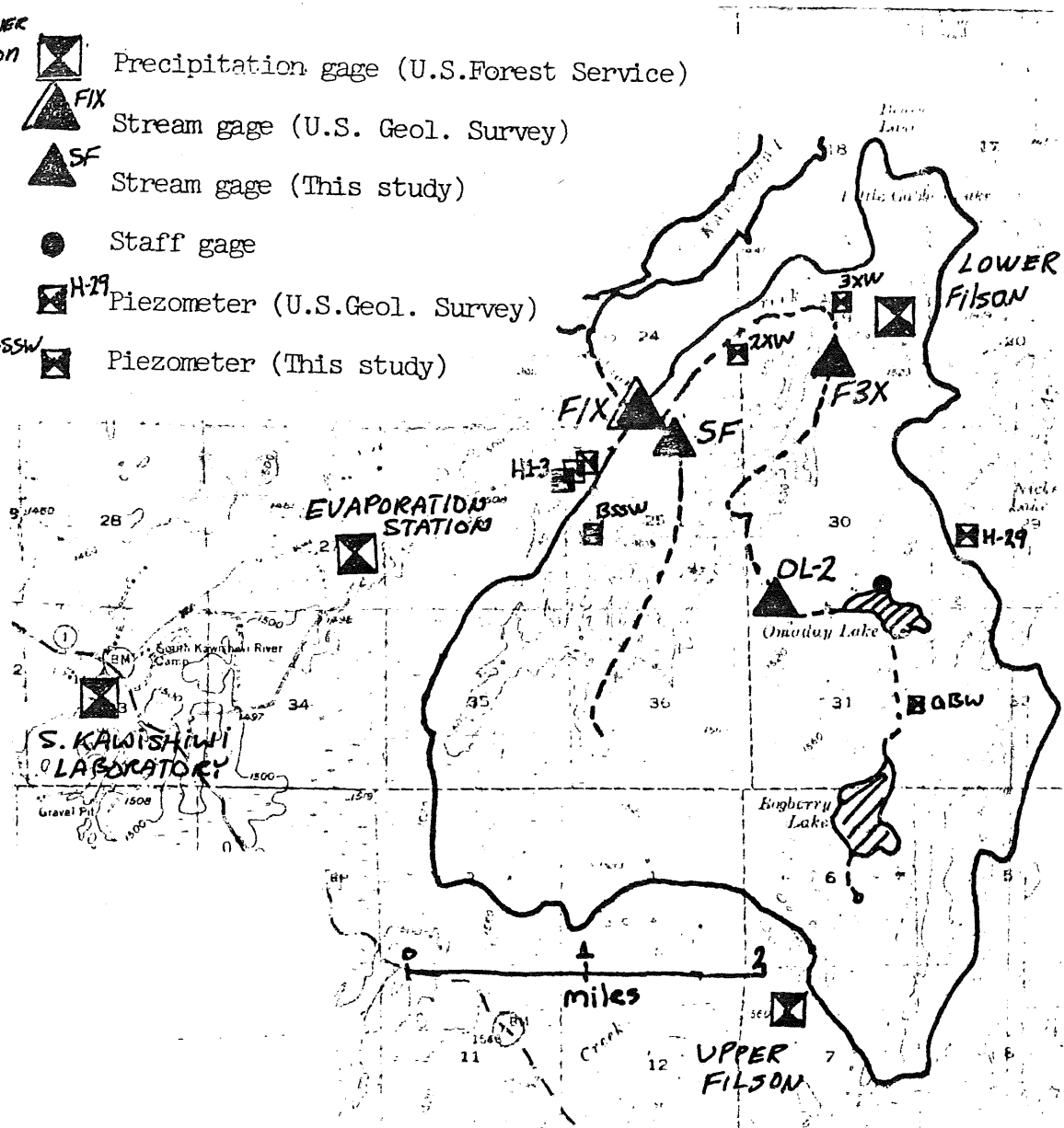
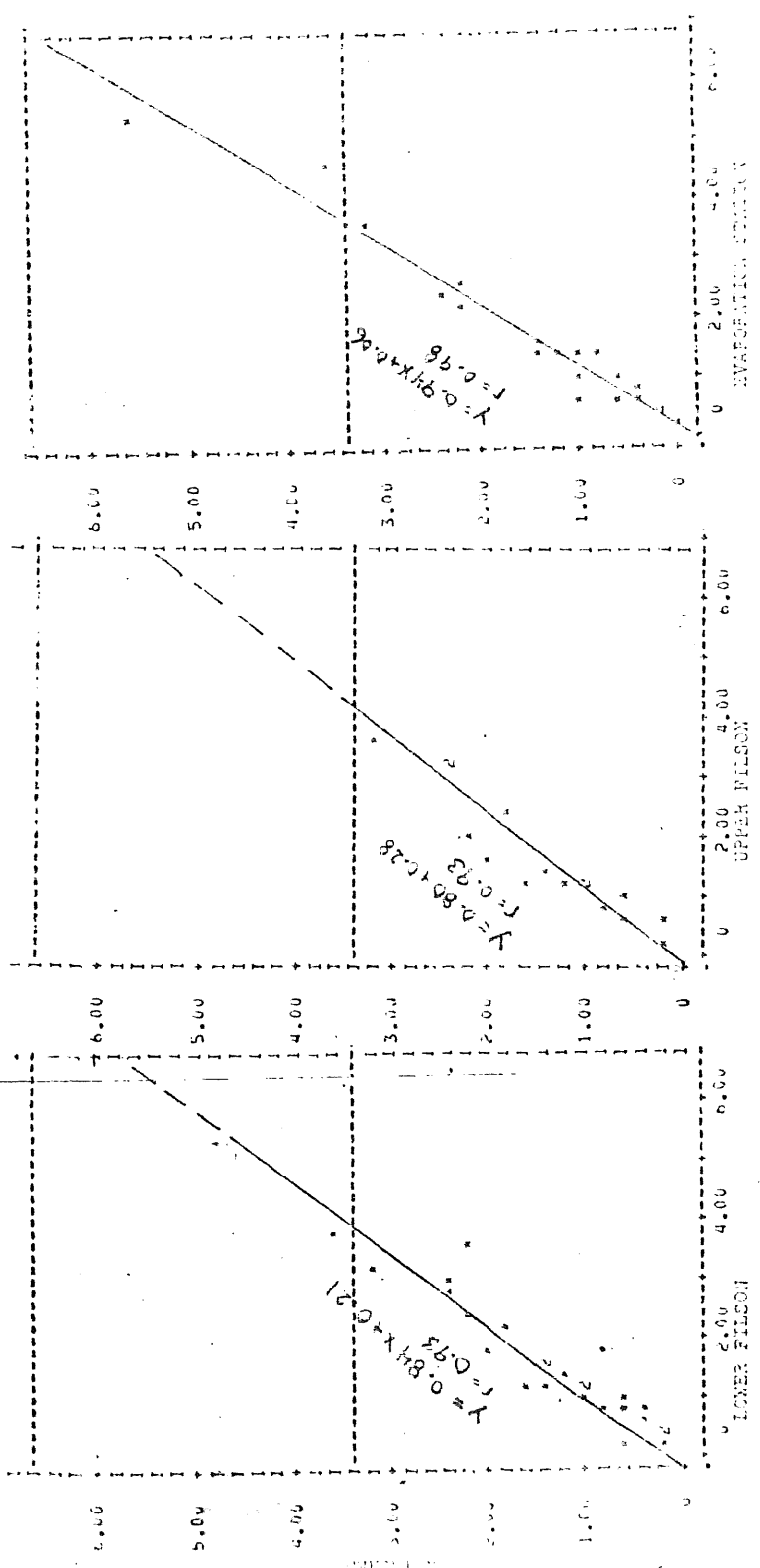


Figure 9 Hydrologic Monitoring Network



CUMULATIVE TWO-WEEK PRECIPITATION, IN INCHES

Figure 10 Correlations between cumulative two-week precipitation at Kawichiki Laboratory and Upper Filson and Lower Filson and Evaporation Station for prevailing rain in U.S. Data collected between 9/1/57 and 9/30/77 by U.S. Forest Service. Raw data are given in Appendix

which consisted of either three feet wide culvert sections or a half inch plywood box. Two inch diameter PVC pipe extended from the base of the stilling wells to the center of the stream channel to assure record during low flow conditions.

Discharge was measured ^(Tables 4-6) by use of a pygmy current-meter and correlated with measurements from staff gauges installed near the recording gauges. Rating curves (fig. II) were established in accordance with procedures of Carter and Davidian (1965). At the SF location, two rating curves were extrapolated from the data to account for a backwater effect caused by a beaver dam built during July, 1977. Previous rating curves were available for the OL-1 location from the U.S. Forest Service.

Ground water levels were measured by four piezometers installed in peat for this study and by five U.S. Geological Survey piezometers previously installed in till ^(fig. Appendix 4). The piezometers in the peat consisted of five to six feet lengths of one and one-quarter inch PVC pipe. A screen was made by slotting one foot of the pipe at one end and wrapping the end with fine nylon mesh to prevent coarse particles from entering the slots. Two inch diameter

TABLE 4 DISCHARGE-STAGE DATA FOR SF STREAM GAGE

<u>Date</u>	<u>Discharge (cubic feet per second)</u>	<u>Stage (feet)</u>
5-10	1.2	1.00
5-27	3.4	1.56
6-11	2.6	1.62
6-25	4.4	1.90
7-02	2.5	1.65
9-12	17.	3.38
9-19*	4.8	2.70
10-15*	16.	3.26
11-05*	1.9	1.73

TABLE 5 DISCHARGE-STAGE DATA FOR F3X STREAM GAGE

<u>Date</u>	<u>Discharge (cubic feet per second)</u>	<u>Stage (feet)</u>
5-10	0.91	1.10
5-28	2.2	1.30
6-11	6.3	1.62
6-25	7.0	1.70
10-15	23	2.54
11-05	3.0	1.42

TABLE 6 DISCHARGE-STAGE DATA FOR OL2 STREAM GAGE

<u>Date</u>	<u>Discharge (cubic feet per second)</u>	<u>Stage (feet)</u>
8-19-75	0.079	3.62
8-19-75	0.083	3.66
10-07-75	.281	3.91
10-07-75	.302	3.91
11-11-75	1.02	4.05
11-19-75	9.07	4.54
4-15-76	25.8	5.13
4-30-76	4.8	4.33

* gage affected by backwater

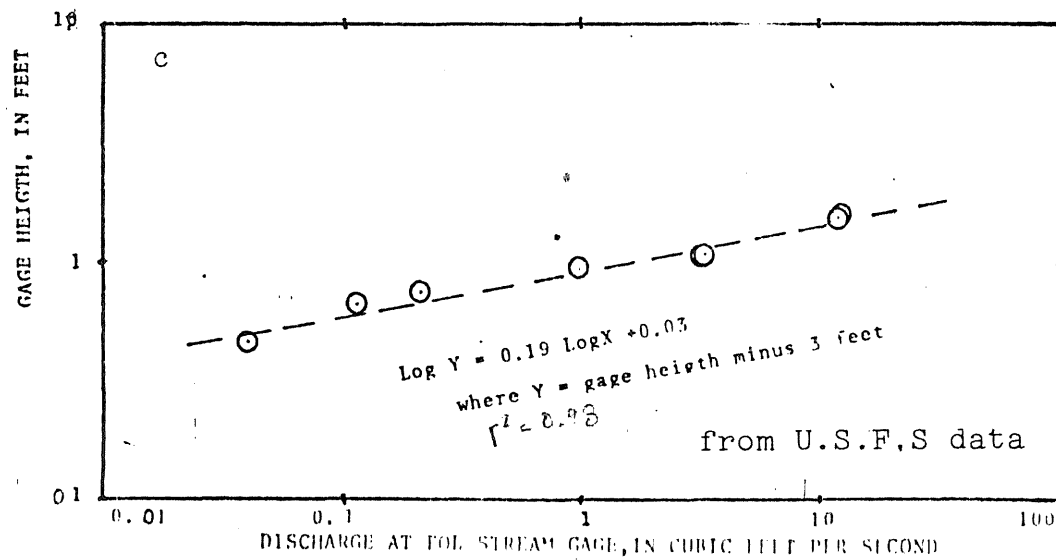
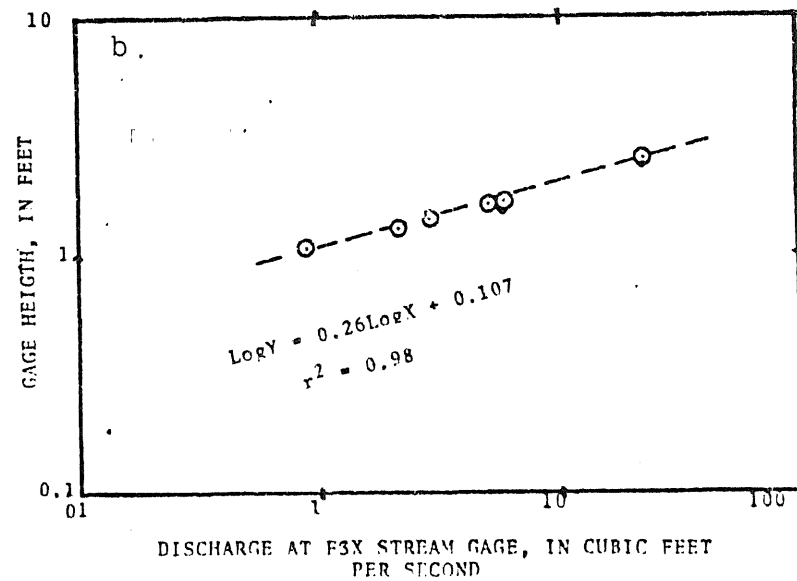
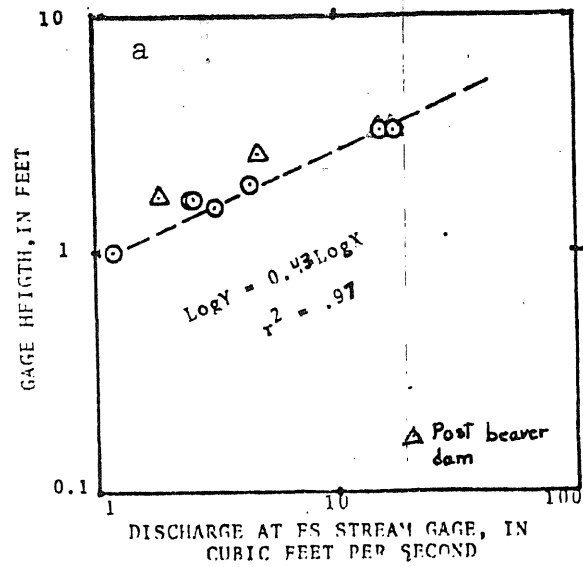


Figure 11 Rating curves for SF (a), F3X(b), and OL-2(c) stream gages

holes were augered by hand into the peat to a depth of three feet, and a two inch diameter PVC casing equipped with a slip cap was placed in the hole. The piezometer was fitted with a plastic collar 1 7/8 inches in diameter immediately above the screen, and then placed in the casing. The casing was pulled out of the hole, leaving the slip-cap at the base of the piezometer. Finally, a slurry of bentonite and water was placed into the hole on top of the collar and the remaining void space was filled with sediment removed from the hole. This arrangement assured that water samples collected from the piezometers were from the base of the peat.

The piezometers installed by the U.S. Geological Survey consisted of two inch PVC pipe and screen installed with the use of a power auger (Olcott and Siegel, 1978). Stratigraphic logs of the materials penetrated are given in the appendix.

Due to vandalism and instrument malfunction, the field record of discharge at OL2, F3X and SF gauges was incomplete. Missing data were generated by use of least squares regression equations (fig. 12) between measured discharges at these gauges and the U.S. Geological Survey gauge at FLX.^(Table 7) Correlation coefficients were high enough to assure accuracy to within an estimated 15 percent

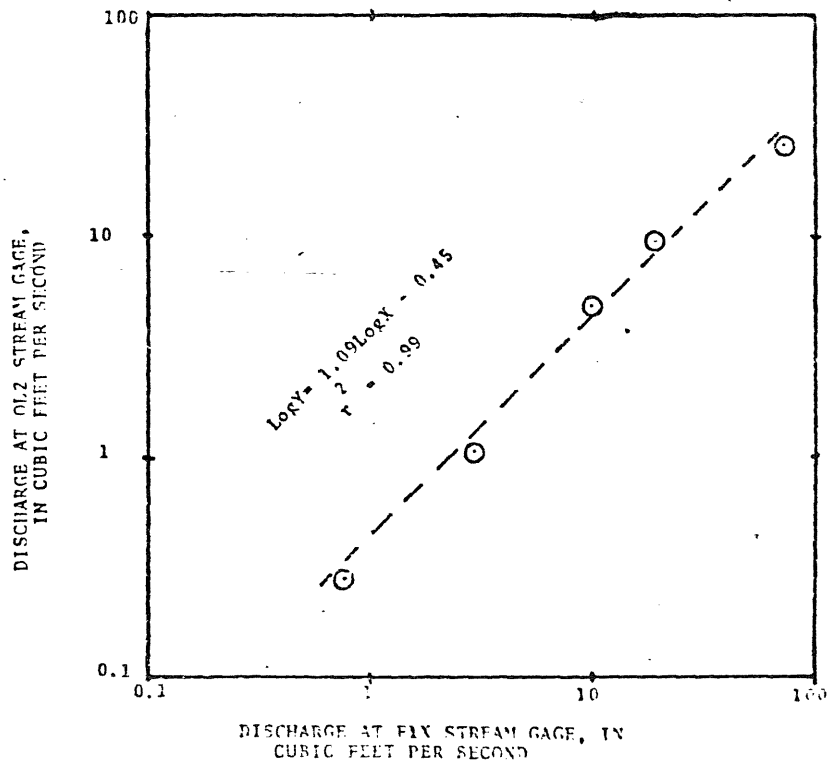
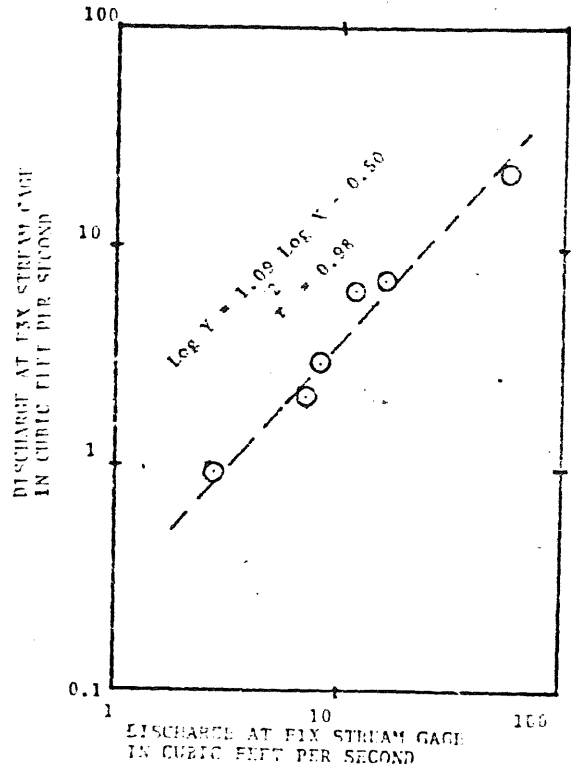
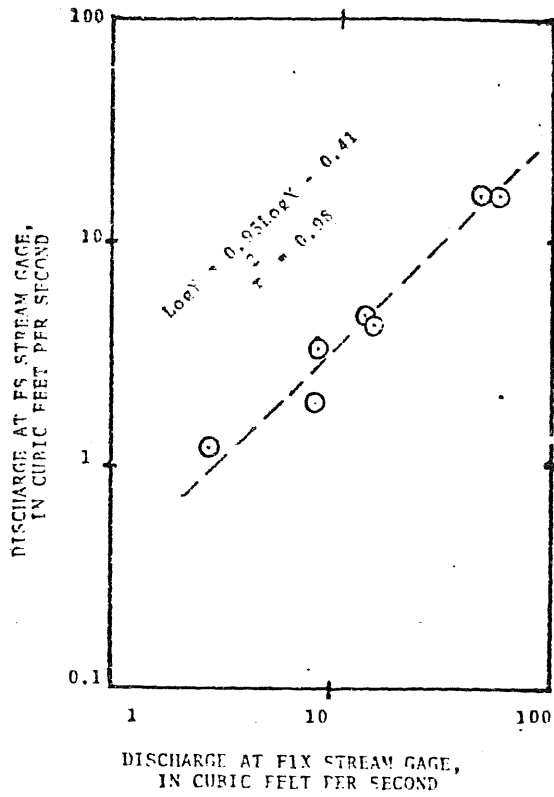


Figure 12 Correlations between discharge at FlX and SF(a), F3X (b) and OL-2(c)

Table 7

DISCHARGE AT STREAM GAGES, IN CUBIC FEET PER SECOND

<u>DATE</u>	<u>F1X</u> ¹	<u>F3X</u> ²	<u>FS</u> ²	<u>FOL</u>
10-07-75	0.78	--	--	0.29 (USFS)
11-11-75	3.8	--	--	1.02 (do)
11-19-75	18	--	--	9.07 (do)
4-15-76	72	--	--	25.8 (do)
4-30-76	4.8	--	--	10.0 (do)
5-10-77	2.8	0.91	1.2	--
5-27-77	8.8	--	3.4	--
5-28-77	7.0	2.2	--	--
6-11-77	12.	6.3	2.6	--
6-25-77	16.	7.0	4.4	--
7-02-77	8.5	--	2.5	--
9-12-77	46	--	17 (USFS)	--
9-19-77	14	--	4.8 (USFS)	--
10-15-77	58	23	16	--
11-05-77	8.2	3.0	1.9	--

1

U.S. Geol. Survey

2

Siegel, except where noted

of the true discharge. This compares favorably with an estimated 10 to 20 percent accuracy for the rating curves proper, given the natural controls of the gauged reaches.

Daily mean discharges at F1X, SE and OL2/F3X gauges are given in appendix 3 . The correlation equations for OL2 and F3X gauges with F1X are virtually identical. Increase in discharge between OL2 and F3X is undetectable because of the limitations of the instrumentation.

RESULTS

During 1977, precipitation occurred as brief and intense storm events. Filson Creek responded to these events quickly, once the soil and till was saturated by March and April snowmelt and rain. Because of the drought condition in 1976, the soil moisture deficit at the beginning of 1977 was abnormally high and the rapid response of Filson Creek to spring snowmelt did not occur as in 1975 and 1976 (figure 13). It is consequently necessary to determine the amount of recharge that went into soil moisture and surface depression storage before calculating the hydrologic budget for the watershed during 1977.

Total recharge to the watershed during early 1977 can be calculated as the sum of the water content of the snow cover at the end of March and the amount of precipitation during March and April. Snow course data by the U.S. Forest Service in the Superior National Forest indicated that the water content of the snow pack at the end of March was $2.0 \pm .1$ inches (Ramquist, 1978, personal communication). Precipitation in April was 1.1 inches. The total recharge of 3.1 inches resulted in only 0.2 inches of runoff at the FlX gage. The difference between the recharge and discharge gives an estimate of the soil moisture deficit, depression storage, and groundwater deficit of about 2.9 inches, neglecting

DISCHARGE, IN CUBIC FEET PER SECOND

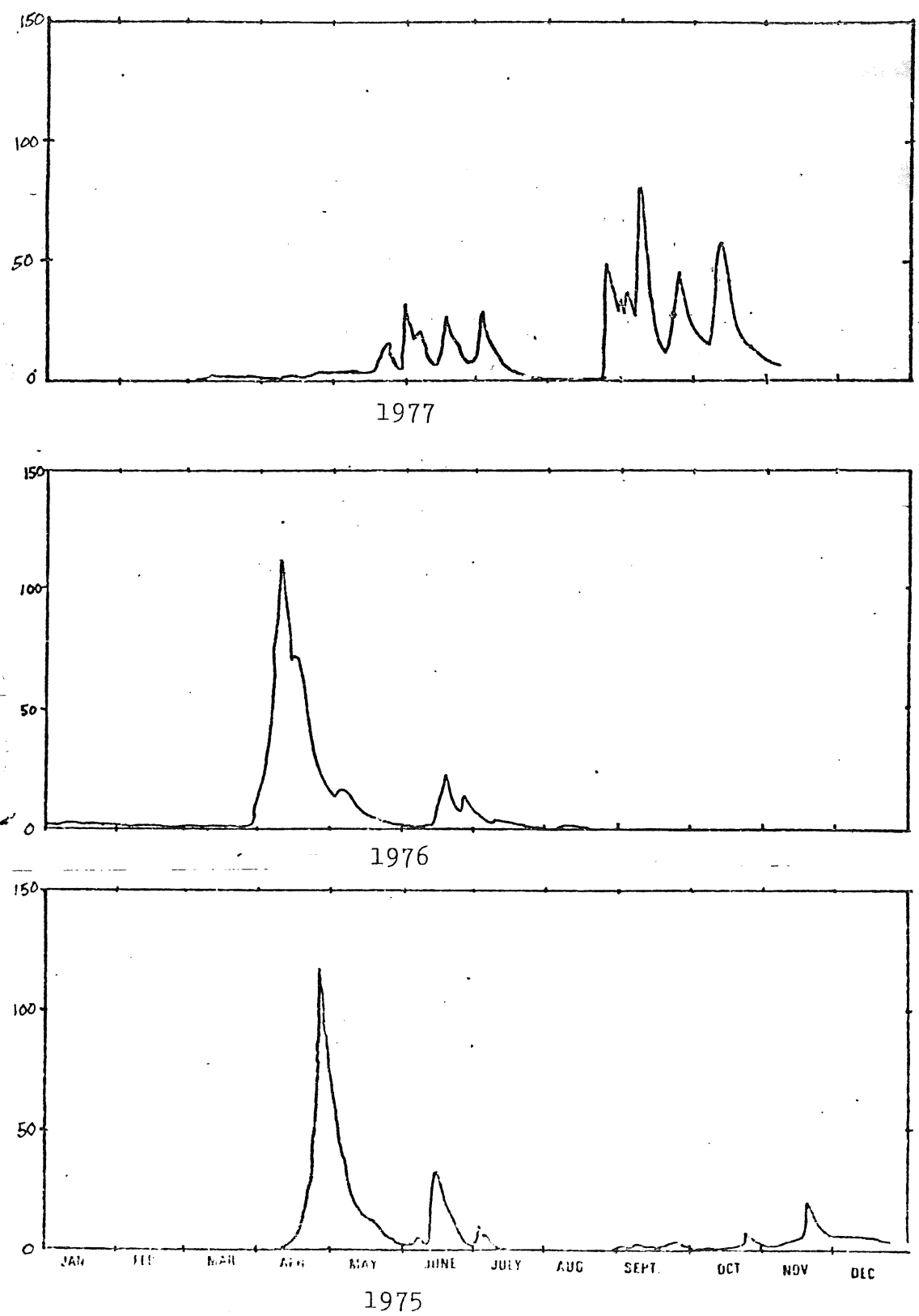


Figure 13 Annual discharge at FlX gage

evapotranspiration which would be minimal because of the very dry soil conditions at the beginning of the growing season.

The soil moisture deficit^{in 2.4 in} probably extended through early May when precipitation events still were not reflected in discharge hydrographs at Filson Creek. The estimate of 2.9 inches may be consequently on the small side.

Part of the recharge to Filson Creek watershed went into surface storage, mainly Omaday and Bogberry Lakes. Precise lake level measurements are not available prior to May 1977. However, correlation between the lake level measurements from May to June with discharge at FlX gage gives a correlation coefficient of 0.98 (figure 14). A plot of lake level versus time (figure 15) indicates that the linear rise in stage from spring recharge stopped at the end of June and then declined during the summer months until the heavy precipitation events in the fall. This decline is attributable to less recharge during the summer months and the increase in evaporation from the lake surface. From the correlation between lake level and FlX discharge, it is estimated that the lake level at ice melt in March was about 0.3 feet, relative to the staff gage installed in May. The rise in lake level of about 0.7 feet from snowmelt and precipitation until the end of April is equivalent to storage of 0.2 ± 0.05 inches. This figure subtracted from the combined soil moisture deficit and depression storage of 2.9 inches gives an estimated soil moisture deficit at the beginning of 1977 of from^{about} 2.6 to 2.7 inches.

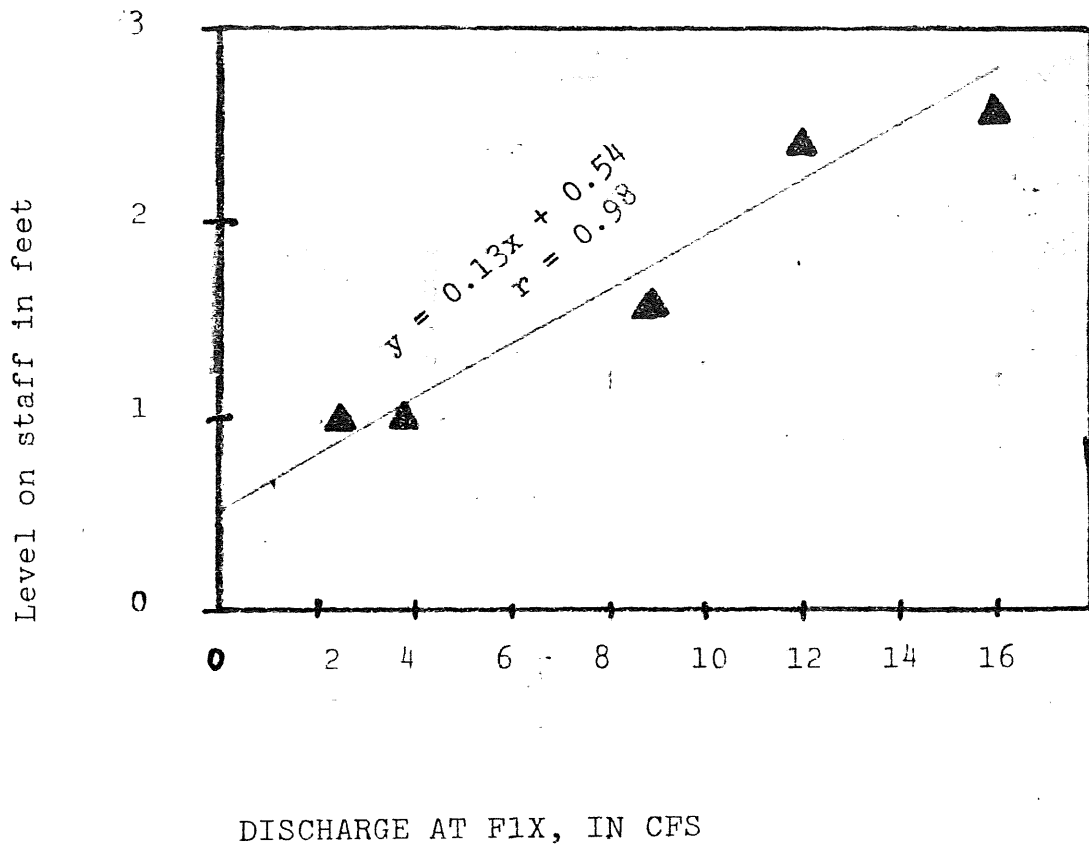


Figure 14 - Correlation between Omaday Lake level and Discharge at Flx Gage

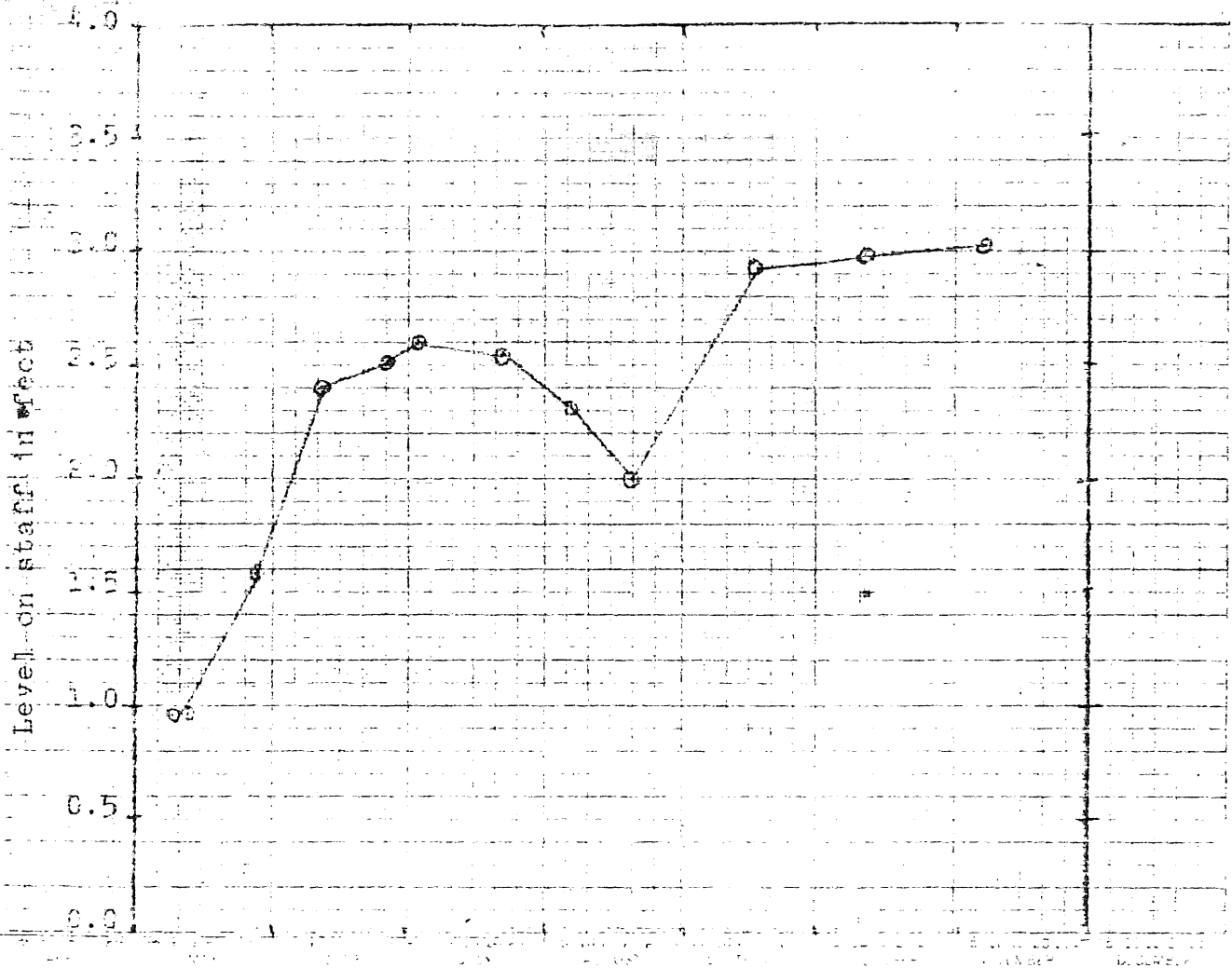


Figure 15 - Water Levels at Omaday Lake, 1977

Soil moisture deficit can also be calculated by using the Thornthwaith water balance method (Thornthwaith and Mather, 1955, 1957).

The Thornthwaith method contains assumptions which may not be directly applicable to Filson Creek Watershed. These assumptions are:

- 1) that there is negligible surface depression storage
- 2) that the rate of ET is uniform over the watershed
- 3) that the effects of humidity and wind velocity are minimal

Filson Creek watershed contains different types of vegetation and soils, lake storage, and indeterminate amount of surface depression storage. Despite these limitations, the method should give an approximation of the magnitude of the combined soil, lake and surface depression storage deficit that existed at the end of the 1976 drought.

The water balance equation used in the method is:

$$P = RO + INF + ET + PERC$$

where: P = monthly precipitation

RO = monthly discharge at FLX

INF = infiltration into the soil and till

ET = evapotranspiration

PERC = recharge to the groundwater system

The method uses mean monthly temperature to calculate a total annual heat index, a dimensionless measure of the amount of heat energy received on the surface of the watershed in one year. The index is applied in an equation that calculated a monthly potential evapotranspiration which is adjusted for latitude differences in the amount of sunlight. The actual evapotranspiration is limited by the moisture content in the soil. In the accounting procedure, the monthly potential evapotranspiration is subtracted from the monthly infiltration to determine when infiltration is in excess. When (INF-PE) is negative, soil moisture loss will occur. For use on a monthly basis, the negative values are summed from month to month. In order to determine how much evapotranspiration will remain in the soil after a given amount of ET, the cumulative negative values are applied to soil moisture retention tables developed from experimental work (Thorthwaith, 1957). Finally, actual ET is calculated through use of the equation:

$$AE = PET + ((INF - PET) - \text{change in soil moisture storage})$$

Tabulation begins after the spring snowmelt in 1976 when it was assumed that the soil and till were fully saturated. It was assumed that 85% of the Filson Creek watershed is covered by an average of 6 inches of till, classified from grain size analyses as ^{gravelly silt to} gravelly, sandy loam (figure 16). This soil type would have an available water capacity of about 0.10 inches/inch (Irrigation Guide for Minnesota, 1979). The average thickness cited is an estimate based on Prettyman (1976) and the author's experience in the field.

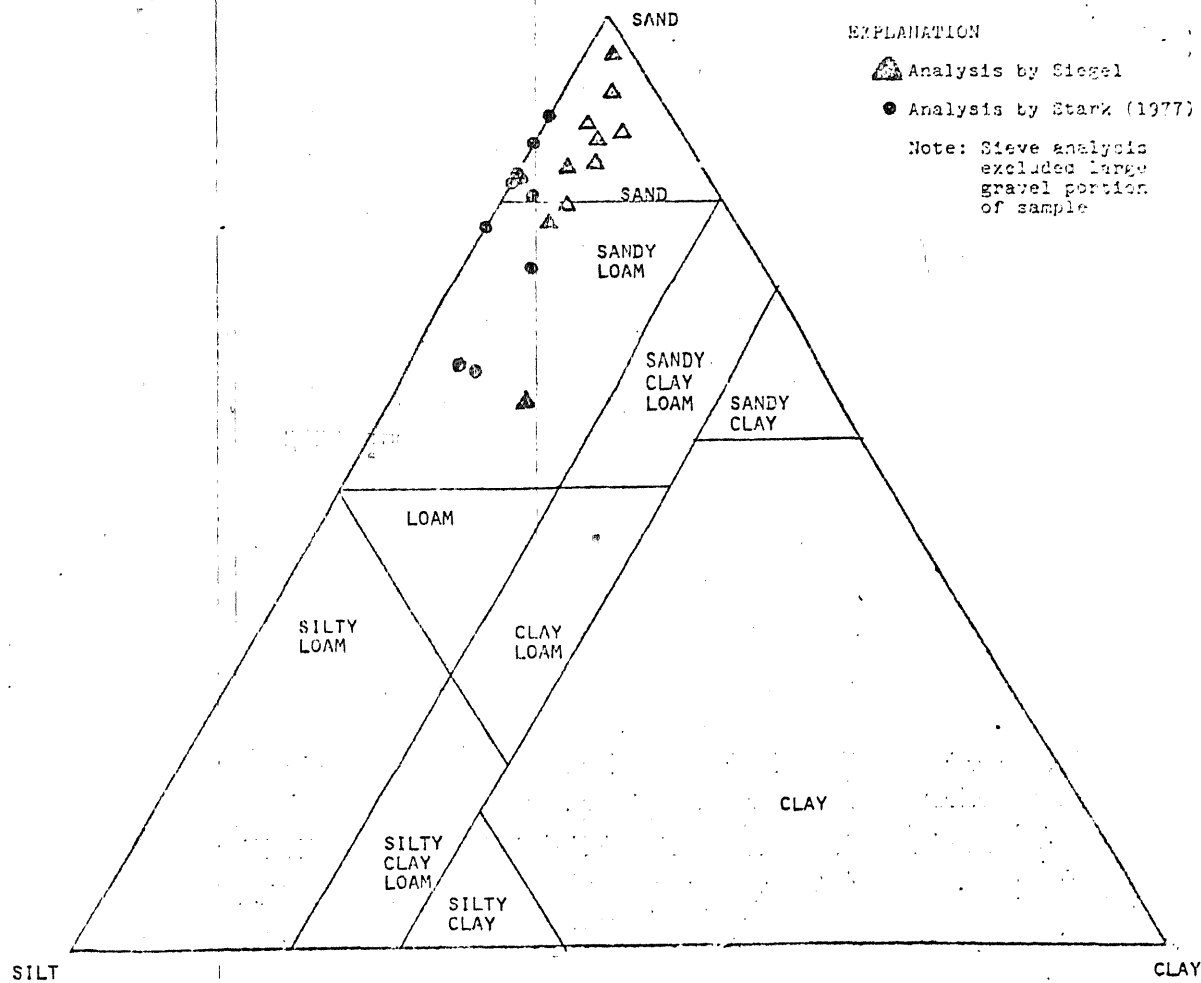


Figure 16 - GRAIN SIZE DISTRIBUTION OF TILL SAMPLES COLLECTED FROM
FILSON CREEK WATERSHED

The remaining 15 percent of the watershed was assumed to be covered by woody peat having a specific yield of 0.25 inch/inch (Boelter, 1966). Available moisture content, defined as the difference between the specific yield and the wilting point, is used in the Thornthwaith calculations. The concept of available moisture content generally applies to soils above the water table, and not to wetlands having the water table at or near the land surface. It is assumed, therefore, that the storage capacity of the peat lies somewhere between the porosity, of about 0.9 inch/inch, and the specific yield. For the Thornthwaith calculations, an average value of 0.6 inch/inch is assumed, realizing that the error may be as much as 30 percent.

Normalizing for the entire watershed gives an average soil cover having an "available moisture content" of between three and four inches/inch.

Table 8 gives the monthly water balance from the end of May to November 1976. The data indicate that the soil moisture deficit at freeze-up in November was between 2.8 and 3.5 inches, which compares favorably with the estimate obtained from the difference between precipitation and runoff during March and April 1977.

Table 8

Soil Moisture at the End of 1976, calculated
by the Thornthwaith Water Budget Method

Month *	Temp., in C ^o	i	Unadj. PE	PE	P	RO	INF (P-RO)	(INF-PE)	Σ -(INF-PE)	Storage
June	17.2	6.61	3.5	4.6	5.9	0.9	5.0	0.4	---	4.0
July	18.9	7.49	3.8	5.1	1.9	0.2	1.7	-3.4	-5.5	1.6
Aug	17.9	6.90	3.6	4.4	1.7	0.0	1.7	-2.7	-6.1	0.8
Sept	12.8	4.15	2.6	2.7	1.7	0.0	1.7	-1.0	-7.0	0.6
Oct	2.1	0.27	0.2	0.2	1.2	0.0	1.2	-1.0	-8.0	0.5
Nov	-6.7	---	---	---	---	0.0	---	---	---	---

* Assuming an available soil moisture capacity of 4.0 inches

Month *	Temp., in C ^o	i	Unadj. PE	PE	P	RO	INF (P-RO)	(INF-PE)	Σ -(INF-PE)	Storage
June	17.2	6.61	3.5	4.6	5.9	0.9	5.0	0.4	---	3.0
July	18.9	7.49	3.8	5.1	1.9	0.2	1.7	-3.4	-5.4	0.9
Aug	17.9	6.90	3.6	4.4	1.7	0.0	1.7	-2.7	-6.1	0.4
Sept	12.8	4.15	2.6	2.7	1.7	0.0	1.7	-1.0	-7.0	0.5
Oct	2.1	0.27	0.2	0.2	1.2	0.0	1.2	-1.0	-8.0	0.2
Nov	-6.7	---	---	---	---	0.0	---	---	---	---

* Assuming an available soil moisture capacity of 3.0-inches

A third method of estimating the storage deficit is to assume that most of the storage capacity is located in the wetlands.

At the end of October 1976 when the piezometers were installed for the study, saturation in the wetlands along Filson Creek and between Omaday and Bogberry Lakes was from 1 3/4 and 2 feet below the land surface. Much of the peat above the water table was visually "powder dry". Assuming a porosity of 90 percent (Boelter, 1964, 1969) for the woody, fibric peat typical of the fens, a value of between 2.8 and 3.2 inches of storage capacity is obtained.

Assuming an average value of 0.5 feet of till with an average porosity of 10 percent over 85% of the watershed, an additional 0.5 inches of storage is obtained. Summing gives a range in available storage between 3.3 and 3.7 inches. This range is undoubtedly high because upon recharge, all the void spaces in the ^{dry and till} peat would not be immediately filled and because the peat was partly saturated ^{immediately} above the water table.

From these varied methods the best estimates of the storage deficit (lake, surface depression, and soil moisture) available at snowmelt in 1977 range from 2.8 to 3.5 inches. By the end of May 1977, wetlands were not only completely saturated, but contained free standing water. For the remainder of the year, most of the remaining storage in the watershed was within the lake basins and periodic saturation of the thin till overlying the bedrock ridges.

Groundwater levels/^(fig 17) in piezometers BSSW and QBW, located in headwater wetlands, generally fluctuated at or near the land surface. Water levels fluctuated more widely in piezometers 2xW and 3XW which were located in fens immediately adjacent to Filson Creek. The greater amplitude in water level changes probably is a result of many factors, including bank storage after large storms, and low hydraulic conductivity in horizons deeper than three feet. Standing water up to 0.5 feet deep on the fens correlated directly with water levels measured in the piezometers. This standing water remained at 3XW and 2XW even when Filson Creek, only a few feet away, was over a foot lower.

An anomalous dry period^{recorded} during July for 2XW was caused by clogging of the well screen. After the screen was replaced, water levels returned to levels above the land surface comparable to the free standing water.

FLEVATION WITH RESPECT TO LAND SURFACE,
IN FEET

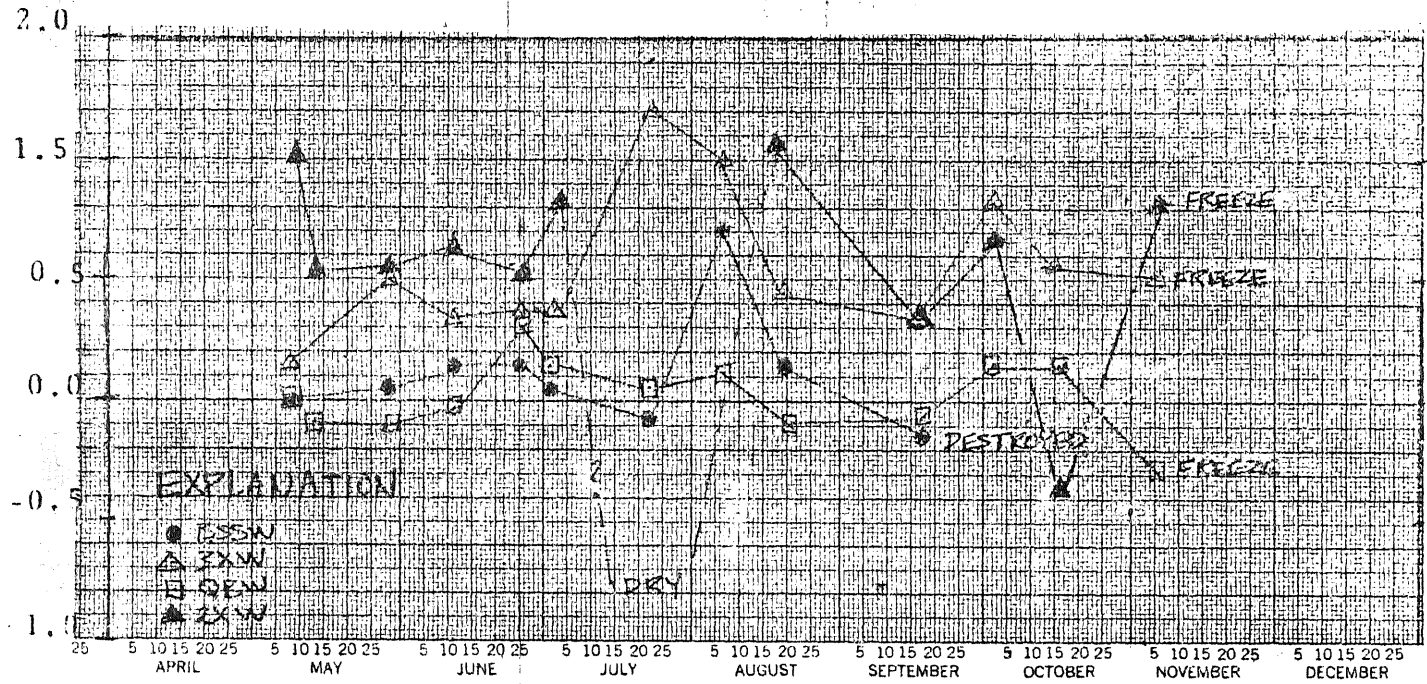


Figure 17 Water Level in Wetland Piezometers, 1977

Table 9 gives the estimated monthly water budget for Filson Creek watershed from May through September 1977 above the FlX gage. Changes in lake storage were estimated from lake level reading at Omaday Lake. Evaporation from the lakes was calculated by taking the product of 0.7, the lake area, and the evaporation measured by the U.S. Forest Service at an evaporation pan located about four miles southwest of FlX gage. Based on the groundwater hydrographs of the wetland piezometers located in the headwater areas, it was assumed that the wetlands were ~~effectively~~ ^{effectively} saturated and groundwater storage remained effectively constant. Evapotranspiration was calculated as a residual. Due to the very flashy precipitation events that occurred in both spring and fall of 1977, the Thornthwaith Water Budget method proved to be non-usable. The method is best used during seasons with one wet and one dry period, ^{such} as occurred in 1976. When multiple precipitation events of similar or greater intensity occur late in the year, the method is very difficult to use and its reliability falls, especially in watersheds as complicated as Filson Creek.

Table 9 Estimated Water Balance for Filson Creek
Watershed between May and October, 1977

Month	Precipitation ³	Stream Runoff	Lake Evaporation ¹	Lake Storage ²	ET
May	5.1	0.6	0.9	+0.2	3.4 ^a
June	4.4	1.9	0.7	+0.2	1.6
July	3.3	1.0	0.8	-0.1	1.6
Aug	5.9	0.7	0.9	0.0	4.3
Sept	6.2	3.8	0.7	+0.1	1.6

¹Evaporation calculated using U.S.F.S. pan data

²Lake storage calculations assumed steep nearshore gradient at lake edge

³Precipitation data from South Kawishiwi Laboratory

^aET value includes unknown amount of soil moisture deficit at beginning of May

WATER QUALITY

Methods

Sampling for water quality was done approximately biweekly from March, 1977 to February, 1978 from a total of four locations on Filson Creek and Omaday Lake, four piezometers placed in wetlands in the Filson Creek Watershed, three precipitation collectors; and six infiltration water collectors placed at the interface between the A and B-soil horizons (fig. 10).

Filson Creek water samples were collected at midstream position near gauging points. Sampling was done by submerging a high density one liter polyethylene bottle below the water surface at approximately 2/3 of the total water depth, and allowing it to completely fill. Sample bottles were pre-rinsed ^{once} with ten percent nitric acid and three times each with de-ionized water and stream water at the time of sampling. Separate samples were taken for total copper and nickel analyses by submerging a fifty milliliter teflon bottle which was pre-rinsed with ultra-pure nitric acid and then washed out with de-ionized water. Samples collected for trace metal analyses were immediately acidified in the field with one milliliter of ultra-pure nitric acid.

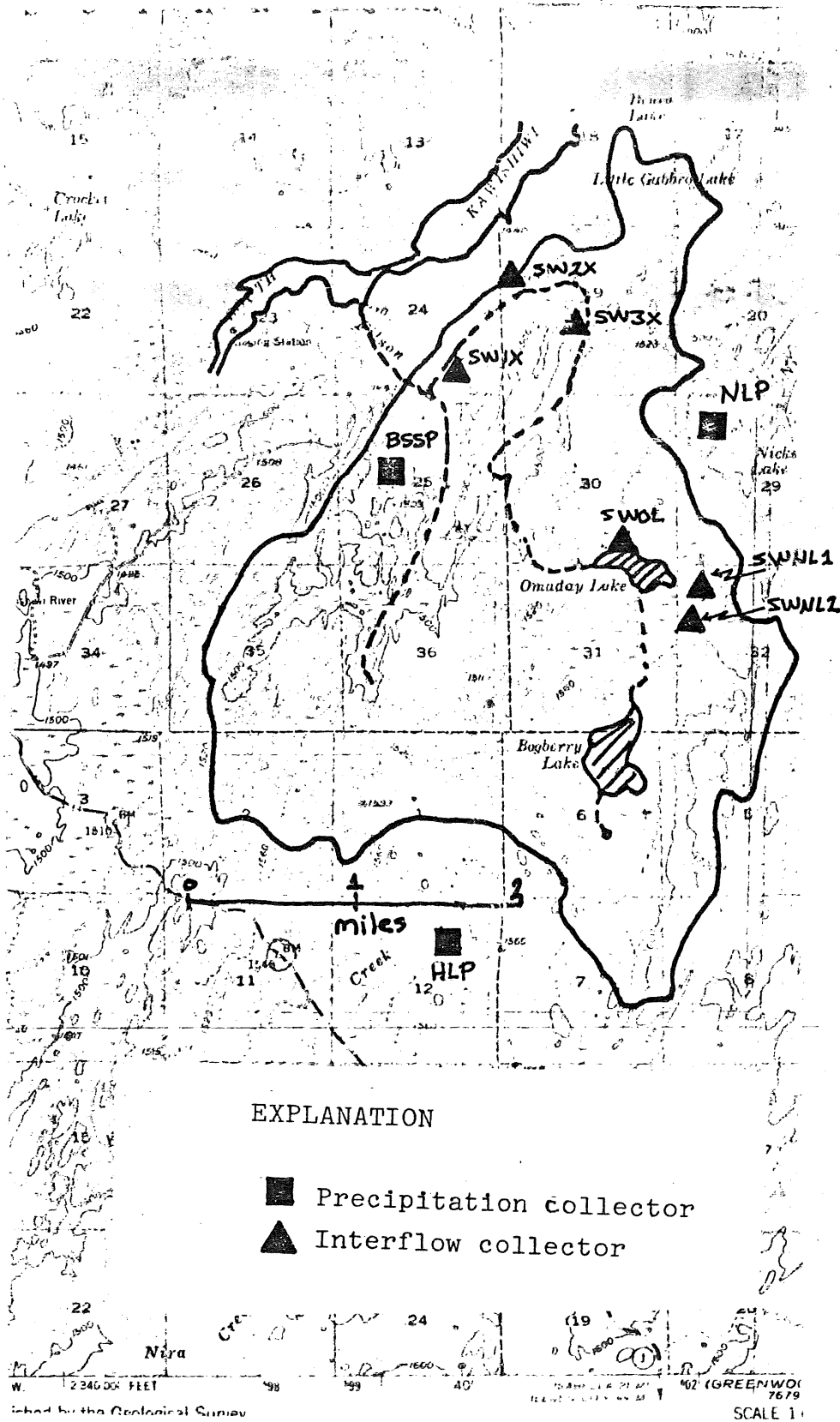


Figure 18 Location of precipitation and shallow interflow water collectors

Omaday Lake was sampled at a depth of two feet at a distance of fifteen feet from the north shore. Because Omaday Lake is less than five feet deep, it is assumed that it is fairly well mixed during the year and that the water samples were representative.

Water temperature and pH were determined in the field for all samples except for surface infiltration and precipitation water, which were assumed to be in equilibrium with ambient temperatures, and neglecting other acid inputs, atmospheric carbon dioxide partial pressure.

Alkalinity was determined in the field for all ground water samples and within six hours of collection for Filson Creek, Omaday Lake, precipitation, and shallow surface interflow water. Titrations and pH measurements were done with a Radiometer #51 field pH meter. Alkalinity was done by potentiometric titration to the inflection point between 5.2 and 4.8 pH. When air temperature was below freezing, pH and alkalinity were determined in the laboratory.

Specific conductance was measured in the laboratory with a Radiometer Specific Conductance meter. Chloride was determined by the mercurimetric method (Brown and others, 1970). Sulfate was determined by the turbidimetric and silica by the ammonium molybdate methods (Standard Methods, 1978). Turbidity and color development were

measured by a Beckman Model 24 spectrophotometer. Color blanks were run for all samples. Total and dissolved calcium, magnesium, sodium and potassium were determined by use of a Perkin Elmer Model 303 Atomic Absorption Spectrophotometer (Brown and others, 1970). When analyzing for calcium and magnesium, lanthanum oxide solution was added to the sample to mask interference from other cations (Perkin Elmer,).

Total copper and nickel were analyzed using a Perkin Elmer Model 360 Atomic Absorption Spectrophotometer equipped with a graphite furnace. Methods used were from the Perkin Elmer (manus). (1970).

Filter blanks were analyzed with each set of samples both for Nucleopore and Gelman 0.45 micron filters and for the Whatman #40 filters. Except for calcium, cations contributed in the filtering procedure were non-detectable in all cases. Anomalous calcium concentrations in filterates early in the project were determined to have come from the liners in the small polyethylene bottles used to store the filtered samples. The replacement of the caps eliminated further contamination.

A summary of the techniques used and the minimum detection limits are given in Table 10. Concurrent with the project, split samples were run in cooperation with the ^{Regional} Copper-Nickel Study Staff and the U.S. Geological

Table 10 Summary of Analytical Techniques

Constituent	Method	Level of Detection	Source
Bicarbonate	Potentiometric titration	1 mg/l	Brown and others (1974)
Sulfate	Turbidimetric	1 mg/l	Standard Methods (1975)
Chloride	Mercuric Nitrate	0.1 mg/l	Brown and others (1974)
Silica	Ammonium molybdate	1 mg/l	Standard Methods (1975)
Calcium	Atomic-absorption		
Magnesium	Atomic-absorption		
Sodium	do.		
Potassium	do.		
pH	Instrument method		Brown and others (1974)
Specific conductance	Wheatstone bridge method		Standard Methods (1975)
Copper	Atomic-absorption		
Nickel	Atomic-absorption		

Survey for quality control assurance. With the exception of sulfate and bicarbonate results for sample CONF, quality control results are favorable (Table 11).

The sulfate discrepancy for sample CONF probably reflects problems associated with the analysis for sulfate at low concentrations in organic rich and colored waters. The U.S. Geological Survey uses the thorin method which, in part, removes metal cations from the sample by ion exchange prior to a titration procedure (Brown and others, 1970). My work and other laboratories cited in Table 11 used the turbidimetric technique, which spectrophotometrically measures the degree of turbidity caused by the precipitation of $BaSO_4$ in a acid solution.

To assess the reliability of the turbidimetric technique beyond the replicate samples, Filson Creek water in which sulfate was non detectable was "spiked" with known amounts of sulfate in a blind test. Results of the test are given in Table 12 . More tests are being done, but the results suggest that the turbidimetric method will give accuracy to \pm 40 percent at low concentrations. Furthermore, the tests suggest that the^e results will tend to be high, and inconclusive below 2 milligrams per liter.

Table II Quality Control Program

Concentration, in milligrams per liter									
Lab	Sample	Ca	Mg	Na	K	Cl	SO ₄	HCO ₃	Si
U.S.G.S.	CONF	3.0	2.0	0.7	0.1	1.8	9,4.8,	8	12
Siegel	CONF	2.9	1.9	1.0	0.4	1.8	5.7,6.3 N.D. ^a	3	12
Minn. Dept. ^a Health	BBL								
Siegel ^b	BBL	9.6	6.6	3.2	0.6	1.5	3.5	26	3.5

^b N.D. Not detected
^a Total cation concentration

Lab	Sample	Cu µg/l	Ni µg/l	SO ₄ mg/l	Alk mg/l
Environmental Research Lab- Duluth	D2511771	1.62±.07	1.44±.13	--	--
	D2511772	1.72±.12	1.08±.16	--	--
Erie Mining Company	D2511771	1	<5	33	24
	D2511772	1	<5	33	16
State Health Department	D2511771	1.8	<1	30	12
	D2511772	1.7	<1	27	11
SERC0	D2511771	1.4	1	14	10
	D2511772	1.4	1	12	6
Eisenreich	D2511771	--	--	--	--
	D2511772	--	--	--	--
Siegal	D2511771	1.1	N.D.	25.6	12.0
	D2511772	0.9	2.0	25.6	12.0

Lab	Sample	Cu µg/l	Ni µg/l	SO ₄ mg/l	Alk mg/l	Silica mg/l
Environmental Research Lab- Duluth	BB328771	3.49±.10	57.8±4.0	-	-	-
	BB328772	3.79±.12	60.4±4.4	-	-	-
Erie Mining Company	BB328771	4	61	62	84	-
	BB328772	3	60	54	86	-
State Health Department	BB328771	7.9	49	54	71	34
	BB328772	8.8	56	55	71	34
Eisenreich	BB328771	2.4	51.2	51.0	-	-
	BB328772	2.4	55.2	53.4	-	-
Siegal	BB328771	3.5	51.1	58.8	67	36.6
	BB328772	2.6	58.8	54.5	62	37.4
Serco	BB328771	2.5	34	54	70	32
	BB328772	2.5	34	56	71	28

Table 12 Sulfate quality control experiment

Known concentration (mg/l)	analytically determined concentration (mg/l)
1	Not detected
2	2.5, 2.6
5	6.8, 7.8, 7.8
7	8.8

Alkalinity of surface water in northeastern Minnesota will often change in time probably because of the breakdown of humic and fulvic acids typical of these waters (Malcolm, 1978)^A. ^{or equilibration with atmospheric CO₂ if the titration is prolonged} Consequently, alkalinity should be ^{rapidly} measured soon after sample collection in the field or results may be anomalously high.

Ground water samples were collected by using a peristaltic pump connected to a pre-rinsed acid-washed Millipore filter apparatus equipped with Nucleopore 0.45 micron acid washed filters. Fifty milliliters of sample were filtered in the field and acidified by the addition of one milliliter of ultra pure nitric acid. In addition, up to one liter of unfiltered sample was collected for anion analyses.

Precipitation was collected in acid-washed, one liter, high density polyethylene bottles equipped with two plastic funnels separated by Whatman #40 filter paper previously washed in acid and de-ionized water. This filter paper was used to prohibit particulate material and insects from entering the bottle.

Surface infiltration water samples were collected by inserting a PVC plastic sheet, one quarter inch thick, 30 inches long and six inches wide between the A and B-soil horizons. Attached to the sheet was a PVC trough which

funneled infiltrating water into an acid washed polyethylene bottle. Both the trough and bottle opening were covered by nylon mesh to keep out large particulate material. The design of the apparatus was similar to that described by Wright (1977).

All water quality samples were chilled at about 40°F during transportation from the field and in subsequent laboratory storage. Twenty-five milliliters of each stream, precipitation and infiltration sample was filtered and acidified in the laboratory using the same techniques outlined for the collection of the ground water samples.

The results of water quality analyses are given in Appendix 5 .

Results

From March, 1977 to February, 1978, concentrations of dissolved calcium, magnesium, sodium and potassium, pH, sulfate, bicarbonate, specific conductance and silica were remarkably similar at the sampling locations on Filson Creek and at Omaday Lake (Table 13). Tests of significance between mean concentrations at the F1X location and at other locations on the creek and Omaday Lake generally show no difference at the 95 percentile level of significance^(Appendix 6). Only the mean chloride and total potassium concentrations are significantly^{differently} both being

Table 13 Summary statistics for water quality of the ... from 12-4-76 to 1-18-77. Concentrations in mg/l except As and Cu (µg/l), Cl, SO₄ (mg/l), specific conductance (µmhos). "D" is dissolved, "T" is total. Cu and Ni values are for total metal.

FLX LOCATION
TOTAL N 25

VARIABLE	VALID N	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	21	6.11	.57	6.60	5.00
HCO3	21	7.28	3.30	13.00	2.00
SO4	21	2.33	3.01	13.00	0
CL	25	.45	.53	2.00	.40
TCA	22	3.72	1.01	5.60	2.10
DCA	16	2.81	.85	4.20	1.50
TMG	22	2.47	.80	3.80	1.60
DMG	20	1.56	.48	2.60	1.10
TNA	24	1.42	.35	2.50	1.10
DNA	22	1.30	.26	1.80	1.00
TK	25	.51	.28	1.20	.20
DK	22	.45	.23	1.00	.20
SILICA	21	8.78	4.26	18.00	1.00
SPCD	23	41.00	8.26	60.00	26.00
CU	15	6.44	1.77	9.00	1.75
NI	14	4.65	1.64	7.50	0.00

F3X LOCATION
TOTAL N 18

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	17		6.12	.27	6.58	5.50
HCO3	17		6.53	6.97	32.00	2.00
SO4	18		1.67	1.78	7.00	0
CL	18		1.28	.49	2.20	.50
TCA	18		3.30	.99	6.00	2.00
DCA	13		2.62	.72	4.00	1.80
TMG	17		2.19	.91	5.40	1.50
DMG	13		1.55	.57	2.10	1.00
TNA	18		1.40	.40	2.50	1.00
DNA	13		1.21	.22	1.70	1.00
TK	18		.70	.24	1.10	.30
DK	13		.62	.17	1.00	.30
SILICA	18		7.21	4.06	15.20	.80
SPCD	16		38.31	7.98	51.00	21.00
CU	9		4.64	1.88	6.80	0.00
NI	8		2.70	1.07	8.20	0.00

OL-2 LOCATION
TOTAL N 17

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	17		5.88	.28	6.53	5.50
HCO3	17		8.29	12.52	56.00	3.00
SO4	16		2.04	3.27	10.00	0
CL	17		1.29	.55	2.70	.40
TCA	17		3.73	.89	6.00	2.50
DCA	14		2.59	.60	3.60	1.70
TMG	17		2.58	1.12	6.20	1.50
DMG	14		1.61	.42	2.50	1.10
TNA	17		1.53	.75	4.20	1.10
DNA	14		1.24	.26	1.80	1.00
TK	17		.72	.40	1.40	.30
DK	14		.64	.29	1.50	.30
SILICA	17		7.22	3.38	14.70	1.20
SPCD	15		42.53	14.01	62.00	23.00
CU	8		2.15	1.77	5.70	.80
NI	8		1.31	.90	4.50	0

Table 13 (continued)

OKADAY LAKE						
		TOTAL N	21			
VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	21		6.27	.31	6.95	5.60
HCO3	21		6.59	5.33	19.00	3.00
SO4	21		2.67	3.76	15.00	0
CL	21		1.28	.48	2.20	.50
TCA	21		3.97	2.05	8.00	2.00
DCA	13		2.69	.71	3.90	1.90
TMG	21		2.65	1.39	5.60	1.50
DNG	13		1.60	.31	2.20	1.20
TNA	21		1.70	.74	3.40	1.10
DNA	13		1.27	.23	1.80	1.00
TK	21		.61	.18	.90	.30
DK	13		.62	.16	.90	.30
SILICA	21		6.18	4.95	16.80	1.00
SPCD	19		39.26	7.56	50.00	21.00
CU	11		0.44	3.60	9.00	
NI	7		0.50	1.00	3.10	

QST LOCATION						
		TOTAL N	11			
VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	11		6.00	.18	6.29	5.62
HCO3	11		6.18	2.82	12.00	3.00
SO4	11		1.91	1.38	4.00	0
CL	11		1.19	.56	2.40	.50
TCA	10		3.22	.36	4.00	2.60
DCA	8		2.41	.60	3.40	1.70
TMG	10		2.27	.51	3.30	1.60
DNG	8		1.61	.39	2.30	1.20
TNA	10		1.30	.15	1.60	1.10
DNA	8		1.20	.09	1.30	1.10
TK	10		1.00	.78	3.10	.60
DK	8		.96	.70	2.90	.60
SILICA	11		6.34	2.72	12.10	2.90
SPCD	9		39.00	7.76	55.00	31.00
CU	7		1.27	1.48	4.50	.30
NI	7		.16	.42	1.10	0

SF LOCATION						
		TOTAL N	17			
VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	17		6.38	.17	6.70	6.10
HCO3	17		9.35	6.29	20.00	3.00
SO4	17		1.59	2.06	10.00	0
CL	17		1.25	.79	3.50	.40
TCA	17		3.48	1.02	6.00	2.00
DCA	12		2.58	.66	3.90	2.00
TMG	16		2.53	1.04	5.50	1.70
DNG	11		1.64	.34	2.10	1.00
TNA	17		1.58	.67	3.80	1.10
DNA	11		1.17	.15	1.50	1.00
TK	17		.65	.45	2.10	.30
DK	11		.45	.09	.60	.30
SILICA	17		11.60	4.56	25.00	0.60
SPCD	14		41.51	12.27	64.00	27.00
CU	10		1.10	1.10	7.00	0.00
NI	10		.62	.70	3.10	0.00

a little lower at FlX than at other locations.

This difference is apparent rather than real because additional samples were collected at the FlX location during storm events in June, 1977 and not at the other locations.

Negative correlation coefficients between concentrations of major ions and discharge at the FlX location show fair concentration-discharge dependences (Table 14). When the additional samples at FlX are removed from the T-Test analysis, the means of chloride and total potassium concentration no longer are significantly different at the FlX location. T-Tests of significance indicate that the mean pH at FlX (6.1) is significantly different than that at OL-2 (5.8), which is immediately downstream from the wetlands at the mouth of Omaday Lake. The lower pH^{at OL-2} is typical of perched wetland streams which do not receive substantial amounts of their base flow from ground water sources (Verry, 1975).

Total concentrations of copper and nickel in 1977 generally increase from headwater locations, OL-1 and SF, to FlX near the mouth (fig. 19). Total nickel concentrations measured at QST and Omaday Lake were, except for one sample, less than one microgram per liter, while mean concentration at FlX and F3X were about 5 and 3 micrograms per liter respectively.

Table 14 Correlation coefficients between discharge and water chemistry at FlX location. Data from this study and Water Resources Data for Minnesota (1976)

Parameter	r	Number of samples	Significance
pH	-.14	37	.207
HCO ₃	-.56	39	.001
SO ₄	-.27	39	.049
Cl	-.30	43	.026
DCA	-.42	30	.001
TCA	-.49	19	.017
DMG	-.43	34	.006
TMG	-.46	19	.023
DNA	-.31	36	.030
TNA	-.56	21	.005
DK	-.01	36	.467
TK	-.29	22	.094
SILICA	+.22	39	.094
SP.COND.	-.37	41	.007

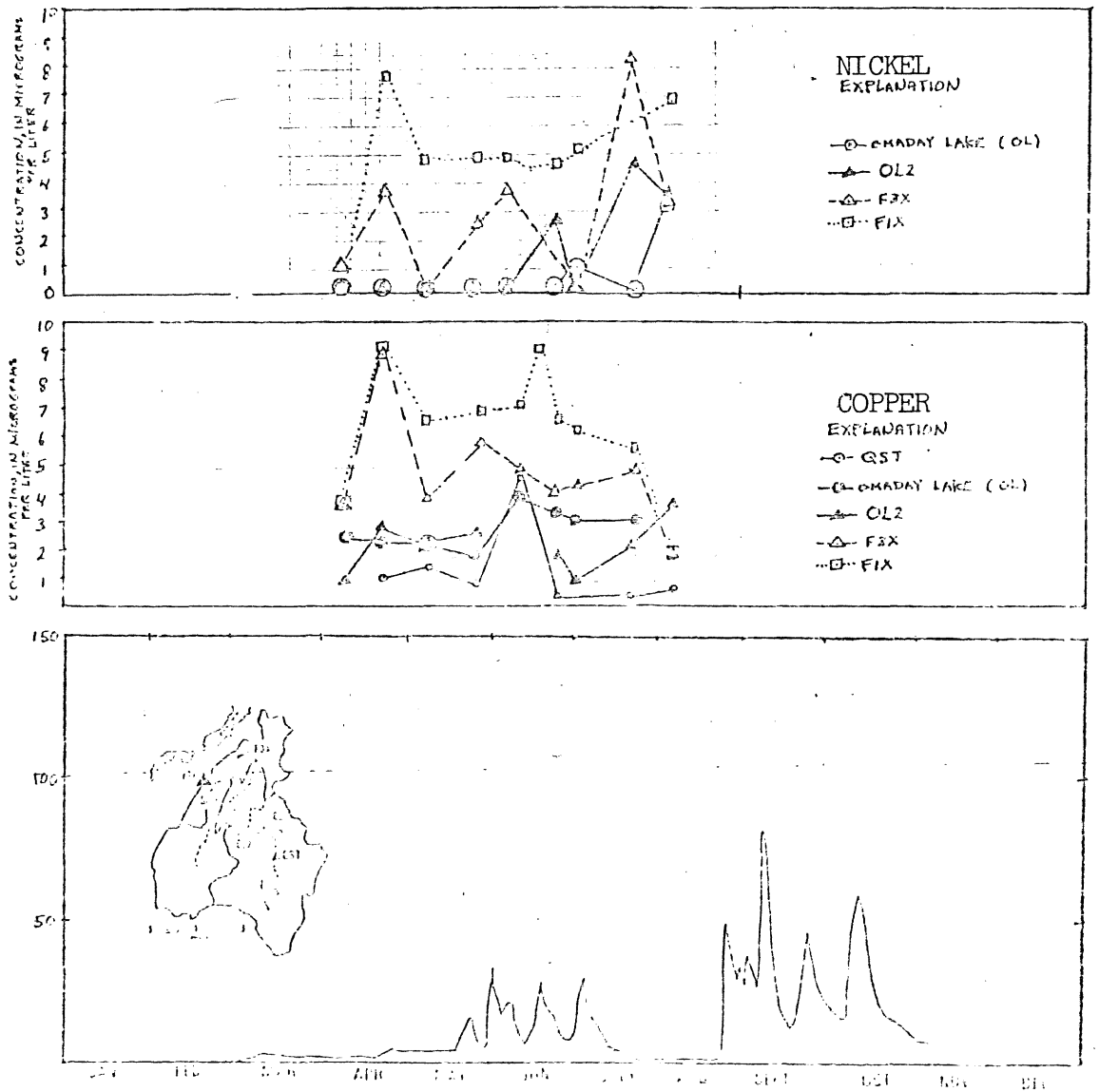


Figure 19 Concentrations of Copper and Nickel in Filson Creek

The smaller copper and nickel concentrations in Filson Creek headwater locations reflect the smaller percentage of sulfide bearing clases in the till and the greater distance of these locations from the contact zone.

Analyses of samples collected at FlX location during a runoff event in mid-June, 1977, show a slight decrease in copper and nickel concentrations during maximum discharge. Similar relationships are not present for other major cations (fig. 20). The observed decrease in copper and nickel is small and may in part reflect analytical limitations.

A comparison of mean concentrations of major cations Cl, HCO₃, pH, silica, and specific conductance for the calendar years 1976 and 1977 shows the difference between Filson Creekwater quality during respective dry and wet years. Concentrations of major cations and silica are higher in base flow conditions than in runoff events.

Comparison of dissolved cation concentrations over time

for samples collected at FlX during 1976 and 1977 show similarity in general trends.

(figure 21). During precipitation events, concentrations decrease.

Base flow in 1977 occurred only during early August, when concentrations of dissolved cations were very similar to concentrations in base flow

during 1976. Cation concentrations measured in baseflow during

January and February 1978 were also similar to that in 1976.

FIGURE 20 CONCENTRATION VARIATIONS FOR SELECTED CONSTITUENTS FOR STORM EVENT ON FEBRUARY 21-22, 1975

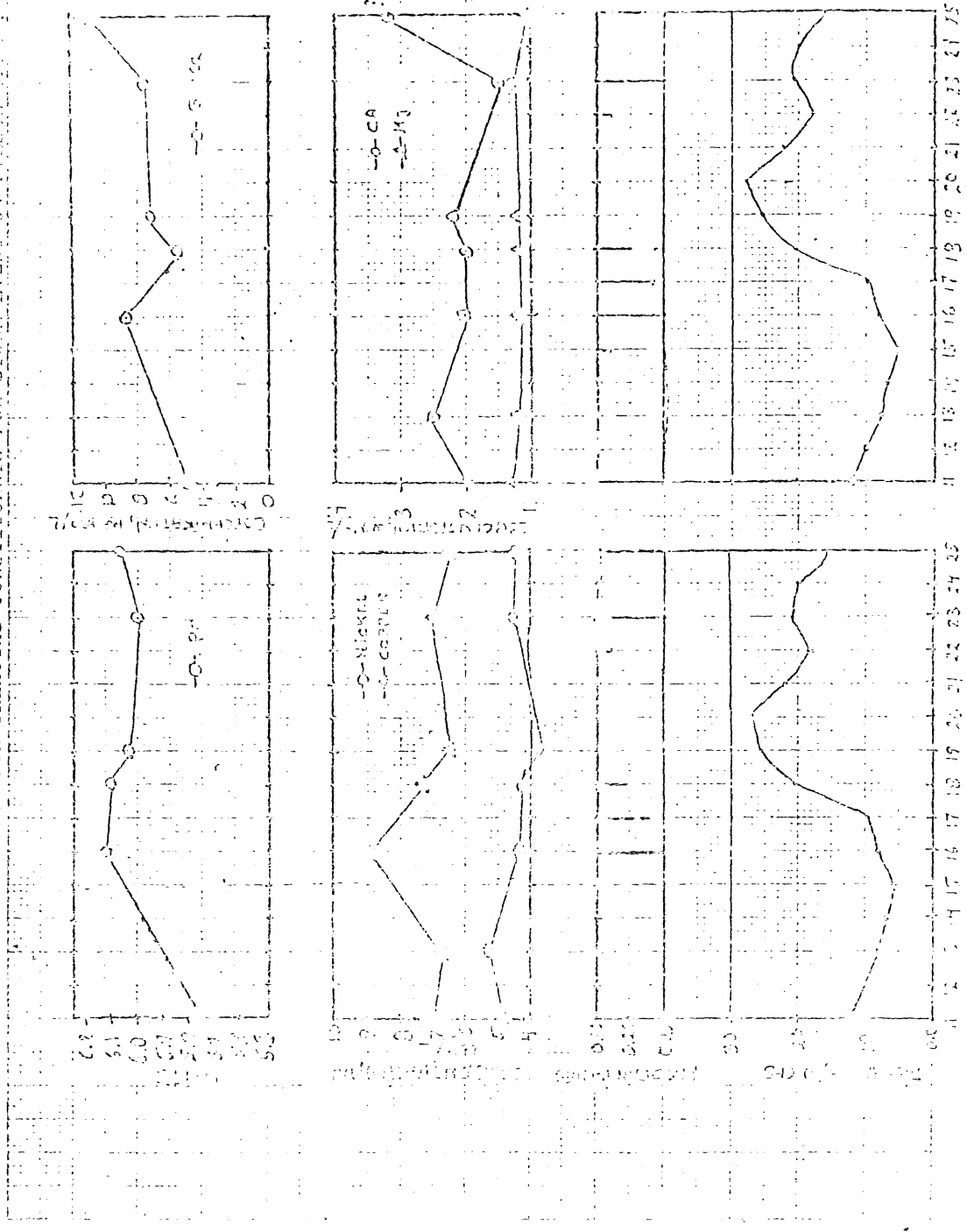


Table 15

COMPARISON BETWEEN MEAN CONCENTRATIONS AT FILSON CREEEK FOR
MAJOR CONSTITUENTS BETWEEN 1976 and 1977 (March to December)

	pH	Ca	Mg	Na	K	HCO ₃	SO ₄	Cl	Silica	Cu	Ni	[N]
1976 ^a	6.36	3.5	2.2	1.3	.4	16	5	1.6	12.7	8.7	6.3	13
1977 ^b	6.09	2.6	1.6	1.1	.5	8	2	0.9	8.6	6	4	[17]

^a U.S.G.S. (1977)^b Siegel, this report

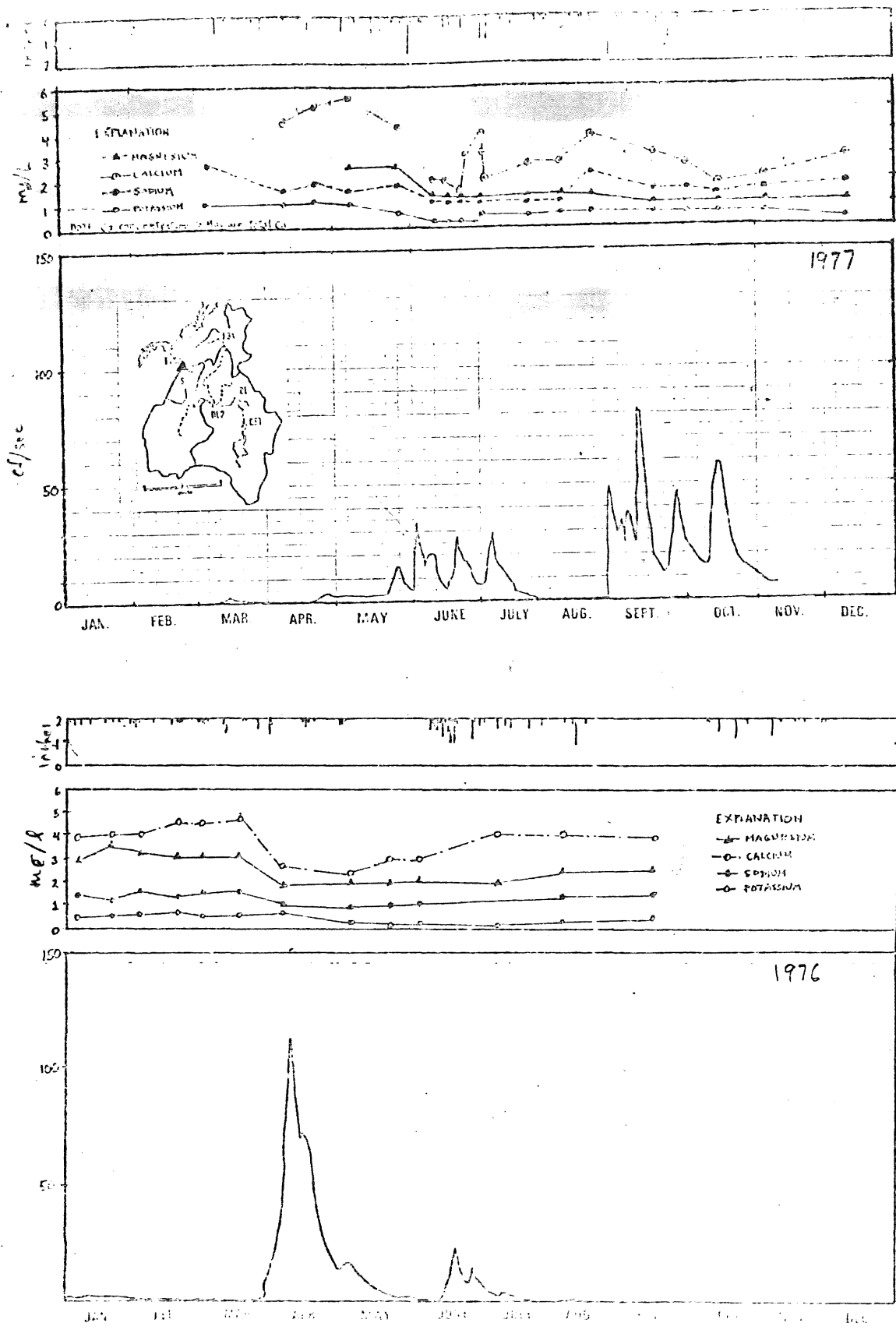


Figure 21 Variations in Cation Concentrations at FLX Location

Comparison between dissolved silica, sulfate and bicarbonate concentrations over time showed differences between 1976 and 1977 (fig. 22). Silica concentrations are depressed during spring and summer months probably due to diatom activity in the headwater lakes and bogs. During base flow winter conditions, silica concentrations rose in 1976 to about 10 mg/l, and in 1977 to a little over 15 mg/l.

Sulfate concentrations in 1976 fluctuated only slightly, while during 1977, a marked increase in sulfate occurred during snowmelt at all locations on Filson Creek as well as Omaday Lake. This increase is attributable to the sulfate load accumulated in the snowpack during the winter. Concurrent with the sulfate increase, pH dropped to its minimum value for the year (fig. 22a), reflecting the pH of the snowpack measured in March 1976 as 4.7. Similar increase in sulfate was not observed in 1976. The lack of sulfate fluctuation in 1976 contrasts with the rapid drop in sulfate concentration to nearly non-detectable limits after the snowmelt in 1976, and conceivably could be related to analytical difficulties mentioned earlier in the report.

Bicarbonate concentrations in 1977 clearly reflect dilution of baseflow by precipitation having little or no alkalinity. Similar dilution by spring snowmelt also occurred in 1976, but not with the same magnitude.

The major dissolved cations in precipitation are calcium and potassium. The major anion is sulfate. Summary statistics of the concentrations of major cations and anions, silica and specific conductance from 16 samples of bulk precipitation (wet and dry) collected from snow and three locations in or near Filson Creek Watershed are given in Table 16.

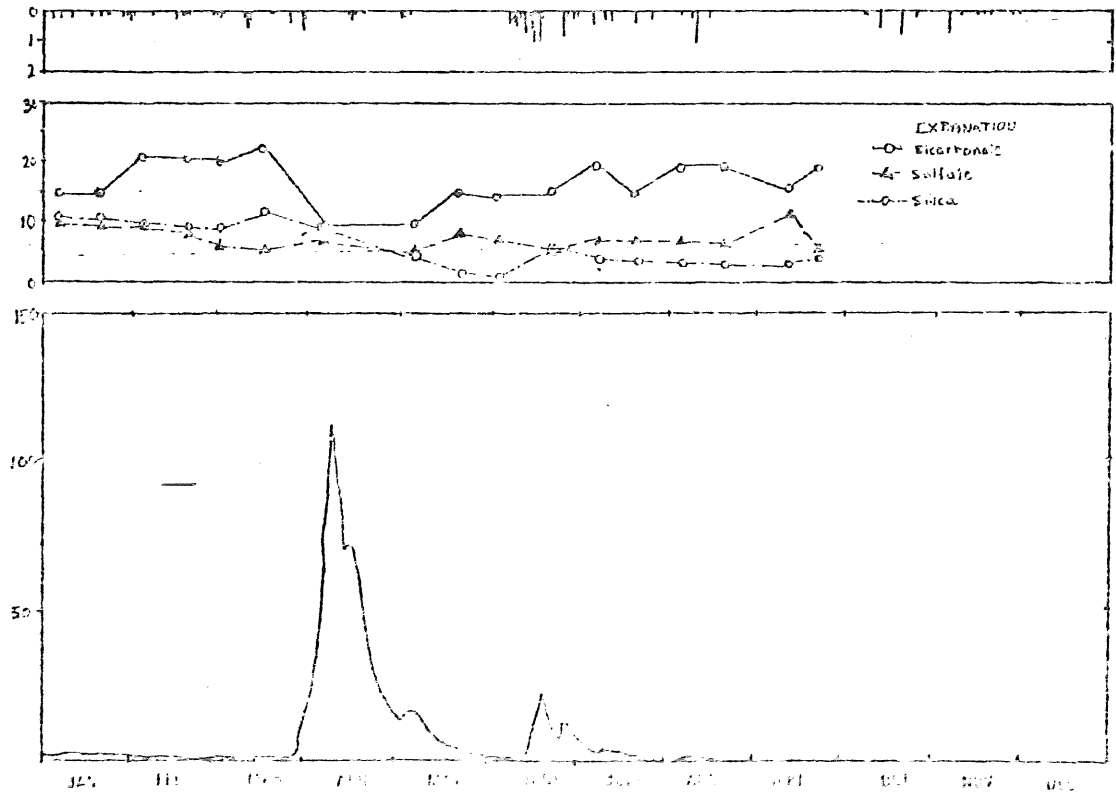
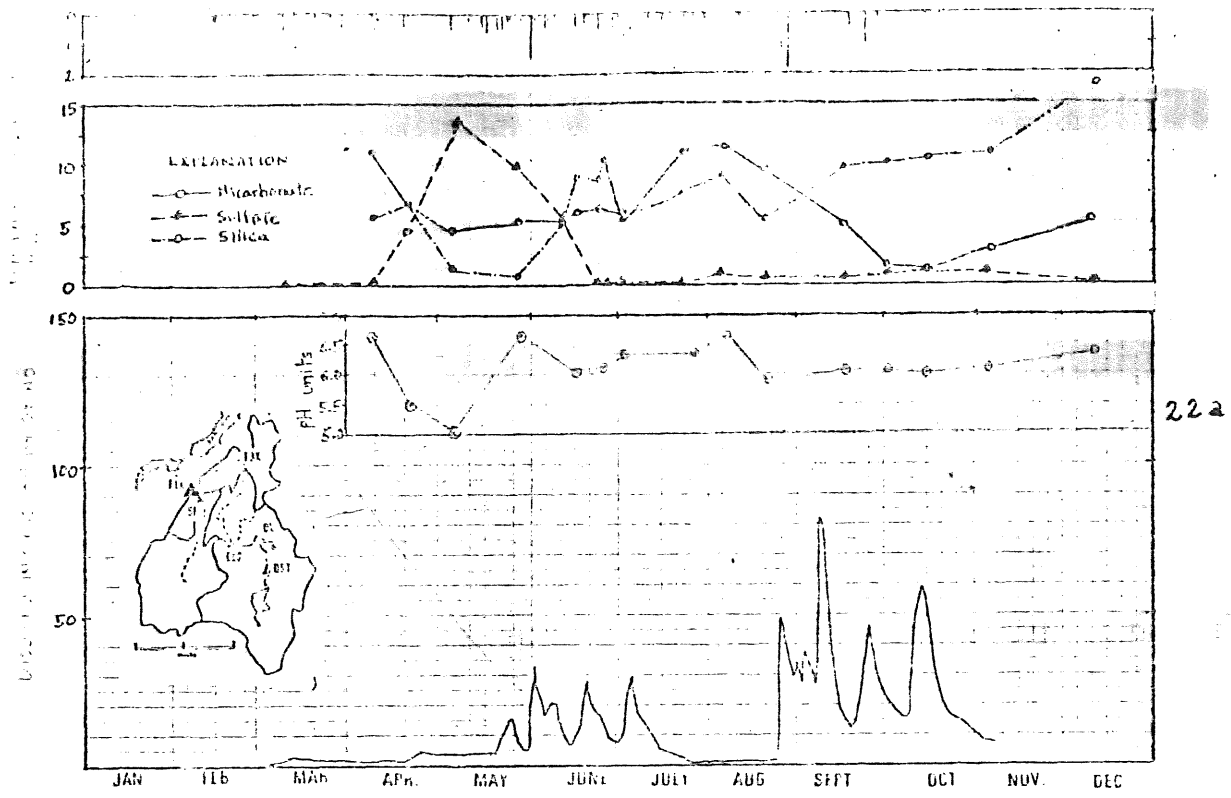


Figure 22 Variations in Anion Concentrations at FLX Location

Table 16 Summary statistics for precipitation water quality collected
in and around Filson Creek Watershed, 1977

VARIABLE	VALID N	TOTAL N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	14	16		5.06	.48	6.30	4.50
HCO3	16			.06	.25	1.00	0
SO4	15			1.84	1.45	5.80	0
CL	14			.52	.26	1.20	.30
TCA	16			.92	.37	1.50	.30
DCA	15			.68	.27	1.30	.30
TMG	16			.09	.09	.40	0
DMG	16			.01	.03	.10	0
TNA	16			.28	.29	1.00	0
DNA	16			.15	.18	.60	0
TK	16			.50	.37	1.40	0
DK	16			.36	.24	.90	0
SILICA	16			.14	.26	.90	0
SPCD	12		19.00	6.70	6.70	34.00	7.00
CU	16		0	0	0	-0	-0
NI	16		0	0	0	-0	-0

Because of vandalism, the collector at HLP location was discontinued in late July, 1977. Specific conductance was used as a measure of particulate contamination not removed by the filter paper in the collectors. Samples were deleted from statistical analysis if the specific conductance exceeded 35 micromhos. Analyses of those samples with greater than 35 umhos conductance showed anomalously high sodium or potassium concentrations. The source of the contamination probably was related to insects which were caught in the filter paper.

Atmospheric mean loading rates derived from bulk precipitation (Table 17) in the Filson Creek Watershed compare favorably with similar data collected by ~~Regional Cu-Ni~~
Study Group (1978).

Surface interflow water was collected at two locations at the base of ~~two~~ stands of aspen and birch within the mixed upland forest community, three locations at the base of an assemblage of mixed conifers, aspen and birch, and one location at the base of a stand of Jack Pine. Although most large particulate material was removed by the nylon mesh covering the apparatus, silt and clay size particles

Table 17 Atmospheric mean loading rates derived from
 bulk precipitation in the Filson Creek Watershed Area-1977
 (all parameters in kg/hect/yr)

	Siegel, this study	Eisenreich (1978) (for Regional Air Quality Group)
Ca	4.9	4.1
Mg	1.7	1.7
Na	1.9	1.9
K	2.4	1.6
SO ₄	16	14
Cl	4.4	7.4

and organic colloidal material were found in approximately 90 percent of the sample bottles at the time of collection. The measured water quality consequently is a composite of interflow water quality and additional leaching from this detritus.

Summary statistics of the interflow water quality are given in Table 18. Mean concentrations of potassium are up to seven times greater in interflow water than in Filson Creek or Omaday Lake water. Dissolved sodium is comparable to that found in precipitation, as are concentrations of chloride and sulfate. Dissolved calcium and magnesium are greater here than in precipitation, but less than in Filson Creek. The occurrence of bicarbonate, not found in precipitation reflects decomposition of organic detritus. Dissolved silica found in the interflow water is probably the result of leaching from small mineral or vegetation particles in detritus found in the bottles.

Ground water quality in the wetlands within Filson Creek Watershed appears to vary according to proximity to the mineralized zone. Ground water collected from 2XW, immediately over the contact zone, is classified as a Ca-Mg-SO₄ water compared to the Ca-Mg-HCO₃ classification of ground water

TABLE 18 SUMMARY STATISTICS FOR ELEMENTS COLLECTED AT BASE OF NATURAL JACK PINE STAND

COLLECTED AT BASE OF NATURAL JACK PINE STAND

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	5		5.57	.55	5.93	4.50
HCO3	6		2.13	2.24	6.20	0
SO4	5		.48	.55	1.00	0
CL	5		.65	.34	1.20	.30
TCA	7		5.04	6.42	19.60	1.70
DCA	7		3.23	3.32	11.70	1.50
TMG	7		.81	.57	2.00	.30
DMG	7		.43	.14	.70	.20
TMA	7		.44	1.04	2.60	0
DMA	7		.03	.05	.10	0
TK	7		4.21	1.42	6.30	2.30
DK	7		4.16	1.42	6.30	2.30
SILICA	4		1.45	1.15	2.60	0
SPCD	7		0	0	0	0
CU	7		0	0	0	0
NI	7		0	0	0	0
DIFF	5		.14	.05	.18	.06

COLLECTED AT BASE OF ASPEN-FIRCH STANDS (SWR1, SWR3)

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	6		5.88	.57	6.51	5.36
HCO3	8		.78	.83	2.00	0
SO4	9		.84	1.14	3.30	0
CL	9		.80	.43	1.60	.30
TCA	6		2.40	1.48	4.60	.80
DCA	9		1.76	1.32	4.00	.70
TMG	6		.83	.63	2.00	.40
DMG	9		.38	.52	1.60	0
TMA	9		.29	.59	1.60	0
DMA	9		0	0	0	0
TK	8		6.89	8.02	20.00	1.30
DK	9		6.20	7.49	20.00	1.30
SILICA	7		.60	.77	2.10	0
SPCD	9		0	0	0	0
CU	9		0	0	0	0
NI	9		0	0	0	0
DIFF	8		.19	.24	.74	.03

COLLECTED AT BASE OF MIXED UPLAND FOREST STAND (SWR1-1 SWR1-2 SWR1)

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	8		5.65	.76	6.37	4.10
HCO3	8		3.95	2.93	7.00	0
SO4	9		.11	.33	1.00	0
CL	8		.87	.35	1.40	.30
TCA	11		3.44	1.69	6.20	1.40
DCA	12		3.27	2.00	8.40	1.20
TMG	11		.85	.44	1.80	0
DMG	12		.78	1.09	3.90	0
TMA	12		.73	1.36	3.75	0
DMA	12		0	0	0	0
TK	12		4.98	3.29	12.30	1.70
DK	12		5.55	3.65	12.30	0
SILICA	8		.71	.95	2.10	0
SPCD	12		0	0	0	0
CU	12		0	0	0	0
NI	12		0	0	0	0
DIFF	8		.27	.07	.21	.04

(Fig 23).

collected from the other piezometers. This difference is even more apparent, considering that ground water from piezometer 3XW, located only about 30 feet north of the contact, is classified as identical to that from QBW, which is farthest removed from the contact zone. Similar sulfate predominance is reported by Siegel and others (1978) for ground water collected from Rainy Lobe till along the contact zone. Summary statistics for ground water chemistry are given in Table 19.

Copper and nickel concentrations vary considerably in wetland ground water over time and between four piezometers sampled (Appendix 5). Not enough analyses are available to statistically correlate these variations with other chemical parameters. Arithmetic means, however, suggest that total nickel concentrations decrease with distance from the contact zone. No similar trend is apparent in the mean copper concentrations.

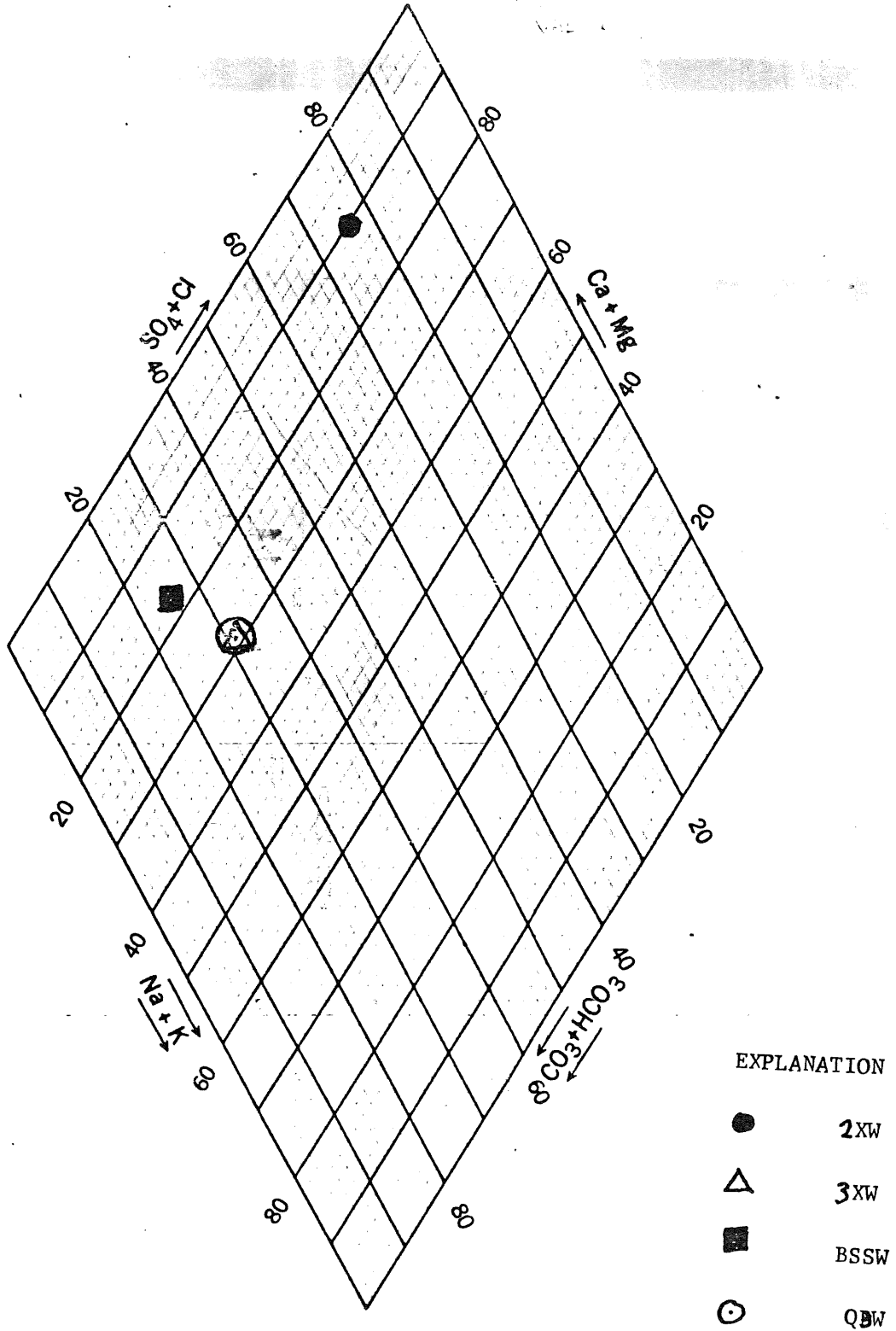


Figure 23 Piper plot of wetland water quality in Filson Creek Watershed. Values plotted are mean values for 1977 samples.

TABLE 19 SUMMARY STATISTICS FOR GROUNDWATER QUALITY IN WETLANDS IN THE
TILSON CREEK WATERSHED

Q1W TOTAL N 10

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	6		5.76	1.14	5.98	5.60
HCO3	9	31.00	23.81	71.00	11.00	
SU4	9	6.36	12.13	40.00	0	
CL	10	2.15	1.68	5.00	2.0	
DLA	10	10.71	5.46	20.00	3.50	
DNA	10	4.05	2.15	8.20	2.00	
DR	10	4.65	1.53	7.40	1.70	
SILICA	9	5.34	5.60	17.80	1.70	
CU	3	26.17	6.15	34.00	14.50	
NI	3	10.63	7.57	21.30	7.90	
	3	14.27	5.20	20.10	10.10	

Q2ND TOTAL N 6

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	6		4.78	2.35	6.08	-6
HCO3	5	29.60	16.01	45.00	-6	
SU4	3	6.09	5.00	11.00	1.00	
CL	6	.95	.67	2.10	-6	
DLA	5	18.30	20.23	54.40	7.50	
DNA	5	3.20	2.48	6.30	-6	
DR	5	3.32	2.28	6.40	-6	
SILICA	4	3.76	4.63	10.40	-6	
CU	5	12.20	6.63	20.20	-6	
NI	5	2.92	2.19	5.90	-6	
	5	20.14	20.60	60.50	4.00	

Q3W TOTAL N 12

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	12		6.14	.23	6.53	5.75
HCO3	10	51.00	24.16	93.00	11.00	
SU4	11	17.25	16.11	44.00	6	
CL	9	1.24	.44	2.00	.80	
DLA	11	10.73	2.48	15.00	8.00	
DNA	11	6.21	1.55	7.70	4.00	
DR	11	6.75	3.25	10.00	4.20	
SILICA	11	17.45	5.69	29.00	10.00	
CU	3	67.67	32.55	105.00	68.00	
NI	4	88.00	57.00	140.00	9.00	

Q4W TOTAL N 8

VARIABLE	VALID N	SUM	MEAN	ST-DEV	MAXIMUM	MINIMUM
PH	8		5.75	1.29	6.20	5.40
HCO3	6	19.17	9.59	35.00	9.0	
SU4	6	40.17	17.60	72.50	20.00	
CL	7	1.97	.44	1.40	2.0	
DLA	6	11.71	3.83	20.00	3.00	
DNA	7	7.77	2.00	11.40	6.10	
DR	6	4.00	2.00	4.50	1.0	
SILICA	6	1.83	2.00	1.70	2.50	
CU	3	10.00	7.20	30.00	17.00	
NI	3	10.00	7.20	30.00	17.00	
	2	10.00	7.20	30.00	17.00	

Chemical Budgets

The difference between the total chemical mass entering the watershed from precipitation and the amount leaving the watershed through discharge over a unit period of time comprises a chemical budget.

The yearly chemical budget for specific constituents can be expressed as:

$$\sum Q_p C_{p_i} - \sum Q_d C_{d_i} = \Delta M_i$$

where: \equiv Q_p = amount of monthly precipitation
 C_{p_i} = concentration of constituent "i" in precipitation
 Q_d = amount of monthly discharge
 C_{d_i} = concentration of constituent "i" in discharge
 ΔM_i = net gain or loss of the constituent from
the watershed

Assuming dynamic equilibrium of the biomass with respect to dissolved calcium, magnesium and sodium, the budgets provide a qualitative index on the relative rates of weathering of the major silicate minerals in the watershed. Because chloride is chemically conservative, a net balance for the year will verify the previously calculated hydrologic budget.

The yearly budget for the major anions and cations for Filson Creek Watershed are given in Table 20 and Figure 24. Monthly net loading on the watershed was obtained from the product of monthly precipitation measured at the Kawishiwi Laboratory and the monthly mean concentration data (if more than one sample were available) of the precipitation sampled around the watershed. It was assumed that snowmelt and precipitation during March and April, 1977 went into storage. During August, the major precipitation events occurred toward the end of the month and extended into September. Because of this, the chemical quality data collected at mid-September were considered representative of precipitation recorded between August 20 and 31. Similarly, quality data from October 1st precipitation were considered representative of precipitation that fell at the end of September.

The results of the yearly budget (minus winter 1978 for which discharge data at FlX are not yet available) indicate that

- 1). Chloride input balances output by about 10%.
- 2). There is a net loss from the watershed of calcium, magnesium, sodium, silica and bicarbonate.
- 3). There is a net gain in the watershed of potassium and sulfate.

When the additional discharge data for winter 1978 are tabulated, the chloride balance will be a little below 10% and net losses will be a little larger.

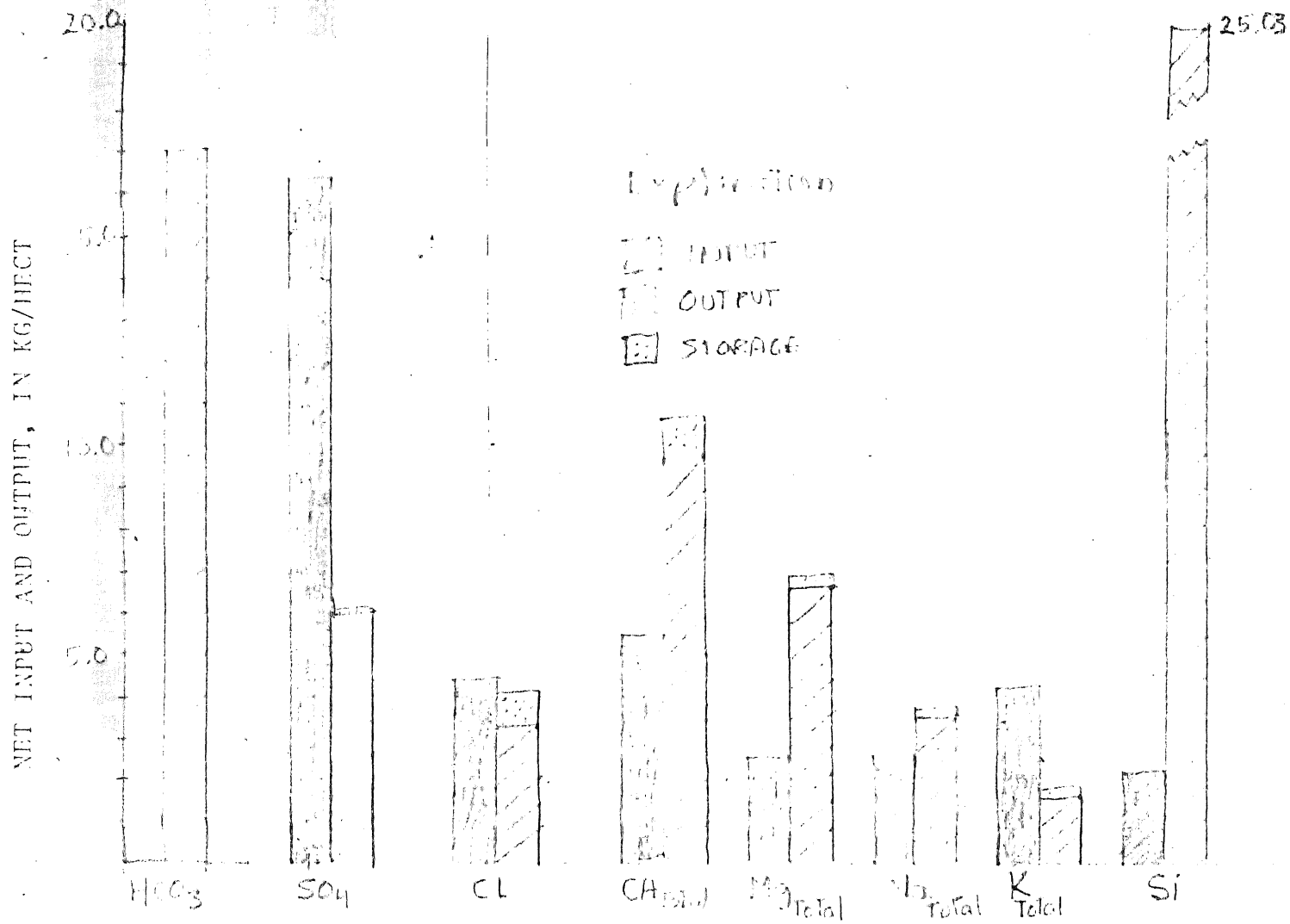


Figure 24 Bar graph showing net gains and losses of major cations and anions for Filson Creek watershed, March-October, 1977

TABLE 20 Net chemical budgets for Filson Creek Watershed through March to October 1977. Snowpack and precipitation in March and April assumed to effectively have gone into storage

	Mass, in Kg/Hectare											
	HCO ₃	SO ₄	Cl	Ca _t	Ca _d	Mg _t	Mg _d	Na _t	Na _d	K _t	K _d	Si
INPUT	0.0	15.0	3.7	4.9	4.9	2.7	1.6	2.4	1.7	3.8	2.4	2.1
OUTPUT	17.2	6.3	3.3	9.8	7.1	6.6	4.4	3.5	3.3	1.5	1.5	25.0
NET LOSS OR GAIN	-17.2	+8.7	+0.4	- 5.1	-2.2	-3.9	-2.7	-1.1	-1.6	2.3	0.9	-22.9

APPENDICES

APPENDIX 1

Percentage of Mineral Species in 40 Size Fraction of Till Collected from Filson Creek Watershed

	Albite	Microcline	Orthopyroxene Hypersthene	Orthopyroxene Enstatite	Biotite	Hornblende	Microcline	Quartz	Opal	Number of specimens		
	2.7	9.1	4.4	1.3	0.7	0.5	39.2	2.4	0.5	3.1	4.8	580
	3.0	17.2	4.3	0.4	1.2	1.6	16.2	3.6	---	1.6	10.9	501
	47.9	23.0	1.8	0.7	---	---	4.8	4.5	---	12.2	5.7	451
	43.0	11.0	3.1	0.2	4.3	1.6	19.5	2.1	---	4.1	5.5	417
	37.3	34.0	0.4	1.5	14.8	1.5	0.9	9.9	---	4.4	4.6	454
	51.3	20.1	3.3	0.7	7.0	---	4.4	1.7	---	1.0	5.2	452
	49.	19.4	2.0	0.6	4.1	---	10.8	3.5	---	2.9	5.5	457
	51.3	11.1	1.0	0.8	6.0	---	22.0	0.7	---	1.8	4.7	457
	43.0	16.5	0.6	---	5.0	---	5.0	7.5	---	2.8	3.0	464
	51.3	15.1	3.5	---	3.8	---	10.8	6.1	---	5.2	4.0	485
	34.5	31.0	1.5	0.6	4.0	---	10.9	4.4	1.9	1.9	11.5	478

APPENDIX 2

Accumulated precipitation at gages in vicinity of Filson Creek Watershed during water year 1976. (All values in inches of water).

Time Period	Recording Gage at So. Kawishiwi Field Lab	Non-Recording Gages on Filson Creek Watershed			Average for Filson Creek Watershed
		Lower Filson	Upper Filson	Evaporation Station	
9/30/75 to 10/14/75	0.4	0.47	0.65	-	0.56
10/14 to 11/3	1.7	1.25	1.12	-	1.19
11/3 to 12/2	2.6	3.00	3.25	-	3.13
12/2 to 12/31	0.9	0.73	0.72	-	.73
12/31/75 to 1/30/76	1.6	1.51	1.47	-	1.49
1/30 to 2/26	0.7	0.73	0.65	-	0.69
2/26 to 3/24	2.1	1.82	1.60	-	1.71
3/24 to 4/30	2.0	2.13	2.48	-	2.31
4/30 to 5/17	0.7	1.08	1.00	1.04	1.04
5/17 to 6/16	3.4	3.24	3.67	3.52	3.48
6/16 to 7/1	2.5	2.72	3.14	2.45	2.77
7/1 to 7/14	1.2	1.15	1.13	1.47	1.25
7/14 to 8/16	2.4	2.37	1.97	2.22	2.19
8/16 to 9/23	1.4	1.34	1.25	1.35	1.31
9/23 to 10/1/76	0.3	0.21	0.28	0.30	0.26
Total precipitation Recorded for Water Year 1976 (in inches of water)	23.9	23.75	24.38	-	24.11

Accumulated precipitation at gages in vicinity of Filson Creek watershed during water year 1977 (All values in inches of water).

Time Period	Recording Gage at St. Kawishiwi Field Lab	Non-Recording Gages on Filson Creek Watershed		
		Lower Filson	Upper Filson	Evaporation Station
10/21 to 10/29	1.2	1.0	1.1	0.5
10/29 to 11/30	0.4	0.4	—	0.4
11/30 to 12/30	0.6	0.7	—	0.7
12/30 to 1/31	0.6	0.7	1.4	0.7
1/22 to 3/14	1.6	1.5	—	1.5
3/14 to 4/15	1.1	1.1	2.9	1.1
4/15 to 5/5	1.0	1.8	—	1.5
5/5 to 5/16	0.8	0.2	2.7	0.5
5/16 to 6/1	3.8	3.7	3.5	4.6
6/1 to 6/30	2.4	3.6	—	2.6
6/20 to 6/30	1.6	1.3	4.5	1.4
6/30 to 7/6	2.6	2.1	—	2.5
7/6 to 8/1	---	1.6	3.2	1.3
8/1 to 8/15	0.2	---	—	0.2
8/15 to 9/1	5.7	---	—	5.4
9/1 to 9/15	—	---	8.9	3.9
9/15 to 9/30	6.2	---	2.7	2.7
Total precipitation Recorded for Water Year 1977 (in inches of water:	29.8	---	30.9	31.5

APPENDIX 3 DAILY MEAN DISCHARGES AT F1X, F3X,
AND OL-2 ON FILSON CREEK, 1975-1977

STATION NUMBER
LATITUDE 475005

05124000
LONGITUDE 091427

FILSON CREEK NEAR ILY, MISSOURI
DRAINAGE AREA

9.002
DAILY

STATE 21
COUNTY 077

DISCHARGE, IN CUBIC FEET PER SECOND, WATER YEAR OCTOBER 1976 TO SEPTEMBER 1977
MEAN VALUES

DAY	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP
1	.04	.04	.08	.00	.00	.00	.00	3.1	34	8.5	1.2	35
2	.06	.12	.07	.00	.00	.00	.00	3.1	30	8.5	1.2	30
3	.06	.15	.06	.00	.00	.00	.00	3.1	24	15	1.2	31
4	.11	.15	.06	.00	.00	.00	.00	2.6	21	17	1.2	34
5	.12	.15	.10	.00	.00	.00	.00	2.6	11	17	1.2	37
6	.22	.15	.15	.00	.00	.00	.22	2.2	24	30	.71	37
7	.12	.18	.20	.00	.00	.01	.22	3.5	20	21	.82	38
8	.02	.15	.18	.00	.00	.01	.31	1.5	20	19	.65	37
9	.02	.15	.18	.00	.00	.01	.42	2.5	18	19	.65	37
10	.02	.24	.15	.00	.00	.01	1.1	2.8	14	19	.60	37
11	.02	.24	.15	.00	.00	.00	1.1	2.5	12	10	.50	37
12	.02	.22	.15	.00	.00	.50	.57	2.5	10	19.5	.07	38
13	.00	.22	.15	.00	.00	2.0	.70	2.5	10	8.5	.07	37
14	.06	.22	.15	.00	.00	4.0	.82	2.5	7.1	8.5	.07	37
15	.06	.22	.15	.00	.00	3.0	.82	2.5	5.3	8.0	.07	37
16	.06	.18	.20	.00	.00	2.0	.82	3.0	6.0	8.0	.07	37
17	.06	.18	.25	.00	.00	2.0	.82	3.0	6.0	8.0	.07	37
18	.04	.22	.28	.00	.00	1.5	.71	2.8	9.5	8.5	.07	37
19	.02	.24	.25	.00	.00	1.0	.80	2.4	20	8.5	.07	37
20	.02	.24	.20	.00	.00	.00	1.1	2.5	27	8.1	.07	37
21	.04	.22	.15	.00	.00	.70	2.5	6.4	22	2.5	1.1	37
22	.12	.22	.06	.00	.00	1.0	3.8	9.5	13	2.1	1.1	37
23	.12	.22	.04	.00	.00	.58	4.3	15	21	2.1	1.1	37
24	.12	.22	.04	.00	.00	.50	5.3	15	20	1.0	1.1	37
25	.12	.22	.01	.00	.00	.80	4.5	12	12	1.0	1.1	37
26	.09	.22	.00	.00	.00	.70	4.1	10	12	1.0	1.1	37
27	.06	.18	.00	.00	.00	1.0	3.5	8.4	10	1.3	1.1	37
28	.09	.18	.00	.00	.00	1.5	4.1	7.0	9.5	1.2	1.1	37
29	.09	.16	.00	.00	.00	1.8	3.8	6.0	8.0	1.2	1.1	37
30	.09	.09	.00	.00	.00	.87	3.3	5.6	8.0	1.2	1.1	37
31	.04	.00	.00	.00	.00	.94	7.6	7.6	8.0	1.2	1.1	37
TOTAL	2.65	5.50	3.42	.00	.00	27.00	51.92	163.0	511.4	253.8	171.39	980.0
MEAN	.084	.176	.110	.000	.000	.87	1.673	5.255	16.7	8.10	5.25	31.1
MAX	.02	.24	.26	.00	.00	4.0	5.3	15	21	2.1	1.1	37
MIN	.02	.04	.00	.00	.00	.00	.22	2.5	5.3	1.0	1.1	37
CFSM	.01	.02	.01	.000	.000	.10	.49	1.82	1.82	1.01	.61	3.11
STDEV	.01	.02	.01	.00	.00	.11	.21	.67	2.07	1.05	.73	3.11
CAL YR 1976	TOTAL	2069.31	MEAN	5.65	MIN	.00	CFSM	.87	TO	8.57		
WTR YR 1977	TOTAL	2169.22	MEAN	5.04	MIN	.00	CFSM	.84	TO	8.45		

1100900
 GREEN CREEK NEAR ILY, 1970

1970, COLLECTED BY J. D. DAVIS
 DATA REVISION 11

PERMISSION DATA FOR WATER YEAR FROM 5:01, 7:01, 1970 - 1971 (1)

DATE	MAX GH (TIME)	MIN GH (TIME)	MEAN GH	TOTIV GH	PER CENT	SHIFT AT AUG - R	DATUM AT CURR - R
10-15	6.10 (2115)	5.56 (1600)	5.77	5.56	57		
10-16	6.11 (2130)	5.60 (0615)	5.86	5.60	58		
10-18	6.09 (0615)	5.60 (2115)	5.86	5.60	58		
10-18	6.11 (0615)	5.60 (2130)	5.84	5.60	57		
10-17	5.70 (0615)	5.25 (2130)	5.60	5.60	58		
10-18	5.71 (2130)	5.60 (0615)	5.60	5.60	58		
10-18	5.60 (0615)	5.25 (2130)	5.60	5.25	51		
10-18	5.61 (0615)	5.25 (2130)	5.60	5.25	51		
10-21	5.55 (0615)	5.25 (2130)	5.51	5.25	49		
10-22	5.47 (0615)	5.25 (2130)	5.37	5.25	47		
10-22	5.46 (0615)	5.25 (2130)	5.37	5.25	48		
10-23	5.40 (0615)	5.25 (2130)	5.31	5.25	47		
10-24	5.40 (0615)	5.30 (2130)	5.31	5.30	48		
10-26	5.38 (0615)	5.30 (2130)	5.33	5.30	49		
10-27	5.34 (0615)	5.35 (2130)	5.37	5.37	49		
10-28	5.35 (0615)	5.32 (0645)	5.35	5.35	49		
10-29	5.37 (0615)	5.31 (0645)	5.31	5.31	49		
10-29	5.31 (0615)	5.29 (0630)	5.29	5.29	49		
10-31	5.29 (0615)	5.28 (0615)	5.29	5.28	49		
11-01	5.28 (0615)	5.20 (0615)	5.26	5.20	49		
11-02	5.28 (0615)	5.25 (0615)	5.26	5.20	49		
11-03	5.25 (0615)	5.25 (0615)	5.25	5.25	49		
11-03	5.25 (0615)	5.25 (0615)	5.25	5.25	49		

LOCATION-- NEARBY, sec. 17, T20N, R10W, Lake County, on east bank 0.1 mile upstream from where
 Spruce Road crosses Tilton Creek, designated F3X
 PERIOD OF RECORD-- April 1977 to November 1977

GAGE-- Water-stage-recorder. Altitude of the gage is 1500 ft., from topographic map

WATER-- Records good to fair

REMARKS FOR PERIOD OF RECORD-- Maximum discharge, 33 cfs. Sept. 9

DISCHARGE, IN CUBIC FEET PER SECOND, APRIL 1977 to NOVEMBER 1977
 MEAN VALUES

DAY	APR	MAY	JUNE	JULY	AUG	SEPT	OCT	NOV
1	.51	1.4	14	3.6	.91	14	12	4.2
2	.33	1.4	12	3.6	.91	12	11	3.8
3	.21	1.4	9.9	6.0	.57	9.8	9.0	3.6
4	.31	1.4	8.6	6.7	.54	15	7.8	3.6
5	.20	1.4	7.4	7.5	.54	15	7.4	3.5
6	.22	1.7	9.8	11	.72	13	7.0	3.4
7	.28	1.7	8.2	9.0	.72	12	6.2	3.3
8	.21	1.7	8.2	6.7	.72	11	6.2	3.2
9	.24	1.6	7.4	5.3	.78	23	6.2	3.6
10	.62	1.3	5.8	4.4	.78	32	6.2	4.6
11	.63	1.2	5.0	4.0	.78	25	6.2	
12	.54	1.2	4.2	3.7	.78	19	12	
13	.47	1.2	3.4	3.7	.78	15	19	
14	.55	1.2	3.0	3.4	.78	12	20	
15	.58	1.2	2.3	2.8	.78	9.8	23	
16	.59	1.4	3.4	2.4	.78	6.0	28	
17	.41	1.3	4.0	2.0	.78	5.3	19	
18	.47	1.2	8.6	2.0	.78	5.3	15	
19	.36	1.3	11	1.7	.78	5.1	13	
20	.63	1.9	11	1.4	.78	4.7	9.9	
21	1.2	2.9	9	1.4	.75	4.3	11	
22	1.7	4.2	7.4	1.3	.53	4.3	9.8	
23	1.9	6.2	8.6	1.2	.50	13	8.2	
24	2.3	6.2	8.2	1.2	.45	17	7.4	
25	2.0	5.0	6.6	1.1	.50	19	7.0	
26	1.8	4.2	5.0	1.0	1.0	18	6.6	
27	2.0	3.7	4.2	1.1	6.2	17	5.8	
28	1.8	3.0	3.8	1.1	19	17	5.0	
29	1.3	2.4	3.6	1.0	13	16	4.6	
30	1.9	2.0	3.4	1.1	11	15	4.6	
31		3.2	---	1.0	12	15	4.2	
TOTAL	20.22	61.9	207	103.4	79.9	430.6	318.3	

LOCATION-- NW 1/4 Sec. 25, T. 13 N., R. 12 W., Lake County, about 0.1 mile upstream from U.S.G.S. gage on Spruce Ecp1, designated SF

PERIOD OF RECORD-- April 1977 to November 1977

GAGE-- Water-stage recorder. Altitude of gage is 1445 ft. from topographic map

REMARKS-- Records fair to good, except after August 1977 when beaver dam created a backwater effect at gage location.

EXTREMES FOR PERIOD OF RECORD-- Maximum discharge 24 cfs on Sept. 9

DISCHARGE, IN CUBIC FEET PER SECOND, APRIL 1977 to NOVEMBER 1977
MEAN VALUES

DAY	APR	MAY	JUNE	JULY	AUG	SEPT	OCT	NOV
1	.22	1.1	2.8	3.6	.46	11	9.2	3.3
2	.15	1.1	6.3	3.6	.38	9.2	8.1	3.0
3	.13	1.1	7.8	4.8	.37	7.5	6.9	2.9
4	.13	1.1	6.9	5.4	.34	11	6.0	2.9
5	.11	1.1	5.8	8.3	.34	11.	5.7	2.8
6	.10	1.3	5.0	9.2	.28	10	5.4	2.7
7	.10	1.4	5.1	6.9	.25	8.9	4.8	2.6
8	.13	1.1	5.4	6.0	.22	8.3	4.8	2.6
9	.25	1.1	4.6	5.1	.22	24	4.8	2.8
10	.43	1.0	4.4		.22	23	4.8	3.6
11	.43	1.0	3.5		.22	18	5.9	
12	.34	1.0	2.6	3.2	.20	14	14	
13	.28	1.0	1.8	2.9	.20	11	15	
14	.17	1.10	1.8	2.1	.20	8.9	17	
15	.25	1.3	2.2	1.3	.15	7.5	17	
16	.20	1.6	2.9	1.3	.29	6.0	15	
17	.12	.80	1.1	1.4	.28	5.1	13	
18	.24	1.1	1.7	1.5	.28	4.5	11	
19	.17	1.5	8.1	1.5	.28	4.5	9.5	
20	.43	1.9	8.3	1.4	.33	4.2	8.3	
21	.21	2.5	6.9	1.4	.53	3.6	7.5	
22	1.4	3.9	5.7	1.6	.33	3.3	6.3	
23	1.5	3.9	6.6	.78	.31	3.0	5.7	
24	1.8	4.1	6.3	.71	.26	6.9	5.4	
25	1.6	4.2	4.9		.31	13	5.1	
26	1.5	4.0	4.6		.78	14	5.1	
27								
28	1.6	2.0	3.9	.50	4.8	13	4.5	
29	1.5	3.2	3.7	.46	14	12	3.9	
30	1.4	3.8	3.7	.46	10	12	3.6	
31	1.2	2.1	3.5	.46	8.1	11	3.6	
1	----	2.3	----	.48	8.9	----	3.3	
Total	19.08	57.6	137.9	84.5	53.8	299.4	29.2	

6

APPENDIX 4 Stratigraphic Logs of Piezometers Installed
in and around the Filson Creek Watershed

Location (1)	Well number (2)	Casing (2 inch plastic)				Materials penetrated		
		Total length, in feet (3)	Screen length, in feet (4)	Total depth below land surface, in feet (5)	Height above land surface, in feet (6)	Altitude at land surface, in feet (7)	Lithology (8)	Depth below land surface, in feet (9)
62-11-25bb	H1	17	None	13.5	3.5	1487.95	Silty clay, sandy to stoney, grey-green Silty clay, sandy to stoney, grey Bedrock	0-6 6-49.5 49.5
62-11-25bb	H2	9	3	6.5	2.5	1489.19	Fill, sandy Peat Silty clay, sandy to stoney, organic material Bedrock	0-3 3-4.5 4.5-6.5 6.5
62-11-25bb	H3	15	None	12	4	1487.07	Fill Silty clay, sandy to stoney, grey-green Bedrock	0-5 5-12 12
62-10-29cd	H29	16	3	13	3	1537 ⁺	Silty clay, sandy with some pebbles and organic material Bedrock	0-14 14

001

Casing (1½ inch PVC plastic)

Materials penetrated

Location	Well number	Total length, in feet	Screen length, in feet	Total depth below land surface, in feet	Lithology	Depth below land surface, in feet
NW¼ NE¼ SE¼, sec 25, T62N, R11W	BSSW	6	1	3	Fibrous brown to black peat, silty	0 - 3
NE¼ SE¼, sec 24, T62N, R11W	2XW	7.3	1	3	Fibrous brown to black peat, sandy Till, sandy	0 - 3.2 3.2 -
SE¼ NW ¼, sec 19, T62N, R11W	3XW	6.1	1	3	Fibrous brown to black peat, sandy. Sand, med. fine, brown	0 - 2.8 2.8 - 3.1
SE¼ NW¼ SE¼, sec 31, T62N, R11W	QBW	3.7	1	3	Fibrous brown to black peat	0 - 3.2

APPENDIX 5 WATER QUALITY DATA FOR
FILSON CREEK WATERSHED STUDY, 1977-1978

WATER QUALITY DATA FOR OL-2 LOCATION

DATE OF SAMPLE	TEMP	PH	HCO3	SO4	CL	TCA	DCA	TMG	DMG	TNA	DNA	TK	DK	SI	SC	TCU	INI
124	0	6.5	56.0	5.7	2.7	6.0	-0	6.2	-0	4.2	-0	.5	-0	5.4	82.0	-0	-0
409	1.0	6.1	9.0	0	1.0	3.6	-0	2.7	-0	1.6	-0	.7	-0	1.8	23.0	.9	0
423	5.5	5.4	6.0	4.0	1.7	3.8	-0	3.1	-0	2.1	-0	1.9	-0	7.1	38.0	2.8	0
508	11.0	5.4	4.0	10.8	.9	4.0	2.6	2.6	2.5	1.7	1.7	.9	.9	3.6	-0	2.3	0
527	17.5	6.1	5.0	9.0	.9	4.3	2.1	2.2	2.1	1.7	1.6	.7	.7	3.8	-0	2.6	0
611	17.5	5.8	3.0	5.0	.5	2.9	1.7	2.0	1.2	1.3	1.2	1.5	1.5	7.8	45.0	-0	-0
625	18.0	5.8	4.0	-0	.4	4.0	2.6	2.0	1.3	1.1	1.1	.3	.3	8.1	40.0	1.8	2.4
702	17.0	6.0	6.0	0	1.6	4.0	1.8	3.0	1.1	1.2	1.0	.7	.6	5.3	39.0	.8	0
723	20.5	5.9	10.0	1.0	1.4	4.8	3.6	3.5	1.6	1.2	1.0	.7	.6	6.6	60.0	2.1	4.5
806	17.0	6.0	8.0	2.0	1.2	4.4	2.0	3.5	1.2	1.2	1.1	.6	.6	5.1	37.0	5.7	3.6
820	15.3	6.1	5.0	1.0	1.0	3.5	2.9	2.1	1.9	1.2	1.2	.6	.6	4.4	32.0	-0	-0
917	15.0	5.8	3.0	1.0	1.2	3.5	3.4	2.1	1.8	1.1	1.1	.6	.6	7.3	39.0	-0	-0
1001	9.8	5.7	3.0	2.0	1.8	3.1	2.7	1.8	1.1	1.1	1.0	.7	.6	9.1	36.0	-0	-0
1015	4.5	5.8	3.0	2.0	1.3	2.4	2.4	1.8	1.8	1.2	1.1	.6	.6	9.6	34.0	-0	-0
1105	4.0	5.8	4.0	2.0	1.4	2.3	2.3	1.8	1.7	1.2	1.2	.6	.6	12.3	37.0	-0	-0
1211	0	5.5	3.0	0	1.0	3.2	2.8	1.9	1.9	1.2	1.2	.3	.3	10.8	41.0	-0	-0
128	-0	5.9	9.0	-0	1.9	3.6	3.4	1.5	1.4	1.7	1.8	.4	.4	14.7	55.0	-0	-0

CL2

WATER QUALITY DATA FOR QST LOCATION

423	.5	6.0	5.0	1.0	1.1	3.2	-0	2.3	-0	1.6	-0	1.5	-0	4.3	32.0	1.0	0
508	12.0	6.0	9.0	4.0	.9	2.8	-0	1.6	-0	1.4	-0	.7	-0	3.4	-0	1.4	0
527	18.0	6.2	6.0	2.0	.9	3.2	-0	2.0	-0	1.4	-0	.6	-0	2.9	-0	.8	0
611	18.0	6.1	5.0	2.0	.5	-0	1.7	-0	1.7	-0	1.2	-0	.7	3.6	34.0	4.5	0
625	18.0	6.3	12.0	0	.5	3.5	2.4	2.0	1.5	1.3	1.1	.6	.6	6.0	46.0	.3	0
722	21.0	6.0	9.0	0	2.4	4.0	2.0	3.3	1.5	1.1	1.1	3.1	2.9	7.0	55.0	.3	0
806	19.0	5.9	5.0	2.0	1.2	3.2	1.8	3.0	1.2	1.1	1.1	.7	.7	7.3	42.0	.6	1.1
820	16.0	5.8	7.0	3.0	1.0	3.2	2.9	2.1	2.0	1.2	1.2	.7	.7	7.1	37.0	-0	-0
917	15.5	6.1	4.0	1.0	1.2	3.4	3.4	2.2	2.3	1.3	1.3	.6	.6	7.2	34.0	-0	-0
1001	8.0	6.0	3.0	4.0	1.8	2.9	2.9	2.0	2.2	1.3	1.3	.7	.7	8.8	40.0	-0	-0
1105	4.0	5.6	3.0	2.0	1.6	2.8	2.2	2.2	2.1	1.3	1.3	.8	.8	12.1	31.0	-0	-0

QST

WATER QUALITY DATA FOR OMADAY LAKE

DATE OF SAMPLE	TEMP	PH	HCO3	SO4	CL	TCA	DCA	TMG	DMG	TNA	DNA	TK	DK	SI	SC	TCU	TNI
1204	0	6.6	19.0	5.0	1.0	8.0	-0	2.8	-0	2.2	-0	.4	-0	1.1	35.0	-0	-0
127	0	6.3	19.0	.5	1.6	7.8	-0	5.8	-0	3.4	-0	.5	-0	1.8	41.0	-0	-0
205	0	6.3	18.0	.5	1.6	7.8	-0	5.8	-0	3.2	-0	.6	-0	1.6	44.0	-0	-0
305	0	6.3	17.0	2.0	1.6	7.8	-0	5.8	-0	3.1	-0	.5	-0	2.2	42.0	9.0	-0
409	2.0	6.9	8.0	0	1.0	4.8	-0	2.7	-0	1.6	-0	.4	-0	1.0	21.0	2.4	0
423	13.2	6.7	8.0	2.0	1.1	4.0	-0	3.1	-0	2.5	-0	.7	-0	1.2	34.0	2.3	.2
508	14.0	6.3	7.0	15.0	.9	2.8	-0	2.7	-0	1.7	-0	.9	-0	1.7	-0	2.3	.3
527	21.0	6.7	7.0	4.0	.9	3.2	-0	2.4	-0	1.1	-0	.9	-0	2.4	-0	1.8	.2
611	17.5	5.8	5.0	6.0	.5	2.6	2.0	2.0	1.4	1.4	1.4	.9	.9	5.5	35.0	3.8	0
625	23.0	6.2	6.0	3.0	.5	2.6	2.0	2.0	1.2	1.2	1.2	.6	.6	8.6	35.0	3.3	0
702	18.0	6.5	6.0	1.0	1.2	2.2	2.0	1.5	1.2	1.2	1.1	.7	.7	4.2	50.0	3.0	.9
725	24.0	6.4	7.0	0	1.6	2.1	1.9	1.7	1.3	1.2	1.1	.7	.7	5.4	50.0	3.0	0
806	19.0	6.4	12.0	0	1.2	2.0	2.0	1.7	1.5	1.2	1.2	.6	.7	5.2	41.0	-0	3.1
820	-0	6.1	7.0	1.0	1.0	3.9	3.9	2.4	2.2	1.2	1.2	.7	.7	4.6	36.0	-0	-0
917	14.0	6.1	4.0	1.0	1.2	3.1	3.1	2.4	1.7	1.1	1.1	.7	.7	7.4	33.0	4.0	-0
1001	12.0	6.1	3.0	6.0	2.2	2.6	2.6	1.8	1.6	1.1	1.0	.8	.7	8.7	38.0	9.0	-0
1015	5.5	6.0	3.0	2.0	1.0	2.9	2.7	1.8	1.8	1.3	1.2	.6	.6	8.9	46.0	-0	-0
1105	4.0	6.3	3.4	2.0	1.6	2.6	2.6	2.0	1.7	1.2	1.2	.6	.6	12.2	29.0	-0	-0
1211	0	5.6	3.0	0	1.0	3.1	3.0	2.1	2.1	1.4	1.4	.3	.3	16.8	41.0	-0	-0
128	-0	6.2	8.0	-0	2.2	3.9	3.9	1.6	1.5	1.8	1.8	.4	.4	14.7	46.0	-0	-0
211	-0	5.9	10.0	-0	1.9	3.6	3.3	1.6	1.6	1.7	1.6	.4	.4	14.6	49.0	-0	-0

WATER QUALITY DATA FOR F3X LOCATION

DATE
OF
SAMPLE

	TEMP	PH	HC03	SO4	CL	TCA	DCA	TMG	DMG	TNA	DNA	TK	DK	SI	SC	TCU	TNI
1204	0	6.3	32.0	0	1.0	4.0	-0	-0	-0	2.1	-0	1.1	-0	3.0	41.0	-0	-0
409	.8	6.3	9.0	2.0	1.0	3.8	-0	3.0	-0	2.5	-0	.9	-0	1.8	21.0	3.6	.8
423	6.4	6.3	10.0	3.0	1.6	4.2	-0	5.4	-0	1.6	-0	.9	-0	10.2	40.0	8.8	3.5
508	9.0	6.2	8.0	2.0	.9	3.2	-0	2.2	-0	1.5	-0	1.1	-0	.8	-0	3.8	0
527	20.0	6.6	7.0	7.0	.9	2.8	-0	2.6	-0	1.6	-0	1.1	-0	4.2	-0	5.3	2.4
611	20.5	5.6	4.0	4.0	.5	2.9	2.1	2.0	1.0	1.2	1.3	.8	.8	6.8	40.0	4.8	3.5
625	22.0	6.1	5.0	0	.5	2.8	2.0	2.0	1.3	1.2	1.1	.6	.6	7.5	47.0	4.0	-0
702	20.3	6.2	5.0	0	1.6	2.1	2.1	1.5	1.1	1.2	1.1	.7	.7	5.7	46.0	4.3	0
723	25.0	6.2	9.0	2.0	1.8	2.3	2.3	1.5	1.2	1.2	1.3	.7	.7	5.4	46.0	4.8	8.2
806	16.5	6.3	14.0	2.0	1.2	2.0	2.0	1.8	1.3	1.2	1.2	.8	.7	5.2	39.0	1.9	3.2
820	15.0	6.3	7.0	1.0	1.0	3.1	3.1	2.3	2.1	1.3	1.1	1.0	1.0	3.5	33.0	-0	-0
917	15.1	5.9	3.0	1.0	1.2	4.0	3.6	2.0	2.1	1.0	1.1	.6	.6	6.8	33.0	-0	-0
1001	8.0	6.1	3.0	2.0	1.9	3.6	2.9	1.9	1.6	1.0	1.1	.6	.6	8.6	39.0	-0	-0
1015	5.8	5.9	2.0	2.0	1.3	2.0	1.8	1.7	1.6	1.1	1.0	.6	.6	9.6	32.0	-0	-0
1105	5.3	5.5	4.0	2.0	1.6	6.0	2.0	2.0	1.7	1.2	1.1	.6	.6	10.4	28.0	-0	-0
1211	0	-0	-0	0	1.0	3.0	2.8	2.0	2.0	1.0	1.0	.3	.3	10.1	32.0	-0	-0
128	-0	6.2	12.0	-0	1.9	3.6	3.4	1.6	1.5	1.6	1.6	.4	.4	15.0	45.0	-0	-0
211	-0	6.1	11.0	-0	2.2	4.0	4.0	1.7	1.7	1.7	1.7	.5	.5	15.2	51.0	-0	-0

WATER QUALITY DATA FOR SF LOCATION

1204	0	6.7	28.0	5.7	3.8	6.0	-0	5.5	-0	3.8	-0	2.1	-0	25.0	69.0	-0	-0
409	.5	6.4	12.0	0	1.0	4.7	-0	4.1	-0	1.2	-0	1.1	-0	7.8	27.0	3.3	0
423	3.8	6.4	16.0	0	1.6	3.3	-0	3.1	-0	2.5	-0	1.2	-0	6.8	38.0	9.0	0
508	9.5	6.3	6.0	10.0	.9	3.6	-0	3.0	-0	1.8	-0	.9	-0	4.6	-0	3.0	0
527	18.5	6.2	5.0	1.8	.9	3.2	-0	3.1	-0	1.7	-0	.4	-0	7.0	-0	3.8	0
612	15.0	6.4	7.0	0	.4	4.2	2.0	2.1	-0	1.6	-0	.4	-0	9.4	-0	4.0	1.5
613	13.0	6.6	8.0	0	.4	4.1	2.0	-0	1.2	1.6	1.1	.4	.4	8.6	55.0	4.0	0
625	19.0	6.3	6.0	1.0	.4	4.1	2.1	2.4	1.0	1.6	1.0	.3	.3	12.3	55.0	4.0	1.1
702	21.0	6.4	9.0	0	1.2	2.0	2.0	1.8	1.4	1.2	1.1	.5	.5	7.5	45.0	4.1	1.5
725	23.7	6.4	15.0	0	1.2	2.6	2.6	2.0	1.8	1.3	1.3	.6	.6	12.1	49.0	5.3	1.4
806	16.0	6.7	12.0	2.0	1.2	2.4	2.4	2.0	1.8	1.3	1.3	.6	.6	11.4	42.0	7.8	3.1
820	14.0	6.5	9.0	1.0	1.0	3.1	3.1	2.3	2.1	1.2	1.2	.4	.4	10.9	33.0	-0	-0
917	13.3	6.3	5.0	1.0	1.2	4.0	3.6	2.0	2.1	1.1	1.1	.4	.4	11.4	38.0	-0	-0
1001	9.0	6.3	3.0	0	1.8	3.6	2.9	1.9	1.6	1.1	1.0	.4	.4	12.0	33.0	-0	-0
1015	5.0	6.2	3.0	0	1.0	2.4	2.4	1.7	1.7	1.1	1.1	.5	.5	12.1	27.0	-0	-0
1105	4.9	6.2	3.0	3.0	1.6	2.2	2.0	1.7	1.6	1.2	1.2	.5	.5	12.1	28.0	-0	-0
128	-0	6.1	12.0	-0	1.9	3.9	3.9	1.7	1.7	1.5	1.5	.4	.4	16.0	42.1	-0	-0

AT
WATER QUALITY DATA FOR FIX LOCATION

DATE OF SAMPLE	TEMP	PH	HCO3	SO4	CL	TCA	DCA	TMG	DMG	TNA	DNA	TK	DK	SI	SC	TCU	TNI
1204	0	6.5	6.0	0	1.0	4.3	-0	3.2	-0	2.5	-0	.7	-0	6.0	40.0	-0	-0
409	0	6.6	12.0	0	1.0	4.5	-0	2.0	-0	1.6	-0	1.0	-0	6.2	27.0	3.6	0
423	9.5	5.5	8.0	4.0	1.5	5.3	-0	3.6	-0	2.0	-0	1.2	-0	7.6	60.0	9.0	7.5
507	9.0	5.0	4.4	13.0	.9	5.6	-0	3.8	2.6	1.8	1.8	1.0	1.0	2.2	52.0	6.5	4.6
527	17.5	6.6	5.1	10.0	.9	4.4	-0	3.8	2.6	1.8	1.8	.6	.6	1.0	42.0	6.8	4.7
611	18.5	5.8	6.0	6.0	.4	2.6	2.0	2.1	1.3	1.2	1.2	.2	.2	5.0	40.0	7.0	4.8
612	-0	-0	-0	-0	.4	2.5	-0	2.1	1.2	1.2	1.2	.2	.2	-0	40.0	-0	-0
613	-0	-0	-0	-0	.4	2.5	-0	1.6	1.2	1.2	1.2	.2	.2	-0	40.0	6.8	5.3
616	-0	6.1	7.0	6.0	.4	2.5	2.0	1.6	1.2	1.1	1.1	.2	.2	8.6	42.0	9.0	4.2
618	-0	-0	-0	-0	.4	-0	-0	-0	-0	1.3	1.3	.2	.2	-0	-0	7.3	4.3
619	-0	-0	-0	-0	.4	-0	-0	-0	-0	1.1	1.1	.2	.2	-0	-0	7.2	4.5
623	-0	6.0	8.0	0	.4	3.2	1.5	2.0	1.3	1.2	1.1	.2	.2	7.8	39.0	7.0	4.5
625	24.0	6.1	7.0	0	.4	3.2	3.2	2.0	1.1	1.2	1.2	.2	.2	11.0	41.0	6.5	4.5
702	21.0	6.3	7.0	0	.6	4.8	4.0	3.0	1.1	1.3	1.3	.6	.6	6.8	43.0	6.3	5.0
703	21.0	6.2	5.0	0	.6	4.0	2.2	3.0	1.2	1.3	1.3	.6	.6	7.1	45.0	6.3	4.5
723	24.2	6.3	12.0	0	1.0	4.8	2.7	3.6	1.2	1.4	1.4	.6	.6	7.7	50.0	5.5	-0
806	17.0	6.6	13.0	2.0	1.2	4.4	2.7	3.6	1.2	1.4	1.4	.7	.7	8.6	43.0	1.9	6.7
820	15.0	5.8	9.0	1.0	1.0	4.1	3.9	2.4	2.4	1.4	1.4	.7	.7	5.8	36.0	-0	-0
917	13.0	6.1	5.0	1.0	1.3	3.3	3.1	2.0	1.6	1.1	1.0	.5	.5	9.5	32.0	-0	-0
1001	9.2	6.1	3.0	2.0	1.8	2.7	2.7	1.7	1.6	1.1	1.0	.6	.6	10.5	37.0	-0	-0
1015	5.0	6.0	2.0	2.0	1.2	-0	1.8	-0	1.4	-0	1.0	.6	.6	11.2	24.0	-0	-0
1105	4.5	6.1	3.4	2.0	1.4	2.1	2.1	2.1	1.7	1.2	1.0	.6	.6	12.3	26.0	-0	-0
1111	0	6.3	6.0	0	1.0	3.0	3.0	1.9	1.9	1.2	1.2	.3	.3	18.6	37.0	-0	-0
128	-0	6.2	13.0	-0	1.9	3.8	3.8	1.6	1.6	1.7	1.7	.4	.4	15.5	50.0	-0	-0
211	-0	6.2	11.0	-0	2.2	4.2	4.2	1.7	1.7	1.8	1.8	.5	.5	15.3	52.0	-0	-0

WATER QUALITY DATA FOR SHALLOW INTERFLOW COLLECTORS

DATE OF SAMPLE	TEMP	PH	HCO3	SO4	CL	TCA	DCA	TMG	DMG	TNA	DNA	TK	DK	SI
SW1X 723	-0	6.3	7.0	0	1.4	4.8	2.3	1.4	.4	0	0	3.5	1.8	1.8
SW1X 917	-0	6.0	7.0	0	.7	2.6	2.6	.6	.5	0	0	4.3	0	0
SW1X 1001	-0	-0	-0	0	1.8	5.4	5.4	1.8	1.8	0	0	6.7	0	0
SWNL1 611	-0	5.6	4.0	0	.3	2.0	1.8	.8	.4	0	0	4.2	4.2	0
SWNL1 722	-0	5.8	3.6	0	1.0	2.2	2.0	.7	.2	2.3	0	2.2	2.2	2.1
SWNL1 917	-0	5.2	0	1.0	.8	2.9	2.9	.9	.6	0	0	6.5	2.4	1.4
SWNL1U105	-0	-0	-0	-0	-0	-0	8.4	-0	3.9	0	0	9.2	4.2	-0
SWNL2 625	-0	5.9	3.0	0	.3	1.2	1.2	0	0	0	0	1.7	1.7	-0
SWNL2 806	-0	6.4	7.0	0	1.2	3.4	2.0	.7	.2	2.8	0	3.1	3.1	.4
SWNL2 820	-0	-0	-0	-0	-0	5.2	5.2	.9	.9	0	0	12.3	12.3	-0
SWNL2 723	-0	-0	-0	-0	-0	6.2	3.7	.7	.2	3.7	0	1.8	1.8	-0
SWNL2 917	-0	4.0	0	0	.7	1.9	1.9	.3	.3	0	0	4.4	4.4	0
SW2X 625	-0	5.4	1.0	0	.3	.8	.8	-0	0	0	0	3.0	3.0	-0
SW2X 723	-0	6.5	2.0	0	1.0	2.2	1.1	.4	0	0	0	2.5	2.5	.9
SW2X 820	-0	-0	-0	3.3	1.0	1.2	1.2	.4	.4	0	0	20.0	20.0	0
SW2X 1001	-0	5.8	1.0	1.3	3.0	4.8	4.0	2.0	1.6	0	0	19.6	18.6	0
SW3X 625	-0	5.4	1.0	1.0	.3	-0	.7	-0	.2	0	0	-0	1.7	-0
SW3X 723	-0	-0	0	0	1.2	2.0	.7	.6	0	1.0	0	1.3	1.3	.9
SW3X 806	-0	6.4	2.0	2.0	1.6	3.4	1.5	1.1	.2	1.6	0	2.6	2.6	.3
SW3X 821	-0	-0	0	0	1.2	-0	1.8	-0	.8	0	0	4.4	4.4	2.1
SW3X 917	-0	5.8	0	0	.7	-0	4.0	.5	.3	0	0	1.7	1.7	0
SWOL 611	-0	5.8	6.2	0	.6	2.5	2.2	.7	.5	.2	0	4.1	4.1	-0
SWOL 625	-0	4.5	.6	1.0	.3	2.0	2.0	.5	.4	0	0	2.9	2.9	-0
SWOL 722	-0	5.9	3.0	1.0	1.2	3.2	1.7	.9	.2	.1	.1	2.3	2.3	2.6
SWOL 806	-0	-0	-0	-0	-0	19.6	10.7	2.0	.6	2.8	.1	4.7	4.7	1.1
SWOL 820	-0	5.7	2.0	0	1.2	1.7	1.7	.3	.3	0	0	5.8	5.4	2.1
SWOL 917	-0	5.7	1.0	0	.7	2.0	1.5	.4	.3	0	0	3.4	3.4	0
SWOL 1001	-0	-0	0	-0	-0	4.3	2.8	.9	.7	0	0	6.3	6.3	-0

WATER QUALITY DATA OF PRECIPITATION

DATE OF SAMPLE		TEMP	PH	HCO3	SO4	ACL	ICA	DCA	TMG	DMG	TNA	DNA	TK	DK	SI	SC
BSSP	806	-0	6.30	1.00	5.80	1.20	1.50	1.20	.40	0	.20	.20	1.40	.60	0	34.00
BSSP	723	-0	4.20	0	3.00	2.40	1.90	.90	.20	0	.30	.30	6.80	1.10	.90	52.00
BSSP	820	-0	-0	0	2.60	.80	.90	.80	.10	.10	10.70	10.70	1.10	1.00	.80	77.00
BSSP	508	-0	4.50	0	-0	-0	.50	.50	.10	0	.20	.20	1.30	.90	0	-0
BSSP	527	-0	5.10	0	3.20	.50	1.00	-0	.10	0	.20	.20	.30	.30	.30	-0
BSSP	625	-0	5.20	0	1.40	.30	1.30	.70	.10	0	0	0	.30	.30	0	26.00
BSSP	917	-0	4.80	0	2.50	.70	.70	.70	.10	0	0	0	.30	0	.90	18.00
BSSP	1015	-0	5.10	0	2.00	.60	.60	.60	.10	.10	.20	.20	.20	.20	0	14.00
QP	1015	-0	4.52	0	2.00	.30	.30	.30	0	0	.20	.20	.20	.20	0	49.00
QP	611	-0	4.78	0	0	.30	1.30	-0	.10	0	1.00	0	.60	.50	0	18.00
QP	625	-0	4.90	0	1.00	.30	.80	.70	.10	0	0	0	.50	.40	0	20.00
QP	722	-0	-0	0	1.00	-0	1.00	.80	.10	0	0	0	.50	.40	0	18.00
QP	723	-0	5.00	0	0	.60	1.00	.70	0	0	.10	0	.50	.40	.40	-0
QP	806	-0	5.80	0	3.00	1.20	1.50	1.30	.10	0	.60	.50	2.00	2.00	.20	12.00
QP	1001	-0	5.20	0	2.00	.30	.40	.40	0	0	.60	.60	.30	0	0	20.00
HLP	611	-0	5.20	0	.90	.30	1.00	.60	0	0	.20	.10	.50	.40	0	18.00
HLP	723	-0	-0	0	1.00	.60	1.10	.60	.10	0	.70	0	.50	.50	.50	13.00
SN77	309	-0	4.70	0	1.80	.70	.70	.60	.10	0	.20	.20	0	0	0	7.00

WATER QUALITY DATA OF WETLAND PIERSIMETERS

DATE	TIME	PH	TEMP	COND	CL	TCA	PCA	TMG	DOC	TNA	DHA	TK	DK	ST	SC	TCU	TRI
10/1	6:00	5.1	24.0	11.0	1.1	-0	-0	-0	11.4	-0	-0	-0	-0	-0	128.0	-0	-0
10/1	6:00	6.0	24.0	7.0	1.1	-0	20.0	-0	11.4	-0	2.8	-0	.3	24.4	100.0	-0	116.0
10/1	6:00	6.2	24.0	7.0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0
10/1	6:00	5.8	24.0	30.0	.8	-0	14.0	-0	9.4	-0	4.3	-0	1.7	16.0	150.0	0.5	192.0
10/1	6:00	5.6	24.0	58.0	1.3	-0	16.1	-0	7.1	-0	3.8	-0	1.3	28.0	122.0	-0	-0
10/1	6:00	5.1	24.0	54.0	.6	-0	16.0	-0	6.6	-0	2.6	-0	.5	36.0	176.0	-0	-0
10/1	6:00	5.0	24.0	37.0	1.0	-0	11.7	-0	6.1	-0	1.4	-0	.3	16.7	168.0	-0	-0
10/1	6:00	5.1	24.0	27.0	1.0	-0	10.5	-0	6.6	-0	2.1	-0	.3	13.5	-0	-0	-0
10/1	6:00	5.9	24.0	1.0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0
10/1	6:00	6.5	24.0	34.0	1.6	-0	8.7	-0	5.4	-0	4.2	-0	-0	-0	92.0	-0	-0
10/1	6:00	6.5	24.0	44.0	.9	-0	8.2	-0	4.3	-0	10.0	-0	.7	17.4	96.0	105.0	144.0
10/1	6:00	6.1	24.0	84.0	.8	-0	11.7	-0	7.7	-0	7.0	-0	1.0	11.0	115.0	-0	111.0
10/1	6:00	6.5	24.0	23.0	-0	-0	9.5	-0	6.6	-0	10.0	-0	.8	15.0	167.0	50.0	176.0
10/1	6:00	6.1	24.0	16.0	-0	-0	10.6	-0	5.8	-0	14.0	-0	1.3	17.6	106.0	-0	-0
10/1	6:00	6.4	24.0	23.0	1.2	-0	13.0	-0	6.0	-0	10.1	-0	1.7	16.0	144.0	-0	-0
10/1	6:00	5.8	24.0	1.0	2.0	-0	13.7	-0	5.2	-0	9.1	-0	.9	17.2	144.0	-0	-0
10/1	6:00	6.5	24.0	6.1	1.5	-0	13.3	-0	6.3	-0	11.6	-0	.9	16.1	142.0	-0	-0
10/1	6:00	6.1	24.0	2.0	.9	-0	13.7	-0	7.5	-0	11.1	-0	.9	18.7	214.0	-0	-0
10/1	6:00	5.9	24.0	2.5	1.7	-0	6.1	-0	7.7	-0	4.6	-0	.4	29.0	-0	-0	-0
10/1	6:00	5.9	24.0	5.0	-0	-0	8.0	-0	7.3	-0	4.3	-0	.4	21.8	115.0	-0	-0
10/1	6:00	6.7	24.0	6.0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0
10/1	6:00	6.4	24.0	2.0	.8	-0	7.4	-0	3.1	-0	1.7	-0	2.9	23.0	124.0	7.9	12.0
10/1	6:00	6.6	24.0	2.0	.3	-0	6.9	-0	3.6	-0	4.3	-0	2.7	34.0	297.0	20.7	13.1
10/1	6:00	5.7	24.0	3.4	5.0	-0	7.5	-0	3.1	-0	5.6	-0	2.1	30.0	196.0	21.3	20.1
10/1	6:00	5.7	24.0	3.4	5.0	-0	9.4	-0	4.2	-0	4.6	-0	1.9	27.2	149.0	-0	-0
10/1	6:00	5.6	24.0	40.0	1.9	-0	10.0	-0	2.0	-0	7.4	-0	2.1	14.3	97.0	-0	-0
10/1	6:00	5.6	24.0	0	2.0	-0	5.3	-0	3.1	-0	4.6	-0	3.7	32.6	77.0	-0	-0
10/1	6:00	5.0	24.0	11.0	.0	-0	1.5	-0	4.0	-0	6.4	-0	.7	15.2	69.0	5.0	4.0
10/1	6:00	5.0	24.0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	-0	1.9	4.0
10/1	6:00	6.1	24.0	-0	-0	-0	4.3	-0	4.9	-0	3.1	-0	16.4	-0	-0	-0	-0
10/1	6:00	5.9	24.0	1.0	2.1	-0	10.6	-0	6.3	-0	3.7	-0	6.7	13.4	100.0	5.9	40.8
10/1	6:00	5.4	24.0	6.0	.6	-0	10.7	-0	-0	-0	3.4	-0	.7	20.2	122.0	3.8	66.1
10/1	6:00	5.0	24.0	-0	-0	-0	5.4	-0	-0	-0	-0	-0	-0	-0	75.5	-0	7.3
10/1	6:00	5.0	24.0	8.0	1.2	-0	9.6	-0	5.2	-0	4.0	-0	13.4	29.2	240.0	-0	-0
10/1	6:00	6.0	24.0	5.0	1.2	-0	20.0	-0	6.6	-0	4.4	-0	17.8	31.0	-0	-0	-0
10/1	6:00	6.0	24.0	7.0	2.8	-0	18.0	-0	7.7	-0	3.6	-0	5.1	32.0	157.0	-0	-0
10/1	6:00	5.6	24.0	6.0	1.9	-0	15.0	-0	6.2	-0	6.1	-0	1.7	-0	-0	-0	-0

APPENDIX 6 T-TESTS BETWEEN WATER QUALITY
DATA AT FLX SAMPLING LOCATION AND OTHER SAMPLING LOCATIONS
ON FILSON CREEK

FILSON CK (CREATION DATE = 76/06/24) DATA FOR SURFACE WATER FILSON CK 77
 SAMPLE FIRST

- T - T E S T -

SAMPLE	NUMBER OF CASES	MEAN	STANDARD DEVIATION	STANDARD ERROR	* POOLED VARIANCE ESTIMATE *			* SEPARATE VARIANCE ESTIMATE *				
					F VALUE	2-TAIL PROB.	T VALUE	DEGREES OF FREEDOM	2-TAIL PROB.	T VALUE	DEGREES OF FREEDOM	2-TAIL PROB.
GROUP 1	20	1.5550	.481	.107	2.43	.119	-0.30	31	.767	-0.33	31.00	.745
GROUP 2	13	1.6000	.708	.085								
GROUP 1	24	1.4208	.351	.072	4.43	.001	-1.68	43	.100	-1.61	27.75	.119
GROUP 2	21	1.7048	.739	.161								
GROUP 1	17	1.2055	.265	.056	1.39	.572	.30	33	.767	.31	28.67	.757
GROUP 2	15	1.2692	.225	.062								
GROUP 1	25	.5120	.265	.057	2.59	.050	-1.43	44	.160	-1.49	40.78	.145
GROUP 2	21	.6103	.177	.039								
GROUP 1	22	.4980	.230	.049	2.01	.213	-2.27	33	.030	-2.48	31.80	.019
GROUP 2	15	.5154	.163	.045								
GROUP 1	21	8.7762	1.257	.929	1.35	.507	1.82	40	.076	1.82	39.13	.076
GROUP 2	21	8.1210	1.040	1.080								
GROUP 1	25	41.7000	9.283	1.723	1.20	.707	.70	40	.485	.71	39.55	.481
GROUP 2	19	39.2632	7.556	1.733								

FILE # 1100 CREATION DATE: 7/8/06/24.11 DATA FOR SURFACE RATE FILMS CK 77

T - T E S T

GROUP	TEST	MEAN	STDEV	SE	MIN	MAX	F	F	T	T	T	T
							VALUE	VALUE	VALUE	VALUE	VALUE	VALUE
							2-TAIL	2-TAIL	2-TAIL	2-TAIL	2-TAIL	2-TAIL
							PROB.	PROB.	PROB.	PROB.	PROB.	PROB.
S05	GROUP 1	6.700	0.373	0.01	6.200	7.200	1.42	0.134	-1.53	0.00	-1.53	0.00
	GROUP 2	6.800	0.312	0.00	6.200	7.200	2.61	0.017	-1.50	0.00	-1.50	0.00
S06	GROUP 1	2.750	0.110	0.00	2.500	3.000	1.06	0.311	-1.29	0.00	-1.29	0.00
	GROUP 2	2.800	0.090	0.00	2.500	3.000	1.62	0.010	-2.20	0.00	-2.20	0.00
S07	GROUP 1	3.700	0.210	0.01	3.200	4.200	4.12	0.002	-1.50	0.00	-1.51	0.00
	GROUP 2	3.900	0.180	0.00	3.200	4.200	1.41	0.157	0.50	0.00	0.50	0.00
S08	GROUP 1	2.800	0.110	0.00	2.500	3.000	2.50	0.010	-1.50	0.00	-1.50	0.00
	GROUP 2	2.800	0.100	0.00	2.500	3.000	2.50	0.010	-1.50	0.00	-1.50	0.00

DATA FOR SURFACE WATER FILSON, CK 17

T - T E S T													
TEST	N	MEAN	STANDARD DEVIATION	STANDARD ERROR	* POOLED VARIANCE ESTIMATE *			* SEPARATE VARIANCE ESTIMATE *					
					F VALUE	2-TAIL PROB.	T VALUE	DEGREES OF FREEDOM	2-TAIL PROB.	T VALUE	DEGREES OF FREEDOM	2-TAIL PROB.	
1	21	1.1111	.373	.691	1.95	.179	-.14	36	.890	-.14	35.59	.866	
2	17	1.1111	.267	.645									
3	11	2.0110	3.090	.720	9.46	.002	-.75	36	.471	-.68	21.76	.504	
4	17	2.5234	6.065	1.689									
5	17	2.3333	1.810	.770	4.13	.050	.71	37	.601	.75	30.13	.461	
6	15	1.9667	1.762	.459									
7	25	1.9667	.565	.105	1.12	.277	-2.13	41	.039	-2.14	38.57	.037	
8	18	1.2822	.647	.115									
9	22	3.2182	1.007	.215	1.04	.319	1.32	38	.196	1.32	36.71	.195	
10	18	3.3180	.939	.233									
11	15	2.5667	.647	.162	1.57	.091	.61	27	.542	.63	26.91	.535	
12	13	2.6231	.724	.201									
13	22	2.1111	.563	.171	1.27	.073	1.05	37	.308	1.02	32.11	.317	
14	17	2.1111	.513	.121									

FILE FTELL (CREATION DATE = 7/10/24.) DATA FOR SURFACE WATER FILSON CR 77
 SURFILE FIRST CONF

T - T E S T

GROUP 1 - FIRST 25 CASES
 GROUP 2 - LAST 17 CASES

VARIABLE	NUMBER OF CASES	MEAN	STANDARD DEVIATION	STANDARD ERROR	* POOLED VARIANCE ESTIMATE *					* SEPARATE VARIANCE ESTIMATE *			
					F VALUE	P-TAIL PROB.	T VALUE	DEGREES OF FREEDOM	P-TAIL PROB.	T VALUE	DEGREES OF FREEDOM	P-TAIL PROB.	
1.05	GROUP 1	24	1.5350	.481	.107	*	*	*	*	*	*	*	*
	GROUP 2	11	1.4364	.336	.102	*	*	*	*	*	*	*	*
						2.02	.256	-1.50	29	.624	-1.55	27.00	.568
1.06	GROUP 1	24	1.4218	.351	.072	*	*	*	*	*	*	*	*
	GROUP 2	17	1.5765	.674	.163	*	*	*	*	*	*	*	*
						3.68	.005	-1.96	39	.341	-1.87	22.18	.393
1.07	GROUP 1	22	1.2955	.285	.096	*	*	*	*	*	*	*	*
	GROUP 2	11	1.1727	.169	.045	*	*	*	*	*	*	*	*
						3.15	.065	1.42	31	.165	1.70	30.40	.099
1.08	GROUP 1	25	.5120	.285	.057	*	*	*	*	*	*	*	*
	GROUP 2	17	.6529	.659	.110	*	*	*	*	*	*	*	*
						2.54	.038	-1.24	40	.223	-1.14	24.53	.267
1.09	GROUP 1	22	.4500	.230	.049	*	*	*	*	*	*	*	*
	GROUP 2	11	.4555	.093	.028	*	*	*	*	*	*	*	*
						6.08	.006	-1.06	31	.051	-1.08	30.22	.037
1.10	GROUP 1	21	11.7782	4.257	.929	*	*	*	*	*	*	*	*
	GROUP 2	17	11.4411	4.562	1.107	*	*	*	*	*	*	*	*
						1.15	.759	-1.55	36	.136	-1.54	33.28	.133
1.11	GROUP 1	23	11.5011	8.263	1.723	*	*	*	*	*	*	*	*
	GROUP 2	15	11.5164	12.284	3.283	*	*	*	*	*	*	*	*
						2.21	.098	-1.15	35	.881	-1.14	20.24	.893

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VIA OUR SERVICE CENTER FILED OR 17

		* POPULUS VARIABLE ESTIMATE *		* SEPARATE VARIANCE ESTIMATE *	
DATE	AMOUNT	VALUE	PERCENT	VALUE	PERCENT
7/1	1,0500	1.32	12.57	1.32	12.57
7/2	1,0100	1.32	12.87	1.32	12.87
7/3	1,0500	1.32	12.57	1.32	12.57
7/4	1,0100	1.32	12.87	1.32	12.87
7/5	1,0500	1.32	12.57	1.32	12.57
7/6	1,0100	1.32	12.87	1.32	12.87
7/7	1,0500	1.32	12.57	1.32	12.57
7/8	1,0100	1.32	12.87	1.32	12.87
7/9	1,0500	1.32	12.57	1.32	12.57
7/10	1,0100	1.32	12.87	1.32	12.87
7/11	1,0500	1.32	12.57	1.32	12.57
7/12	1,0100	1.32	12.87	1.32	12.87
7/13	1,0500	1.32	12.57	1.32	12.57
7/14	1,0100	1.32	12.87	1.32	12.87
7/15	1,0500	1.32	12.57	1.32	12.57
7/16	1,0100	1.32	12.87	1.32	12.87
7/17	1,0500	1.32	12.57	1.32	12.57
7/18	1,0100	1.32	12.87	1.32	12.87
7/19	1,0500	1.32	12.57	1.32	12.57
7/20	1,0100	1.32	12.87	1.32	12.87
7/21	1,0500	1.32	12.57	1.32	12.57
7/22	1,0100	1.32	12.87	1.32	12.87
7/23	1,0500	1.32	12.57	1.32	12.57
7/24	1,0100	1.32	12.87	1.32	12.87
7/25	1,0500	1.32	12.57	1.32	12.57
7/26	1,0100	1.32	12.87	1.32	12.87
7/27	1,0500	1.32	12.57	1.32	12.57
7/28	1,0100	1.32	12.87	1.32	12.87
7/29	1,0500	1.32	12.57	1.32	12.57
7/30	1,0100	1.32	12.87	1.32	12.87

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